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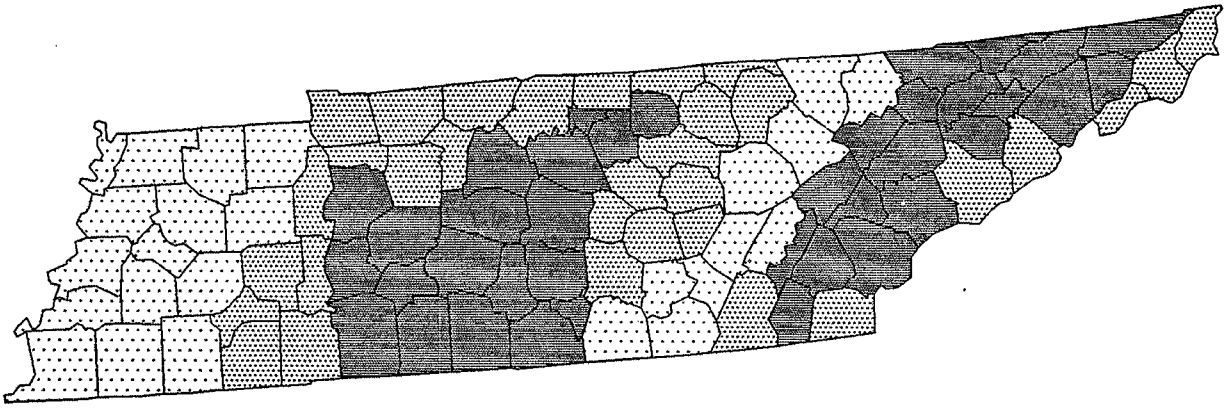
EPA's Map of Radon Zones

TENNESSEE



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EPA'S MAP OF RADON ZONES
TENNESSEE

RADON DIVISION
OFFICE OF RADIATION AND INDOOR AIR
U.S. ENVIRONMENTAL PROTECTION AGENCY

SEPTEMBER, 1993

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OVERVIEW

Sections 307 and 309 of the 1988 Indoor Radon Abatement Act (IRAA) direct EPA to identify areas of the United States that have the potential to produce elevated levels of radon. EPA, the U.S. Geological Survey (USGS), and the Association of American State Geologists (AASG) have worked closely over the past several years to produce a series of maps and documents which address these directives. The EPA Map of Radon Zones is a compilation of that work and fulfills the requirements of sections 307 and 309 of IRAA. The Map of Radon Zones identifies, on a county-by-county basis, areas of the U.S. that have the highest potential for elevated indoor radon levels (greater than 4 pCi/L).

The Map of Radon Zones is designed to assist national, State and local governments and organizations to target their radon program activities and resources. It is also intended to help building code officials determine areas that are the highest priority for adopting radon-resistant building practices. The Map of Radon Zones should not be used to determine if individual homes in any given area need to be tested for radon. **EPA recommends that all homes be tested for radon, regardless of geographic location or the zone designation of the county in which they are located.**

This document provides background information concerning the development of the Map of Radon Zones. It explains the purposes of the map, the approach for developing the map (including the respective roles of EPA and USGS), the data sources used, the conclusions and confidence levels developed for the prediction of radon potential, and the review process that was conducted to finalize this effort.

BACKGROUND

Radon (Rn^{222}) is a colorless, odorless, radioactive gas. It comes from the natural decay of uranium that is found in nearly all soils. It typically moves through the ground to the air above and into homes and other buildings through cracks and openings in the foundation. Any home, school or workplace may have a radon problem, regardless of whether it is new or old, well-sealed or drafty, or with or without a basement. Nearly one out of every 15 homes in the U.S. is estimated to have elevated annual average levels of indoor radon.

Radon first gained national attention in early 1984, when extremely high levels of indoor radon were found in areas of Pennsylvania, New Jersey, and New York, along the Reading Prong-physiographic province. EPA established a Radon Program in 1985 to assist States and homeowners in reducing their risk of lung cancer from indoor radon.

Since 1985, EPA and USGS have been working together to continually increase our understanding of radon sources and the migration dynamics that cause elevated indoor radon levels. Early efforts resulted in the 1987 map entitled "Areas with Potentially High Radon Levels." This map was based on limited geologic information only because few indoor radon measurements were available at the time. The development of EPA's Map of Radon Zones and its technical foundation, USGS' National Geologic Radon Province Map, has been based on additional information from six years of the State/EPA Residential Radon Surveys, independent State residential surveys, and continued expansion of geologic and geophysical information, particularly the data from the National Uranium Resource Evaluation project.

Purpose of the Map of Radon Zones

EPA's Map of Radon Zones (Figure 1) assigns each of the 3141 counties in the United States to one of three zones:

- Zone 1 counties have a predicted average indoor screening level > than 4 pCi/L
- Zone 2 counties have a predicted average screening level ≥ 2 pCi/L and ≤ 4 pCi/L
- Zone 3 counties have a predicted average screening level < 2 pCi/L

The Zone designations were determined by assessing five factors that are known to be important indicators of radon potential: indoor radon measurements, geology, aerial radioactivity, soil parameters, and foundation types.

The predictions of average screening levels in each of the Zones is an expression of radon potential in the lowest liveable area of a structure. This map is unable to estimate actual exposures to radon. EPA recommends methods for testing and fixing individual homes based on an estimate of actual exposure to radon. For more information on testing and fixing elevated radon levels in homes consult these EPA publications: *A Citizen's Guide to Radon, the Consumer's Guide to Radon Reduction and the Home Buyer's and Seller's Guide to Radon.*

EPA believes that States, local governments and other organizations can achieve optimal risk reductions by targeting resources and program activities to high radon potential areas. Emphasizing targeted approaches (technical assistance, information and outreach efforts, promotion of real estate mandates and policies and building codes, etc.) in such areas addresses the greatest potential risks first.

EPA also believes that the use of passive radon control systems in the construction of new homes in Zone 1 counties, and the activation of those systems if necessitated by follow-up testing, is a cost effective approach to achieving significant radon risk reduction.

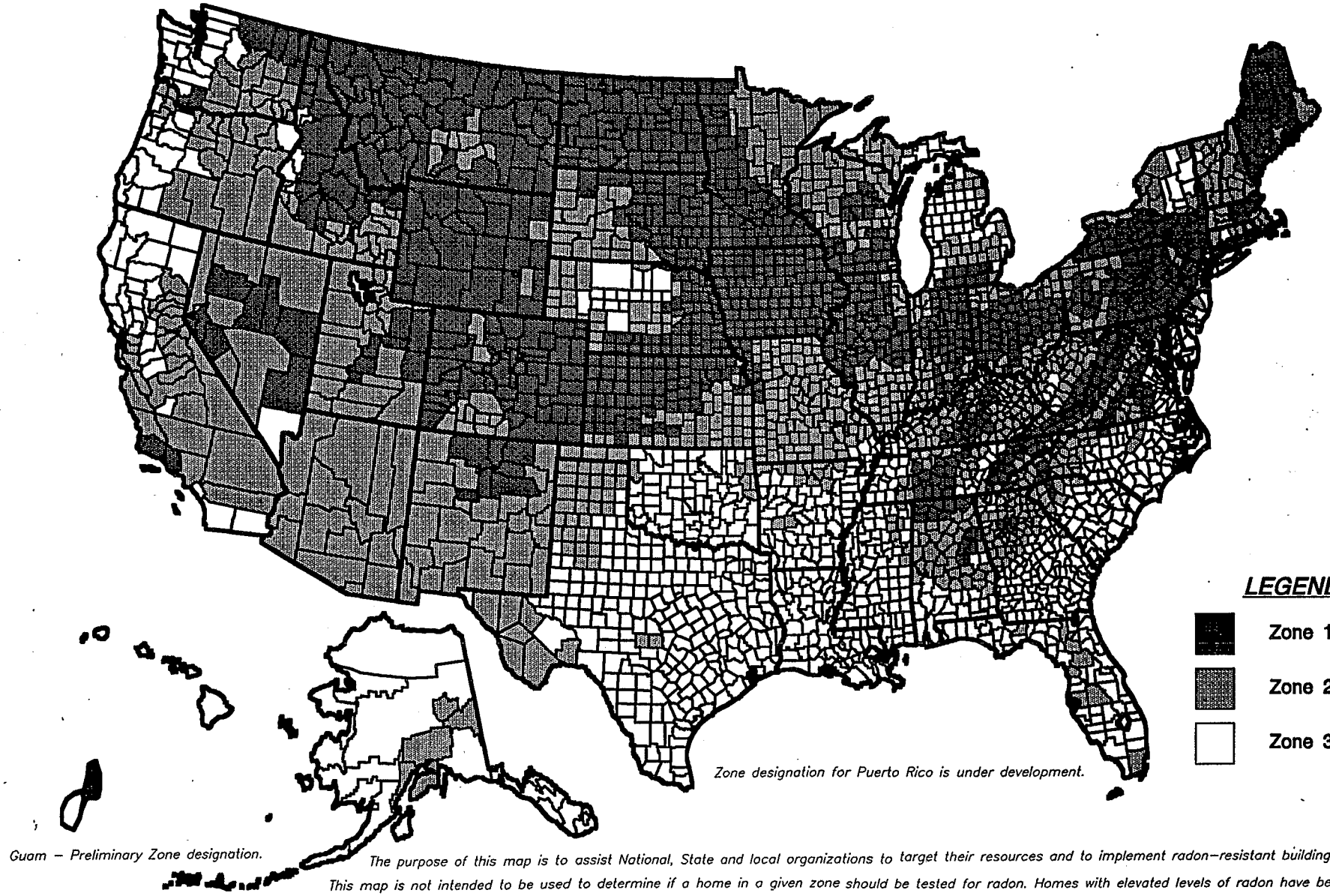
The Map of Radon Zones and its supporting documentation establish no regulatory requirements. Use of this map by State or local radon programs and building code officials is voluntary. The information presented on the Map of Radon Zones and in the supporting documentation is not applicable to radon in water.

Development of the Map of Radon Zones

The technical foundation for the Map of Radon Zones is the USGS Geologic Radon Province Map. In order to examine the radon potential for the United States, the USGS began by identifying approximately 360 separate geologic provinces for the U.S. The provinces are shown on the USGS Geologic Radon Province Map (Figure 2). Each of the geologic provinces was evaluated by examining the available data for that area: indoor radon measurements, geology, aerial radioactivity, soil parameters, and foundation types. As stated previously, these five factors are considered to be of basic importance in assessing radon

Figure 1

EPA Map of Radon Zones



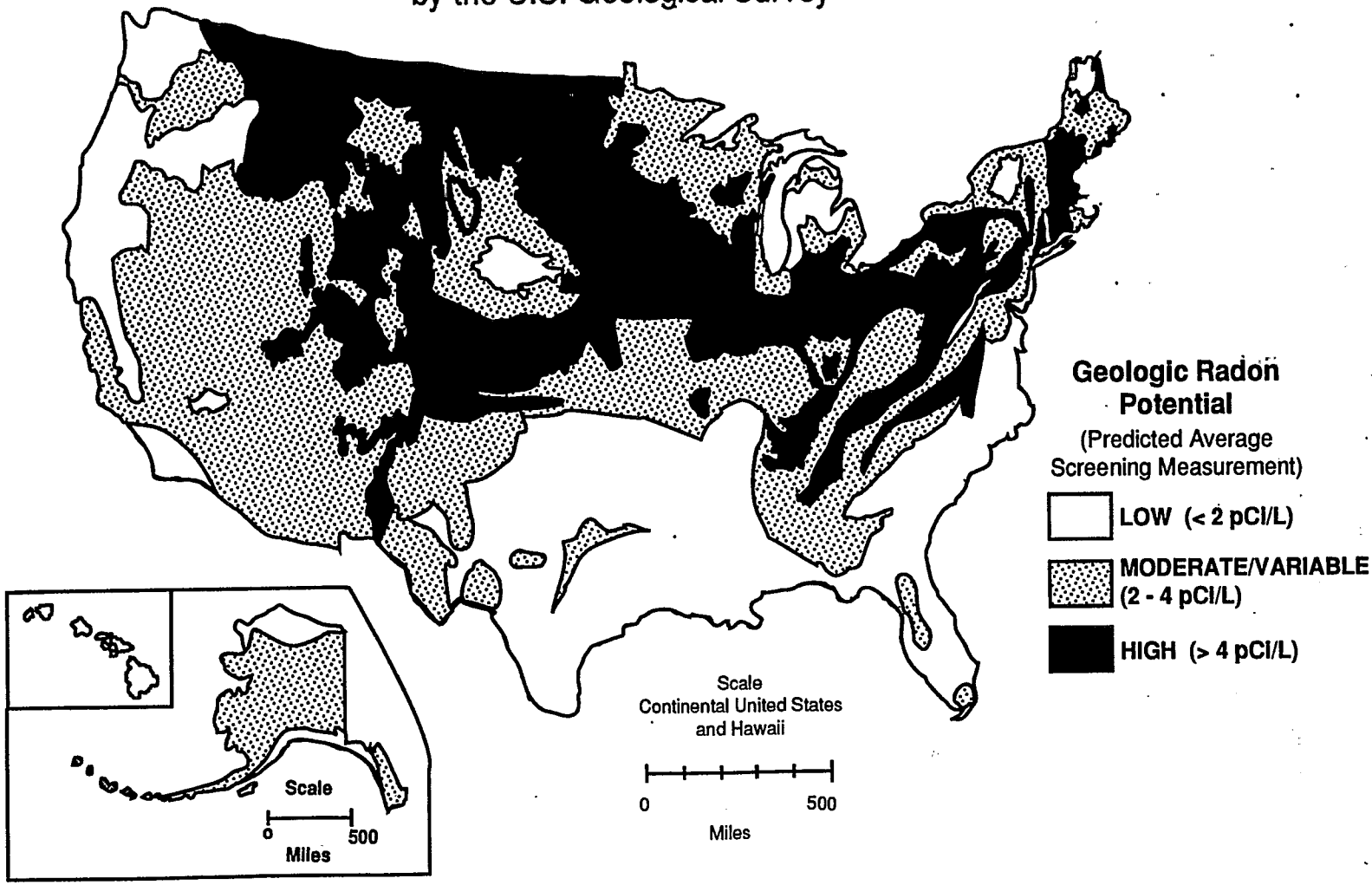
The purpose of this map is to assist National, State and local organizations to target their resources and to implement radon-resistant building codes. This map is not intended to be used to determine if a home in a given zone should be tested for radon. Homes with elevated levels of radon have been found in all three zones. ***All homes should be tested, regardless of geographic location.***

IMPORTANT: Consult the EPA Map of Radon Zones document (EPA-402-R-93-071) before using this map. This document contains information on radon potential variations within counties. EPA also recommends that this map be supplemented with any available local data in order to further understand and predict the radon potential of a specific area.

Figure 2

GENERALIZED GEOLOGIC RADON POTENTIAL OF THE UNITED STATES

by the U.S. Geological Survey



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potential and some data are available for each of these factors in every geologic province. The province boundaries do not coincide with political borders (county and state) but define areas of general radon potential. The five factors were assigned numerical values based on an assessment of their respective contribution to radon potential, and a confidence level was assigned to each contributing variable. The approach used by USGS to estimate the radon potential for each province is described in Part II of this document.

EPA subsequently developed the Map of Radon Zones by extrapolating from the province level to the county level so that all counties in the U.S. were assigned to one of three radon zones. EPA assigned each county to a given zone based on its provincial radon potential. For example, if a county is located within a geologic province that has a predicted average screening level greater than 4 pCi/L, it was assigned to Zone 1. Likewise, counties located in provinces with predicted average screening levels $\geq$ 2 pCi/L and $\leq$ 4 pCi/L, and less than 2 pCi/L, were assigned to Zones 2 and 3, respectively.

If the boundaries of a county fall in more than one geologic province, the county was assigned to a zone based on the predicted radon potential of the province in which most of the area lies. For example, if three different provinces cross through a given county, the county was assigned to the zone representing the radon potential of the province containing most of the county's land area. (In this case, it is not technically correct to say that the predicted average screening level applies to the entire county since the county falls in multiple provinces with differing radon potentials.)

Figures 3 and 4 demonstrate an example of how EPA extrapolated the county zone designations for Nebraska from the USGS geologic province map for the State. As figure 3 shows, USGS has identified 5 geologic provinces for Nebraska. Most of the counties are extrapolated "straight" from their corresponding provinces, but there are counties "partitioned" by several provinces -- for example, Lincoln County. Although Lincoln county falls in multiple provinces, it was assigned to Zone 3 because most of its area falls in the province with the lowest radon potential.

It is important to note that EPA's extrapolation from the province level to the county level may mask significant "highs" and "lows" within specific counties. In other words, within-county variations in radon potential are not shown on the Map of Radon Zones. EPA recommends that users who may need to address specific within-county variations in radon potential (e.g., local government officials considering the implementation of radon-resistant construction codes) consult USGS' Geologic Radon Province Map and the State chapters provided with this map for more detailed information, as well as any locally available data.

Map Validation

The Map of Radon Zones is intended to represent a preliminary assessment of radon potential for the entire United States. The factors that are used in this effort --indoor radon data, geology, aerial radioactivity, soils, and foundation type -- are basic indicators for radon potential. It is important to note, however, that the map's county zone designations are not "statistically valid" predictions due to the nature of the data available for these 5 factors at the county level. In order to validate the map in light of this lack of statistical confidence, EPA conducted a number of analyses. These analyses have helped EPA to identify the best situations in which to apply the map, and its limitations.

Figure 3

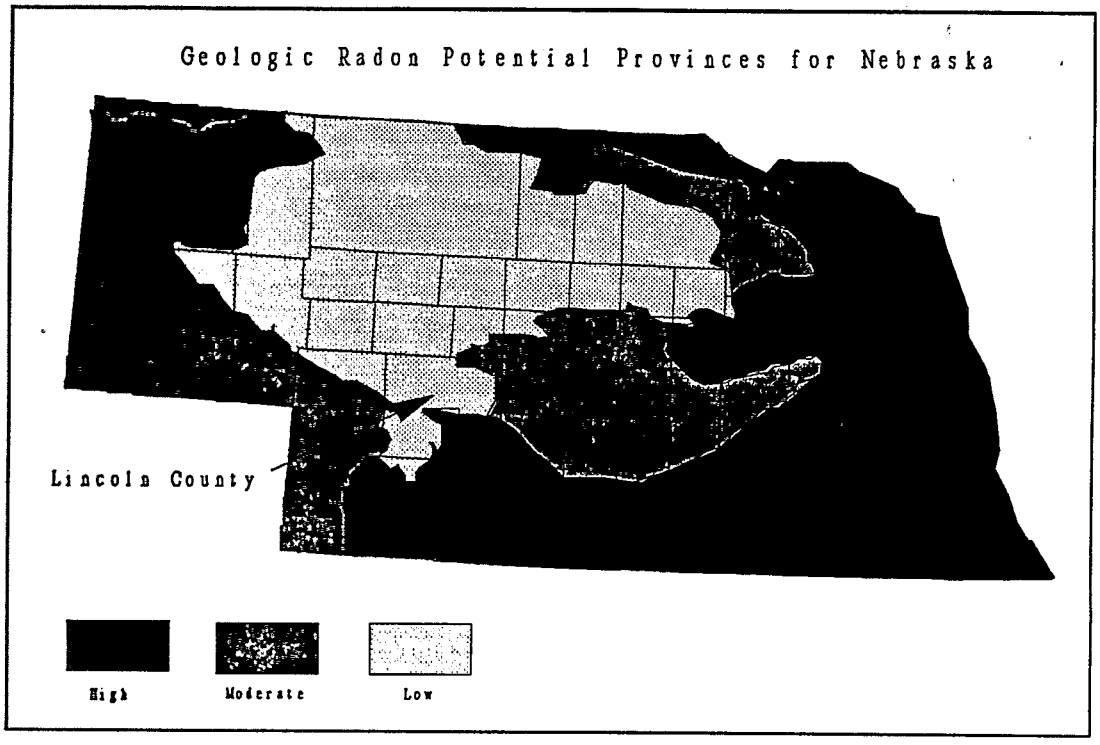
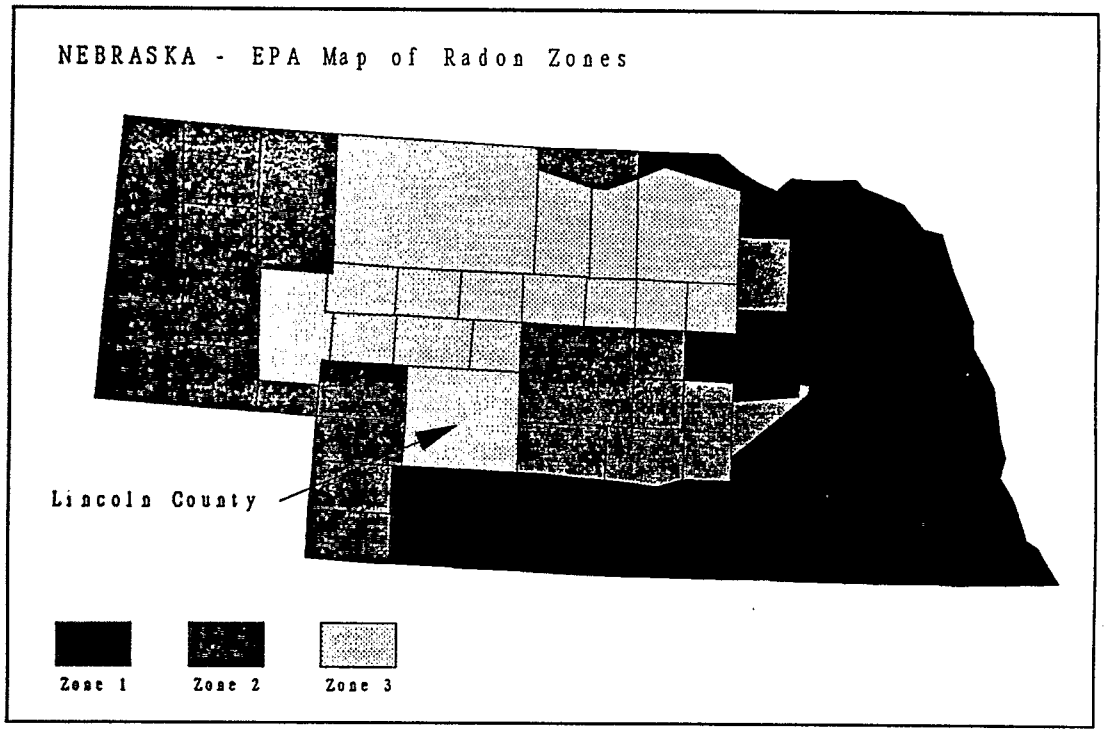


Figure 4



One such analysis involved comparing county zone designations to indoor radon measurements from the State/EPA Residential Radon Surveys (SRRS). Screening averages for counties with at least 100 measurements were compared to the counties' predicted radon potential as indicated by the Map of Radon Zones. EPA found that 72% of the county screening averages were correctly reflected by the appropriate zone designations on the Map. In all other cases, they only differed by 1 zone.

Another accuracy analysis used the annual average data from the National Residential Radon Survey (NRRS). The NRRS indicated that approximately 6 million homes in the United States have annual averages greater than or equal to 4 pCi/L. By cross checking the county location of the approximately 5,700 homes which participated in the survey, their radon measurements, and the zone designations for these counties, EPA found that approximately 3.8 million homes of the 5.4 million homes with radon levels greater than or equal to 4 pCi/L will be found in counties designated as Zone 1. A random sampling of an equal number of counties would have only found approximately 1.8 million homes greater than 4 pCi/L. In other words, this analysis indicated that the map approach is three times more efficient at identifying high radon areas than random selection of zone designations.

Together, these analyses show that the approach EPA used to develop the Map of Radon Zones is a reasonable one. In addition, the Agency's confidence is enhanced by results of the extensive State review process -- the map generally agrees with the States' knowledge of and experience in their own jurisdictions. However, the accuracy analyses highlight two important points: the fact that elevated levels will be found in Zones 2 and 3, and that there will be significant numbers of homes with lower indoor radon levels in all of the Zones. For these reasons, users of the Map of Radon Zones need to supplement the Map with locally available data whenever possible. Although all known "hot spots", i.e., localized areas of consistently elevated levels, are discussed in the State-specific chapters, accurately defining the boundaries of the "hot spots" on this scale of map is not possible at this time. Also, unknown "hot spots" do exist.

The Map of Radon Zones is intended to be a starting point for characterizing radon potential because our knowledge of radon sources and transport is always growing. Although this effort represents the best data available at this time, EPA will continue to study these parameters and others such as house construction, ventilation features and meteorology factors in order to better characterize the presence of radon in U.S homes, especially in high risk areas. These efforts will eventually assist EPA in refining and revising the conclusions of the Map of Radon Zones. And although this map is most appropriately used as a targeting tool by the aforementioned audiences -- **the Agency encourages all residents to test their homes for radon, regardless of geographic location or the zone designation of the county in which they live. Similarly, the Map of Radon Zones should not to be used in lieu of testing during real estate transactions.**

Review Process

The Map of Radon Zones has undergone extensive review within EPA and outside the Agency. The Association of American State Geologists (AASG) played an integral role in this review process. The AASG individual State geologists have reviewed their State specific information, the USGS Geologic Radon Province Map, and other materials for their geologic content and consistency.

In addition to each State geologist providing technical comments, the State radon offices were asked to comment on their respective States' radon potential evaluations. In particular, the States were asked to evaluate the data used to assign their counties to specific zones. EPA and USGS worked with the States to resolve any issues concerning county zone designations. In a few cases, States have requested changes in county zone designations. The requests were based on additional data from the State on geology, indoor radon measurements, population, etc. Upon reviewing the data submitted by the States, EPA did make some changes in zone designations. These changes, which do not strictly follow the methodology outlined in this document, are discussed in the respective State chapters.

EPA encourages the States and counties to conduct further research and data collection efforts to refine the Map of Radon Zones. EPA would like to be kept informed of any changes the States, counties, or others make to the maps. Updates and revisions will be handled in a similar fashion to the way the map was developed. States should notify EPA of any proposed changes by forwarding the changes through the Regional EPA offices that are listed in Part II. Depending on the amount of new information that is presented, EPA will consider updating this map periodically. The State radon programs should initiate proper notification of the appropriate State officials when the Map of Radon Zones is released and when revisions or updates are made by the State or EPA.

THE USGS/EPA RADON POTENTIAL ASSESSMENTS: AN INTRODUCTION

by

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BACKGROUND

The Indoor Radon Abatement Act of 1988 (15 U.S.C. 2661-2671) directed the U.S. Environmental Protection Agency (EPA) to identify areas of the United States that have the potential to produce harmful levels of indoor radon. These characterizations were to be based on both geological data and on indoor radon levels in homes and other structures. The EPA also was directed to develop model standards and techniques for new building construction that would provide adequate prevention or mitigation of radon entry. As part of an Interagency Agreement between the EPA and the U.S. Geological Survey (USGS), the USGS has prepared radon potential estimates for the United States. This report is one of ten booklets that document this effort. The purpose and intended use of these reports is to help identify areas where states can target their radon program resources, to provide guidance in selecting the most appropriate building code options for areas, and to provide general information on radon and geology for each state for federal, state, and municipal officials dealing with radon issues. *These reports are not intended to be used as a substitute for indoor radon testing, and they cannot and should not be used to estimate or predict the indoor radon concentrations of individual homes, building sites, or housing tracts. Elevated levels of indoor radon have been found in every State, and EPA recommends that all homes be tested for indoor radon.*

Booklets detailing the radon potential assessment for the U.S. have been developed for each State. USGS geologists are the authors of the geologic radon potential booklets. Each booklet consists of several components, the first being an overview to the mapping project (Part I), this introduction to the USGS assessment (Part II), including a general discussion of radon (occurrence, transport, etc.), and details concerning the types of data used. The third component is a summary chapter outlining the general geology and geologic radon potential of the EPA Region (Part III). The fourth component is an individual chapter for each state (Part IV). Each state chapter discusses the state's specific geographic setting, soils, geologic setting, geologic radon potential, indoor radon data, and a summary outlining the radon potential rankings of geologic areas in the state. A variety of maps are presented in each chapter—geologic, geographic, population, soils, aerial radioactivity, and indoor radon data by county. Finally, the booklets contain EPA's map of radon zones for each state and an accompanying description (Part V).

Because of constraints on the scales of maps presented in these reports and because the smallest units used to present the indoor radon data are counties, some generalizations have been made in order to estimate the radon potential of each area. Variations in geology, soil characteristics, climatic factors, homeowner lifestyles, and other factors that influence radon concentrations can be quite large within any particular geologic area, so these reports cannot be used to estimate or predict the indoor radon concentrations of individual homes or housing

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tracts. Within any area of a given geologic radon potential ranking, there are likely to be areas where the radon potential is lower or higher than that assigned to the area as a whole, especially in larger areas such as the large counties in some western states.

In each state chapter, references to additional reports related to radon are listed for the state, and the reader is urged to consult these reports for more detailed information. In most cases the best sources of information on radon for specific areas are state and local departments of health, state departments responsible for nuclear safety or environmental protection, and U.S. EPA regional offices. More detailed information on state or local geology may be obtained from the state geological surveys. Addresses and telephone numbers of state radon contacts, geological surveys, and EPA regional offices are listed in Appendix C at the end of this chapter.

RADON GENERATION AND TRANSPORT IN SOILS

Radon (^{222}Rn) is produced from the radioactive decay of radium (^{226}Ra), which is, in turn, a product of the decay of uranium (^{238}U) (fig. 1). The half-life of ^{222}Rn is 3.825 days. Other isotopes of radon occur naturally, but, with the exception of thoron (^{220}Rn), which occurs in concentrations high enough to be of concern in a few localized areas, they are less important in terms of indoor radon risk because of their extremely short half-lives and less common occurrence. In general, the concentration and mobility of radon in soil are dependent on several factors, the most important of which are the soil's radium content and distribution, porosity, permeability to gas movement, and moisture content. These characteristics are, in turn, determined by the soil's parent-material composition, climate, and the soil's age or maturity. If parent-material composition, climate, vegetation, age of the soil, and topography are known, the physical and chemical properties of a soil in a given area can be predicted.

As soils form, they develop distinct layers, or horizons, that are cumulatively called the soil profile. The A horizon is a surface or near-surface horizon containing a relative abundance of organic matter but dominated by mineral matter. Some soils contain an E horizon, directly below the A horizon, that is generally characterized by loss of clays, iron, or aluminum, and has a characteristically lighter color than the A horizon. The B horizon underlies the A or E horizon. Important characteristics of B horizons include accumulation of clays, iron oxides, calcium carbonate or other soluble salts, and organic matter complexes. In drier environments, a horizon may exist within or below the B horizon that is dominated by calcium carbonate, often called caliche or calcrete. This carbonate-cemented horizon is designated the K horizon in modern soil classification schemes. The C horizon underlies the B (or K) and is a zone of weathered parent material that does not exhibit characteristics of A or B horizons; that is, it is generally not a zone of leaching or accumulation. In soils formed in place from the underlying bedrock, the C horizon is a zone of unconsolidated, weathered bedrock overlying the unweathered bedrock.

The shape and orientation of soil particles (soil structure) control permeability and affect water movement in the soil. Soils with blocky or granular structure have roughly equivalent permeabilities in the horizontal and vertical directions, and air and water can infiltrate the soil relatively easily. However, in soils with platy structure, horizontal permeability is much greater than vertical permeability, and air and moisture infiltration is generally slow. Soils with prismatic or columnar structure have dominantly vertical permeability. Platy and prismatic structures form in soils with high clay contents. In soils with shrink-swell clays, air

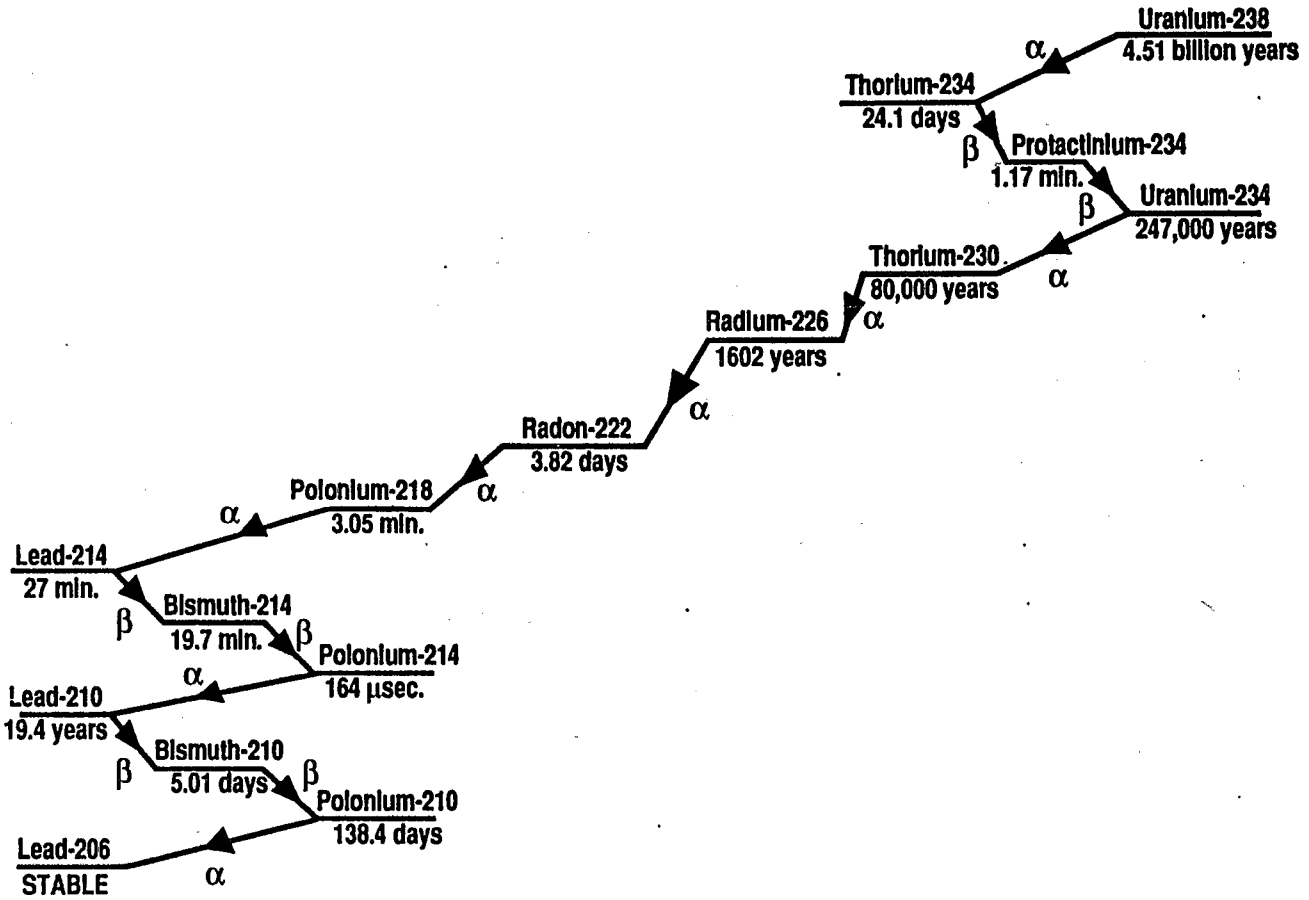

Figure 1. The uranium-238 decay series, showing the half-lives of elements and their modes of decay (after Wanty and Schoen, 1991). α denotes alpha decay, β denotes beta decay.

and moisture infiltration rates and depth of wetting may be limited when the cracks in the surface soil layers swell shut. Clay-rich B horizons, particularly those with massive or platy structure, can form a capping layer that impedes the escape of soil gas to the surface (Schumann and others, 1992). However, the shrinkage of clays can act to open or widen cracks upon drying, thus increasing the soil's permeability to gas flow during drier periods.

Radon transport in soils occurs by two processes: (1) diffusion and (2) flow (Tanner, 1964). Diffusion is the process whereby radon atoms move from areas of higher concentration to areas of lower concentration in response to a concentration gradient. Flow is the process by which soil air moves through soil pores in response to differences in pressure within the soil or between the soil and the atmosphere, carrying the radon atoms along with it. Diffusion is the dominant radon transport process in soils of low permeability, whereas flow tends to dominate in highly permeable soils (Sextro and others, 1987). In low-permeability soils, much of the radon may decay before it is able to enter a building because its transport rate is reduced. Conversely, highly permeable soils, even those that are relatively low in radium, such as those derived from some types of glacial deposits, have been associated with high indoor radon levels in Europe and in the northern United States (Akerblom and others, 1984; Kunz and others, 1989; Sextro and others, 1987). In areas of karst topography formed in carbonate rock (limestone or dolomite) environments, solution cavities and fissures can increase soil permeability at depth by providing additional pathways for gas flow.

Not all radium contained in soil grains and grain coatings will result in mobile radon when the radium decays. Depending on where the radium is distributed in the soil, many of the radon atoms may remain imbedded in the soil grain containing the parent radium atom, or become imbedded in adjacent soil grains. The portion of radium that releases radon into the pores and fractures of rocks and soils is called the emanating fraction. When a radium atom decays to radon, the energy generated is strong enough to send the radon atom a distance of about 40 nanometers (1 nm = 10^{-9} meters), or about 2×10^{-6} inches—this is known as alpha recoil (Tanner, 1980). Moisture in the soil lessens the chance of a recoiling radon atom becoming imbedded in an adjacent grain. Because water is more dense than air, a radon atom will travel a shorter distance in a water-filled pore than in an air-filled pore, thus increasing the likelihood that the radon atom will remain in the pore space. Intermediate moisture levels enhance radon emanation but do not significantly affect permeability. However, high moisture levels can significantly decrease the gas permeability of the soil and impede radon movement through the soil.

Concentrations of radon in soils are generally many times higher than those inside of buildings, ranging from tens of pCi/L to more than 100,000 pCi/L, but typically in the range of hundreds to low thousands of pCi/L. Soil-gas radon concentrations can vary in response to variations in climate and weather on hourly, daily, or seasonal time scales. Schumann and others (1992) and Rose and others (1988) recorded order-of-magnitude variations in soil-gas radon concentrations between seasons in Colorado and Pennsylvania. The most important factors appear to be (1) soil moisture conditions, which are controlled in large part by precipitation; (2) barometric pressure; and (3) temperature. Washington and Rose (1990) suggest that temperature-controlled partitioning of radon between water and gas in soil pores also has a significant influence on the amount of mobile radon in soil gas.

Homes in hilly limestone regions of the southern Appalachians were found to have higher indoor radon concentrations during the summer than in the winter. A suggested cause for this phenomenon involves temperature/pressure-driven flow of radon-laden air from subsurface

solution cavities in the carbonate rock into houses. As warm air enters solution cavities that are higher on the hillslope than the homes, it cools and settles, pushing radon-laden air from lower in the cave or cavity system into structures on the hillslope (Gammage and others, 1993). In contrast, homes built over caves having openings situated below the level of the home had higher indoor radon levels in the winter, caused by cooler outside air entering the cave, driving radon-laden air into cracks and solution cavities in the rock and soil, and ultimately, into homes (Gammage and others, 1993).

RADON ENTRY INTO BUILDINGS

A driving force (reduced atmospheric pressure in the house relative to the soil, producing a pressure gradient) and entry points must exist for radon to enter a building from the soil. The negative pressure caused by furnace combustion, ventilation devices, and the stack effect (the rising and escape of warm air from the upper floors of the building, causing a temperature and pressure gradient within the structure) during cold winter months are common driving forces. Cracks and other penetrations through building foundations, sump holes, and slab-to-foundation wall joints are common entry points.

Radon levels in the basement are generally higher than those on the main floor or upper floors of most structures. Homes with basements generally provide more entry points for radon, commonly have a more pronounced stack effect, and typically have lower air pressure relative to the surrounding soil than nonbasement homes. The term "nonbasement" applies to slab-on-grade or crawl space construction.

METHODS AND SOURCES OF DATA

The assessments of radon potential in the booklets that follow this introduction were made using five main types of data: (1) geologic (lithologic); (2) aerial radiometric; (3) soil characteristics, including soil moisture, permeability, and drainage characteristics; (4) indoor radon data; and (5) building architecture (specifically, whether homes in each area are built slab-on-grade or have a basement or crawl space). These five factors were evaluated and integrated to produce estimates of radon potential. Field measurements of soil-gas radon or soil radioactivity were not used except where such data were available in existing, published reports of local field studies. Where applicable, such field studies are described in the individual state chapters.

GEOLOGIC DATA

The types and distribution of lithologic units and other geologic features in an assessment area are of primary importance in determining radon potential. Rock types that are most likely to cause indoor radon problems include carbonaceous black shales, glauconite-bearing sandstones, certain kinds of fluvial sandstones and fluvial sediments, phosphorites, chalk, karst-producing carbonate rocks, certain kinds of glacial deposits, bauxite, uranium-rich granitic rocks, metamorphic rocks of granitic composition, silica-rich volcanic rocks, many sheared or faulted rocks, some coals, and certain kinds of contact metamorphosed rocks. Rock types least likely to cause radon problems include marine quartz sands, non-carbonaceous shales and siltstones, certain kinds of clays, silica-poor metamorphic and

igneous rocks, and basalts. Exceptions exist within these general lithologic groups because of the occurrence of localized uranium deposits, commonly of the hydrothermal type in crystalline rocks or the "roll-front" type in sedimentary rocks. Uranium and radium are commonly sited in heavy minerals, iron-oxide coatings on rock and soil grains, and organic materials in soils and sediments. Less common are ...ium associated with phosphate and carbonate complexes in rocks and soils, and uranium minerals.

Although many cases of elevated indoor radon levels can be traced to high radium and (or) uranium concentrations in parent rocks, some structural features, most notably faults and shear zones, have been identified as sites of localized uranium concentrations (Deffeyes and MacGregor, 1980) and have been associated with some of the highest reported indoor radon levels (Gundersen, 1991). The two highest known indoor radon occurrences are associated with sheared fault zones in Boyertown, Pennsylvania (Gundersen and others, 1988a; Smith and others, 1987), and in Clinton, New Jersey (Henry and others, 1991; Muessig and Bell, 1988).

NURE Aerial Radiometric Data

Aerial radiometric data are used to quantify the radioactivity of rocks and soils. Equivalent uranium (eU) data provide an estimate of the surficial concentrations of radon parent materials (uranium, radium) in rocks and soils. Equivalent uranium is calculated from the counts received by a gamma-ray detector from the 1.76 MeV (mega-electron volts) emission energy corresponding to bismuth-214 (^{214}Bi), with the assumption that uranium and its decay products are in secular equilibrium. Equivalent uranium is expressed in units of parts per million (ppm). Gamma radioactivity also may be expressed in terms of a radium activity; 3 ppm eU corresponds to approximately 1 picocurie per gram (pCi/g) of radium-226. Although radon is highly mobile in soil and its concentration is affected by meteorological conditions (Kovach, 1945; Klusman and Jaacks, 1987; Schery and others, 1984; Schumann and others, 1992), statistical correlations between average soil-gas radon concentrations and average eU values for a wide variety of soils have been documented (Gundersen and others, 1988a, 1988b; Schumann and Owen, 1988). Aerial radiometric data can provide an estimate of radon source strength over a region, but the amount of radon that is able to enter a home from the soil is dependent on several local factors, including soil structure, grain size distribution, moisture content, and permeability, as well as type of house construction and its structural condition.

The aerial radiometric data used for these characterizations were collected as part of the Department of Energy National Uranium Resource Evaluation (NURE) program of the 1970s and early 1980s. The purpose of the NURE program was to identify and describe areas in the United States having potential uranium resources (U.S. Department of Energy, 1976). The NURE aerial radiometric data were collected by aircraft in which a gamma-ray spectrometer was mounted, flying approximately 122 m (400 ft) above the ground surface. The equivalent uranium maps presented in the state chapters were generated from reprocessed NURE data in which smoothing, filtering, recalibrating, and matching of adjacent quadrangle data sets were performed to compensate for background, altitude, calibration, and other types of errors and inconsistencies in the original data set (Duval and others, 1989). The data were then gridded and contoured to produce maps of eU with a pixel size corresponding to approximately 2.5 x 2.5 km (1.6 x 1.6 mi).

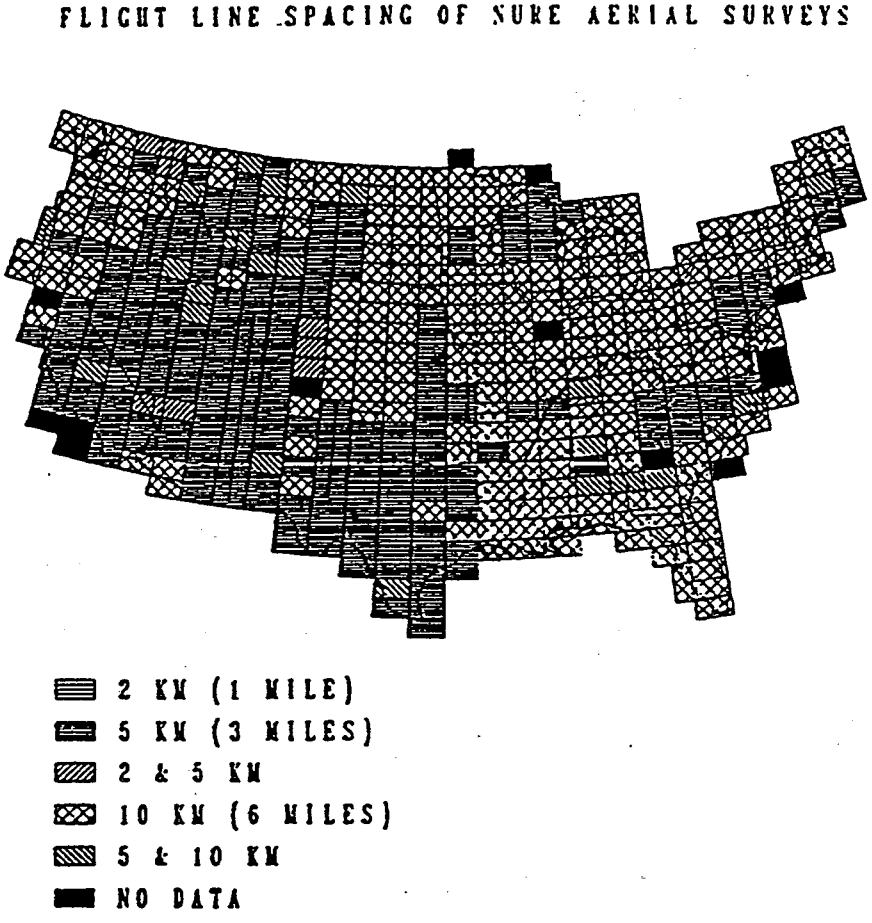

Figure 2. Nominal flightline spacings for NURE aerial gamma-ray surveys covering the contiguous United States (from Duval and others, 1990). Rectangles represent 1°x2° quadrangles.

Figure 2 is an index map of NURE 1° x 2° quadrangles showing the flight-line spacing for each quadrangle. In general, the more closely spaced the flightlines are, the more area was covered by the aerial gamma survey, and thus, more detail is available in the data set. For an altitude of 400 ft above the ground surface and with primary flightline spacing typically between 3 and 6 miles, less than 10 percent of the ground surface of the United States was actually measured by the airborne gamma-ray detectors (Duval and others, 1989), although some areas had better coverage than others due to the differences in flight-line spacing between areas (fig. 2). This suggests that some localized uranium anomalies may not have been detected by the aerial surveys, but the good correlations of eU patterns with geologic outcrop patterns indicate that, at relatively small scales (approximately 1:1,000,000 or smaller) the National eU map (Duval and others, 1989) gives reasonably good estimates of average surface uranium concentrations and thus can assist in the prediction of radon potential of rocks and soils, especially when augmented with additional geologic and soil data.

The shallow (20-30 cm) depth of investigation of gamma-ray spectrometers, either ground-based or airborne (Duval and others, 1971; Durrance, 1986), suggests that gamma-ray data may sometimes underestimate the radon-source strength in soils in which some of the radionuclides in the near-surface soil layers have been transported downward through the soil profile. In such cases the concentration of radioactive minerals in the A horizon would be lower than in the B horizon, where such minerals are typically concentrated. The concentration of radionuclides in the C horizon and below may be relatively unaffected by surface solution processes. Under these conditions the surface gamma-ray signal may indicate a lower radon source concentration than actually exists in the deeper soil layers, which are most likely to affect radon levels in structures with basements. The redistribution of radionuclides in soil profiles is dependent on a combination of climatic, geologic, and geochemical factors. There is reason to believe that correlations of eU with actual soil radium and uranium concentrations at a depth relevant to radon entry into structures may be regionally variable (Duval, 1989; Schumann and Gundersen, 1991). Given sufficient understanding of the factors cited above, these regional differences may be predictable.

Soil Survey Data

Soil surveys prepared by the U.S. Soil Conservation Service (SCS) provide data on soil characteristics, including soil-cover thickness, grain-size distribution, permeability, shrink-swell potential, vegetative cover, generalized groundwater characteristics, and land use. The reports are available in county formats and State summaries. The county reports typically contain both generalized and detailed maps of soils in the area.

Because of time and map-scale constraints, it was impractical to examine county soil reports for each county in the United States, so more generalized summaries at appropriate scales were used where available. For State or regional-scale radon characterizations, soil maps were compared to geologic maps of the area, and the soil descriptions, shrink-swell potential, drainage characteristics, depth to seasonal high water table, permeability, and other relevant characteristics of each soil group noted. Technical soil terms used in soil surveys are generally complex; however, a good summary of soil engineering terms and the national distribution of technical soil types is the "Soils" sheet of the National Atlas (U.S. Department of Agriculture, 1987).

Soil permeability is commonly expressed in SCS soil surveys in terms of the speed, in inches per hour (in/hr), at which water soaks into the soil, as measured in a soil percolation test. Although in/hr are not truly units of permeability, these units are in widespread use and are referred to as "permeability" in SCS soil surveys. The permeabilities listed in the SCS surveys are for water, but they generally correlate well with gas permeability. Because data on gas permeability of soils is extremely limited, data on permeability to water is used as a substitute except in cases in which excessive soil moisture is known to exist. Water in soil pores inhibits gas transport, so the amount of radon available to a home is effectively reduced by a high water table. Areas likely to have high water tables include river valleys, coastal areas, and some areas overlain by deposits of glacial origin (for example, loess).

Soil permeabilities greater than 6.0 in/hr may be considered high, and permeabilities less than 0.6 in/hr may be considered low in terms of soil-gas transport. Soils with low permeability may generally be considered to have a lower radon potential than more permeable soils with similar radium concentrations. Many well-developed soils contain a clay-rich B horizon that may impede vertical soil gas transport. Radon generated below this horizon cannot readily escape to the surface, so it would instead tend to move laterally, especially under the influence of a negative pressure exerted by a building.

Shrink-swell potential is an indicator of the abundance of smectitic (swelling) clays in a soil. Soils with a high shrink-swell potential may cause building foundations to crack, creating pathways for radon entry into the structure. During dry periods, desiccation cracks in shrink-swell soils provide additional pathways for soil-gas transport and effectively increase the gas permeability of the soil. Soil permeability data and soil profile data thus provide important information for regional radon assessments.

INDOOR RADON DATA

Two major sources of indoor radon data were used. The first and largest source of data is from the State/EPA Residential Radon Survey (Ronca-Battista and others, 1988; Dziuban and others, 1990). Forty-two states completed EPA-sponsored indoor radon surveys between 1986 and 1992 (fig. 3). The State/EPA Residential Radon Surveys were designed to be comprehensive and statistically significant at the state level, and were subjected to high levels of quality assurance and control. The surveys collected screening indoor radon measurements, defined as 2-7 day measurements using charcoal canister radon detectors placed in the lowest livable area of the home. The target population for the surveys included owner-occupied single family, detached housing units (White and others, 1989), although attached structures such as duplexes, townhouses, or condominiums were included in some of the surveys if they met the other criteria and had contact with the ground surface. Participants were selected randomly from telephone-directory listings. In total, approximately 60,000 homes were tested in the State/EPA surveys.

The second source of indoor radon data comes from residential surveys that have been conducted in a specific state or region of the country (e.g. independent state surveys or utility company surveys). Several states, including Delaware, Florida, Illinois, New Hampshire, New Jersey, New York, Oregon, and Utah, have conducted their own surveys of indoor radon. The quality and design of a state or other independent survey are discussed and referenced where the data are used.

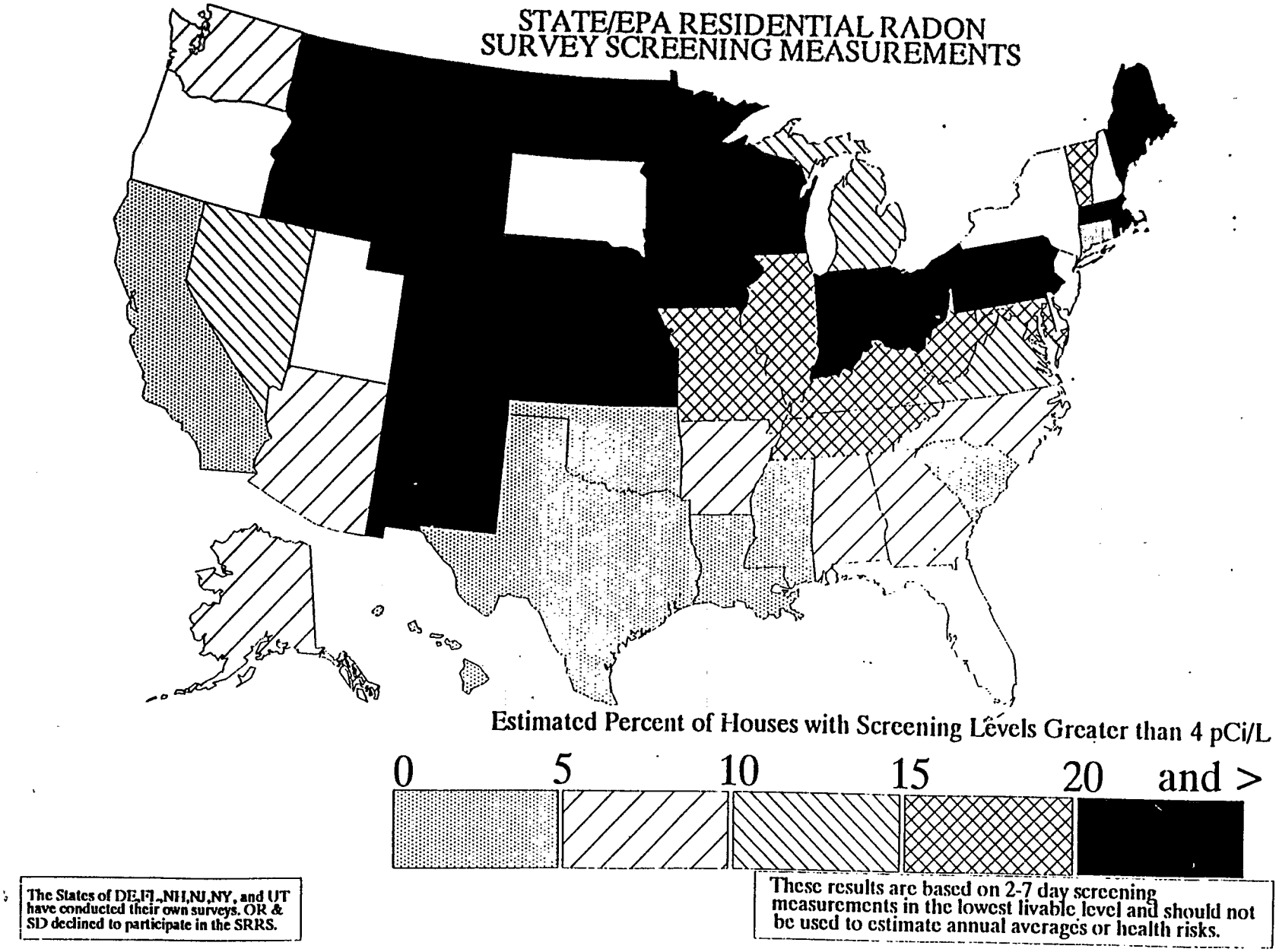

Figure 3. Percent of homes tested in the State/EPA Residential Radon Survey with screening indoor radon levels exceeding 4 pCi/L.

Data for only those counties with five or more measurements are shown in the indoor radon maps in the state chapters, although data for all counties with a nonzero number of measurements are listed in the indoor radon data tables in each state chapter. In total, indoor radon data from more than 100,000 homes nationwide were used in the compilation of these assessments. Radon data from State or regional indoor radon surveys, public health organizations, or other sources are discussed in addition to the primary data sources where they are available. Nearly all of the data used in these evaluations represent short-term (2-7 day) screening measurements from the lowest livable space of the homes. Specific details concerning the nature and use of indoor radon data sets other than the State/EPA Residential Radon Survey are discussed in the individual State chapters.

RADON INDEX AND CONFIDENCE INDEX

Many of the geologic methods used to evaluate an area for radon potential require subjective opinions based on the professional judgment and experience of the individual geologist. The evaluations are nevertheless based on established scientific principles that are universally applicable to any geographic area or geologic setting. This section describes the methods and conceptual framework used by the U.S. Geological Survey to evaluate areas for radon potential based on the five factors discussed in the previous sections. The scheme is divided into two basic parts, a Radon Index (RI), used to rank the general radon potential of the area, and the Confidence Index (CI), used to express the level of confidence in the prediction based on the quantity and quality of the data used to make the determination. This scheme works best if the areas to be evaluated are delineated by geologically-based boundaries (geologic provinces) rather than political ones (state/county boundaries) in which the geology may vary across the area.

Radon Index. Table 1 presents the Radon Index (RI) matrix. The five factors—indoor radon data, geology, aerial radioactivity, soil parameters, and house foundation type—were quantitatively ranked (using a point value of 1, 2, or 3) for their respective contribution to radon potential in a given area. At least some data for the 5 factors are consistently available for every geologic province. Because each of these main factors encompass a wide variety of complex and variable components, the geologists performing the evaluation relied heavily on their professional judgment and experience in assigning point values to each category and in determining the overall radon potential ranking. Background information on these factors is discussed in more detail in the preceding sections of this introduction.

Indoor radon was evaluated using unweighted arithmetic means of the indoor radon data for each geologic area to be assessed. Other expressions of indoor radon levels in an area also could have been used, such as weighted averages or annual averages, but these types of data were not consistently available for the entire United States at the time of this writing, or the schemes were not considered sufficient to provide a means of consistent comparison across all areas. For this report, charcoal-canister screening measurement data from the State/EPA Residential Radon Surveys and other carefully selected sources were used, as described in the preceding section. To maintain consistency, other indoor radon data sets (vendor, state, or other data) were not considered in scoring the indoor radon factor of the Radon Index if they were not randomly sampled or could not be statistically combined with the primary indoor radon data sets. However, these additional radon data sets can provide a means to further refine correlations between geologic factors and radon potential, so they are

TABLE 1. RADON INDEX MATRIX. "ppm eU" indicates parts per million of equivalent uranium, as indicated by NURE aerial radiometric data. See text discussion for details.

INCREASING RADON POTENTIAL →

FACTOR	POINT VALUE		
	1	2	3
INDOOR RADON (average)	< 2 pCi/L	2 - 4 pCi/L	> 4 pCi/L
AERIAL RADIOACTIVITY	< 1.5 ppm eU	1.5 - 2.5 ppm eU	> 2.5 ppm eU
GEOLOGY*	negative	variable	positive
SOIL PERMEABILITY	low	moderate	high
ARCHITECTURE TYPE	mostly slab	mixed	mostly basement

*GEOLOGIC FIELD EVIDENCE (GFE) POINTS: GFE points are assigned in addition to points for the "Geology" factor for specific, relevant geologic field studies. See text for details.

Geologic evidence supporting:	HIGH radon	+2 points
	MODERATE	+1 point
	LOW	-2 points
	No relevant geologic field studies	0 points

SCORING:

Radon potential category	Point range	Probable average screening indoor radon for area
LOW	3-8 points	< 2 pCi/L
MODERATE/VARIABLE	9-11 points	2 - 4 pCi/L
HIGH	12-17 points	> 4 pCi/L

POSSIBLE RANGE OF POINTS = 3 to 17

TABLE 2. CONFIDENCE INDEX MATRIX

INCREASING CONFIDENCE →

FACTOR	POINT VALUE		
	1	2	3
INDOOR RADON DATA	sparse/no data	fair coverage/quality	good coverage/quality
AERIAL RADIOACTIVITY	questionable/no data	glacial cover	no glacial cover
GEOLOGIC DATA	questionable	variable	proven geol. model
SOIL PERMEABILITY	questionable/no data	variable	reliable, abundant

SCORING:

LOW CONFIDENCE	4 - 6 points
MODERATE CONFIDENCE	7 - 9 points
HIGH CONFIDENCE	10 - 12 points

POSSIBLE RANGE OF POINTS = 4 to 12

included as supplementary information and are discussed in the individual State chapters. If the average screening indoor radon level for an area was less than 2 pCi/L, the indoor radon factor was assigned 1 point, if it was between 2 and 4 pCi/L, it was scored 2 points, and if the average screening indoor radon level for an area was greater than 4 pCi/L, the indoor radon factor was assigned 3 RI points.

Aerial radioactivity data used in this report are from the equivalent uranium map of the conterminous United States compiled from NURE aerial gamma-ray surveys (Duval and others, 1989). These data indicate the gamma radioactivity from approximately the upper 30 cm of rock and soil, expressed in units of ppm equivalent uranium. An approximate average value of eU was determined visually for each area and point values assigned based on whether the overall eU for the area falls below 1.5 ppm (1 point), between 1.5 and 2.5 ppm (2 points), or greater than 2.5 ppm (3 points).

The geology factor is complex and actually incorporates many geologic characteristics. In the matrix, "positive" and "negative" refer to the presence or absence and distribution of rock types known to have high uranium contents and to generate elevated radon in soils or indoors. Examples of "positive" rock types include granites, black shales, phosphatic rocks, and other rock types described in the preceding "geologic data" section. Examples of "negative" rock types include marine quartz sands and some clays. The term "variable" indicates that the geology within the region is variable or that the rock types in the area are known or suspected to generate elevated radon in some areas but not in others due to compositional differences, climatic effects, localizeddistribution of uranium, or other factors. Geologic information indicates not only how much uranium is present in the rocks and soils but also gives clues for predicting general radon emanation and mobility characteristics through additional factors such as structure (notably the presence of faults or shears) and geochemical characteristics (for example, a phosphate-rich sandstone will likely contain more uranium than a sandstone containing little or no phosphate because the phosphate forms chemical complexes with uranium). "Negative", "variable", and "positive" geology were assigned 1, 2, and 3 points, respectively.

In cases where additional reinforcing or contradictory geologic evidence is available, Geologic Field Evidence (GFE) points were added to or subtracted from an area's score (Table 1). Relevant geologic field studies are important to enhancing our understanding of how geologic processes affect radon distribution. In some cases, geologic models and supporting field data reinforced an already strong (high or low) score; in others, they provided important contradictory data. GFE points were applied for geologically-sound evidence that supports the prediction (but which may contradict one or more factors) on the basis of known geologic field studies in the area or in areas with geologic and climatic settings similar enough that they could be applied with full confidence. For example, areas of the Dakotas, Minnesota, and Iowa that are covered with Wisconsin-age glacial deposits exhibit a low aerial radiometric signature and score only one RI point in that category. However, data from geologic field studies in North Dakota and Minnesota (Schumann and others, 1991) suggest that eU is a poor predictor of geologic radon potential in this area because radionuclides have

been leached from the upper soil layers but are present and possibly even concentrated in deeper soil horizons, generating significant soil-gas radon. This positive supporting field evidence adds two GFE points to the score, which helps to counteract the invalid conclusion suggested by the radiometric data. No GFE points are awarded if there are no documented field studies for the area.

"Soil permeability" refers to several soil characteristics that influence radon concentration and mobility, including soil type, grain size, structure, soil moisture, drainage, slope, and permeability. In the matrix, "low" refers to permeabilities less than about 0.6 in/hr; "high" corresponds to greater than about 6.0 in/hr, in U.S. Soil Conservation Service (SCS) standard soil percolation tests. The SCS data are for water permeability, which generally correlates well with the gas permeability of the soil except when the soil moisture content is very high. Areas with consistently high water tables were thus considered to have low gas permeability. "Low, "moderate", and "high" permeability were assigned 1, 2, and 3 points, respectively.

Architecture type refers to whether homes in the area have mostly basements (3 points), mostly slab-on-grade construction (1 point), or a mixture of the two. Split-level and crawl space homes fall into the "mixed" category (2 points). Architecture information is necessary to properly interpret the indoor radon data and produce geologic radon potential categories that are consistent with screening indoor radon data.

The overall RI for an area is calculated by adding the individual RI scores for the 5 factors, plus or minus GFE points, if any. The total RI for an area falls in one of three categories—low, moderate or variable, or high. The point ranges for the three categories were determined by examining the possible combinations of points for the 5 factors and setting rules such that a majority (3 of 5 factors) would determine the final score for the low and high categories, with allowances for possible deviation from an ideal score by the other two factors. The moderate/variable category lies between these two ranges. A total deviation of 3 points from the "ideal" score was considered reasonable to allow for natural variability of factors—if two of the five factors are allowed to vary from the "ideal" for a category, they can differ by a minimum of 2 (1 point different each) and a maximum of 4 points (2 points different each). With "ideal" scores of 5, 10, and 15 points describing low, moderate, and high geologic radon potential, respectively, an ideal low score of 5 points plus 3 points for possible variability allows a maximum of 8 points in the low category. Similarly, an ideal high score of 15 points minus 3 points gives a minimum of 12 points for the high category. Note, however, that if both other factors differ by two points from the "ideal", indicating considerable variability in the system, the total point score would lie in the adjacent (i.e., moderate/variable) category.

Confidence Index. Except for architecture type, the same factors were used to establish a Confidence Index (CI) for the radon potential prediction for each area (Table 2). Architecture type was not included in the confidence index because house construction data are readily and reliably available through surveys taken by agencies and industry groups including the National Association of Home Builders, U.S. Department of Housing and Urban Development, and the Federal Housing Administration; thus it was not considered necessary

to question the quality or validity of these data. The other factors were scored on the basis of the quality and quantity of the data used to complete the RI matrix.

Indoor radon data were evaluated based on the distribution and number of data points and on whether the data were collected by random sampling (State/EPA Residential Radon Survey or other state survey data) or volunteered vendor data (likely to be nonrandom and biased toward population centers and/or high indoor radon levels). The categories listed in the CI matrix for indoor radon data ("sparse or no data", "fair coverage or quality", and "good coverage/quality") indicate the sampling density and statistical robustness of an indoor radon data set. Data from the State/EPA Residential Radon Survey and statistically valid state surveys were typically assigned 3 Confidence Index points unless the data were poorly distributed or absent in the area evaluated.

Aerial radioactivity data are available for all but a few areas of the continental United States and for part of Alaska. An evaluation of the quality of the radioactivity data was based on whether there appeared to be a good correlation between the radioactivity and the actual amount of uranium or radium available to generate mobile radon in the rocks and soils of the area evaluated. In general, the greatest problems with correlations among eU, geology, and soil-gas or indoor radon levels were associated with glacial deposits (see the discussion in a previous section) and typically were assigned a 2-point Confidence Index score. Correlations among eU, geology, and radon were generally sound in unglaciated areas and were usually assigned 3 CI points. Again, however, radioactivity data in some unglaciated areas may have been assigned fewer than 3 points, and in glaciated areas may be assigned only one point, if the data were considered questionable or if coverage was poor.

To assign Confidence Index scores for the geologic data factor, rock types and geologic settings for which a physical-chemical, process-based understanding of radon generation and mobility exists were regarded as having "proven geologic models" (3 points); a high confidence could be held for predictions in such areas. Rocks for which the processes are less well known or for which data are contradictory were regarded as "variable" (2 points), and those about which little is known or for which no apparent correlations have been found were deemed "questionable" (1 point).

The soil permeability factor was also scored based on quality and amount of data. The three categories for soil permeability in the Confidence Index are similar in concept, and scored similarly, to those for the geologic data factor. Soil permeability can be roughly estimated from grain size and drainage class if data from standard, accepted soil percolation tests are unavailable; however, the reliability of the data would be lower than if percolation test figures or other measured permeability data are available, because an estimate of this type does not encompass all the factors that affect soil permeability and thus may be inaccurate in some instances. Most published soil permeability data are for water; although this is generally closely related to the air permeability of the soil, there are some instances when it may provide an incorrect estimate. Examples of areas in which water permeability data may not accurately reflect air permeability include areas with consistently high levels of soil moisture, or clay-rich soils, which would have a low water permeability but may have a

significantly higher air permeability when dry due to shrinkage cracks in the soil. These additional factors were applied to the soil permeability factor when assigning the RI score, but may have less certainty in some cases and thus would be assigned a lower CI score.

The Radon Index and Confidence Index give a general indication of the relative contributions of the interrelated geologic factors influencing radon generation and transport in rocks and soils, and thus, of the potential for elevated indoor radon levels to occur in a particular area. However, because these reports are somewhat generalized to cover relatively large areas of States, it is highly recommended that more detailed studies be performed in local areas of interest, using the methods and general information in these booklets as a guide.

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APPENDIX A
GEOLOGIC TIME SCALE

Subdivisions (and their symbols)						Age estimates of boundaries in mega-annum (Ma) [1]
Eon or Eonothem	Era or Erathem	Period, System, Subperiod, Subsystem		Epoch or Series		
Phanerozoic [2]	Cenozoic [2] (Cz)	Quaternary [2] (Q)		Holocene		0.010
				Pleistocene		1.6 (1.6–1.9)
		Tertiary (T)	Neogene [2] Subperiod or Subsystem (N)	Pliocene		5 (4.9–5.3)
				Miocene		24 (23–26)
			Paleogene [2] Subperiod or Subsystem (Pg)	Oligocene		38 (34–38)
				Eocene		55 (54–56)
				Paleocene		66 (63–66)
	Mesozoic [2] (Mz)	Cretaceous (K)		Late	Upper	96 (95–97)
				Early	Lower	138 (135–141)
		Jurassic (J)		Late	Upper	
				Middle	Middle	
				Early	Lower	205 (200–215)
		Triassic (Tr)		Late	Upper	
				Middle	Middle	
				Early	Lower	~240
	Paleozoic [2] (Pz)	Permian (P)		Late	Upper	
				Early	Lower	290 (290–305)
		Carboniferous Systems (C)	Pennsylvanian (IP)	Late	Upper	
				Middle	Middle	
				Early	Lower	~330
			Mississippian (M)	Late	Upper	
				Early	Lower	360 (360–365)
		Devonian (D)		Late	Upper	
				Middle	Middle	
				Early	Lower	410 (405–415)
		Silurian (S)		Late	Upper	
				Middle	Middle	
				Early	Lower	435 (435–440)
		Ordovician (O)		Late	Upper	
				Middle	Middle	
				Early	Lower	500 (495–510)
		Cambrian (€)		Late	Upper	
				Middle	Middle	
				Early	Lower	~570 [3]
Proterozoic (P)	Late Proterozoic (Z)	None defined				900
	Middle Proterozoic (Y)	None defined				1600
	Early Proterozoic (X)	None defined				2500
Archean (A)	Late Archean (W)	None defined				3000
	Middle Archean (V)	None defined				3400
	Early Archean (U)	None defined				3800 ?
pre-Archean (pA) [4]						

[1] Ranges reflect uncertainties of isotopic and biostratigraphic age assignments. Age boundaries not closely bracketed by existing data shown by ~. Decay constants and isotopic ratios employed are cited in Steiger and Jäger (1977). Designation m.y. used for an interval of time.

[2] Modifiers (lower, middle, upper or early, middle, late) when used with these items are informal divisions of the larger unit; the first letter of the modifier is lowercase.

[3] Rocks older than 570 Ma also called Precambrian (p€), a time term without specific rank.

[4] Informal time term without specific rank.

APPENDIX B
GLOSSARY OF TERMS

Units of measure

pCi/L (picocuries per liter)- a unit of measure of radioactivity used to describe radon concentrations in a volume of air. One picocurie (10^{-12} curies) is equal to about 2.2 disintegrations of radon atoms per minute. A liter is about 1.06 quarts. The average concentration of radon in U.S. homes measured to date is between 1 and 2 pCi/L.

Bq/m^3 (Becquerels per cubic meter)- a metric unit of radioactivity used to describe radon concentrations in a volume of air. One becquerel is equal to one radioactive disintegration per second. One pCi/L is equal to 37 Bq/m^3.

ppm (parts per million)- a unit of measure of concentration by weight of an element in a substance, in this case, soil or rock. One ppm of uranium contained in a ton of rock corresponds to about 0.03 ounces of uranium. The average concentration of uranium in soils in the United States is between 1 and 2 ppm.

in/hr (inches per hour)- a unit of measure used by soil scientists and engineers to describe the permeability of a soil to water flowing through it. It is measured by digging a hole 1 foot (12 inches) square and one foot deep, filling it with water, and measuring the time it takes for the water to drain from the hole. The drop in height of the water level in the hole, measured in inches, is then divided by the time (in hours) to determine the permeability. Soils range in permeability from less than 0.06 in/hr to greater than 20 in/hr, but most soils in the United States have permeabilities between these two extremes.

Geologic terms and terms related to the study of radon

aerial radiometric, aeroradiometric survey A survey of radioactivity, usually gamma rays, taken by an aircraft carrying a gamma-ray spectrometer pointed at the ground surface.

alluvial fan A low, widespread mass of loose rock and soil material, shaped like an open fan and deposited by a stream at the point where it flows from a narrow mountain valley out onto a plain or broader valley. May also form at the junction with larger streams or when the gradient of the stream abruptly decreases.

alluvium, alluvial General terms referring to unconsolidated detrital material deposited by a stream or other body of running water.

alpha-track detector A passive radon measurement device consisting of a plastic film that is sensitive to alpha particles. The film is etched with acid in a laboratory after it is exposed. The etching reveals scratches, or "tracks", left by the alpha particles resulting from radon decay , which can then be counted to calculate the radon concentration. Useful for long-term (1-12 months) radon tests.

amphibolite A mafic metamorphic rock consisting mainly of pyroxenes and(or) amphibole and plagioclase.

argillite, argillaceous Terms referring to a rock derived from clay or shale, or any sedimentary rock containing an appreciable amount of clay-size material, i.e., argillaceous sandstone.

arid Term describing a climate characterized by dryness, or an evaporation rate that exceeds the amount of precipitation.

basalt A general term for a dark-colored mafic igneous rocks that may be of extrusive origin, such as volcanic basalt flows, or intrusive origin, such as basalt dikes.

batholith A mass of plutonic igneous rock that has more than 40 square miles of surface exposure and no known bottom.

carbonate A sedimentary rock consisting of the carbonate (CO_3) compounds of calcium, magnesium, or iron, e.g. limestone and dolomite.

carbonaceous Said of a rock or sediment that is rich in carbon, is coaly, or contains organic matter.

charcoal canister A passive radon measurement device consisting of a small container of granulated activated charcoal that is designed to adsorb radon. Useful for short duration (2-7 days) measurements only. May be referred to as a "screening" test.

chert A hard, extremely dense sedimentary rock consisting dominantly of interlocking crystals of quartz. Crystals are not visible to the naked eye, giving the rock a milky, dull luster. It may be white or gray but is commonly colored red, black, yellow, blue, pink, brown, or green.

clastic pertaining to a rock or sediment composed of fragments that are derived from preexisting rocks or minerals. The most common clastic sedimentary rocks are sandstone and shale.

clay A rock containing clay mineral fragments or material of any composition having a diameter less than 1/256 mm.

clay mineral One of a complex and loosely defined group of finely crystalline minerals made up of water, silicate and aluminum (and a wide variety of other elements). They are formed chiefly by alteration or weathering of primary silicate minerals. Certain clay minerals are noted for their small size and ability to absorb substantial amounts of water, causing them to swell. The change in size that occurs as these clays change between dry and wet is referred to as their **"shrink-swell"** potential.

concretion A hard, compact mass of mineral matter, normally subspherical but commonly irregular in shape; formed by precipitation from a water solution about a nucleus or center, such as a leaf, shell, bone, or fossil, within a sedimentary or fractured rock.

conglomerate A coarse-grained, clastic sedimentary rock composed of rock and mineral fragments larger than 2 mm, set in a finer-grained matrix of clastic material.

cuesta A hill or ridge with a gentle slope on one side and a steep slope on the other. The formation of a cuesta is controlled by the different weathering properties and the structural dip of the rocks forming the hill or ridge.

daughter product A nuclide formed by the disintegration of a radioactive precursor or "parent" atom.

delta, deltaic Referring to a low, flat, alluvial tract of land having a triangular or fan shape, located at or near the mouth of a river. It results from the accumulation of sediment deposited by a river at the point at which the river loses its ability to transport the sediment, commonly where a river meets a larger body of water such as a lake or ocean.

dike A tabular igneous intrusion of rock, younger than the surrounding rock, that commonly cuts across the bedding or foliation of the rock it intrudes.

diorite A plutonic igneous rock that is medium in color and contains visible dark minerals that make up less than 50% of the rock. It also contains abundant sodium plagioclase and minor quartz.

dolomite A carbonate sedimentary rock of which more than 50% consists of the mineral dolomite ($CaMg(CO_3)_2$), and is commonly white, gray, brown, yellow, or pinkish in color.

drainage The manner in which the waters of an area pass, flow off of, or flow into the soil. Also refers to the water features of an area, such as lakes and rivers, that drain it.

eolian Pertaining to sediments deposited by the wind.

esker A long, narrow, steep-sided ridge composed of irregular beds of sand and gravel deposited by streams beneath a glacier and left behind when the ice melted.

evapotranspiration Loss of water from a land area by evaporation from the soil and transpiration from plants.

extrusive Said of igneous rocks that have been erupted onto the surface of the Earth.

fault A fracture or zone of fractures in rock or sediment along which there has been movement.

fluvial, fluvial deposit Pertaining to sediment that has been deposited by a river or stream.

foliation A linear feature in a rock defined by both mineralogic and structural characteristics. It may be formed during deformation or metamorphism.

formation A mappable body of rock having similar characteristics.

glacial deposit Any sediment transported and deposited by a glacier or processes associated with glaciers, such as glaciofluvial sediments deposited by streams flowing from melting glaciers.

gneiss A rock formed by metamorphism in which bands and lenses of minerals of similar composition alternate with bands and lenses of different composition, giving the rock a striped or "foliated" appearance.

granite Broadly applied, any coarsely crystalline, quartz- and feldspar-bearing igneous plutonic rock. Technically, granites have between 10 and 50% quartz, and alkali feldspar comprises at least 65% of the total feldspar.

gravel An unconsolidated, natural accumulation of rock fragments consisting predominantly of particles greater than 2 mm in size.

heavy minerals Mineral grains in sediment or sedimentary rock having higher than average specific gravity. May form layers and lenses because of wind or water sorting by weight and size

and may be referred to as a "placer deposit." Some heavy minerals are magnetite, garnet, zircon, monazite, and xenotime.

igneous Said of a rock or mineral that solidified from molten or partly molten rock material. It is one of the three main classes into which rocks are divided, the others being sedimentary and metamorphic.

intermontane A term that refers to an area between two mountains or mountain ranges.

intrusion, intrusive The processes of emplacement or injection of molten rock into pre-existing rock. Also refers to the rock formed by intrusive processes, such as an "intrusive igneous rock".

kame A low mound, knob, hummock, or short irregular ridge formed by a glacial stream at the margin of a melting glacier; composed of bedded sand and gravel.

karst terrain A type of topography that is formed on limestone, gypsum and other rocks by dissolution of the rock by water, forming sinkholes and caves.

lignite A brownish-black coal that is intermediate in coalification between peat and subbituminous coal.

limestone A carbonate sedimentary rock consisting of more than 50% calcium carbonate, primarily in the form of the mineral calcite ($CaCO_3$).

lithology The description of rocks in hand specimen and in outcrop on the basis of color, composition, and grain size.

loam A permeable soil composed of a mixture of relatively equal parts clay, silt, and sand, and usually containing some organic matter.

loess A fine-grained eolian deposit composed of silt-sized particles generally thought to have been deposited from windblown dust of Pleistocene age.

mafic Term describing an igneous rock containing more than 50% dark-colored minerals.

marine Term describing sediments deposited in the ocean, or precipitated from ocean waters.

metamorphic Any rock derived from pre-existing rocks by mineralogical, chemical, or structural changes in response to changes in temperature, pressure, stress, and the chemical environment. Phyllite, schist, amphibolite, and gneiss are metamorphic rocks.

moraine A mound, ridge, or other distinct accumulation of unsorted, unbedded glacial material, predominantly till, deposited by the action of glacial ice.

outcrop That part of a geologic formation or structure that appears at the surface of the Earth, as in "rock outcrop".

percolation test A term used in engineering for a test to determine the water permeability of a soil. A hole is dug and filled with water and the rate of water level decline is measured.

permeability The capacity of a rock, sediment, or soil to transmit liquid or gas.

phosphate, phosphatic, phosphorite Any rock or sediment containing a significant amount of phosphate minerals, i.e., minerals containing PO_4.

physiographic province A region in which all parts are similar in geologic structure and climate, which has had a uniform geomorphic history, and whose topography or landforms differ significantly from adjacent regions.

placer deposit See heavy minerals

residual Formed by weathering of a material in place.

residuum Deposit of residual material.

rhyolite An extrusive igneous rock of volcanic origin, compositionally equivalent to granite.

sandstone A clastic sedimentary rock composed of sand-sized rock and mineral material that is more or less firmly cemented. Sand particles range from 1/16 to 2 mm in size.

schist A strongly foliated crystalline rock, formed by metamorphism, that can be readily split into thin flakes or slabs. Contains mica; minerals are typically aligned.

screening level Result of an indoor radon test taken with a charcoal canister or similar device, for a short period of time, usually less than seven days. May indicate the potential for an indoor radon problem but does not indicate annual exposure to radon.

sediment Deposits of rock and mineral particles or fragments originating from material that is transported by air, water or ice, or that accumulate by natural chemical precipitation or secretion of organisms.

semiarid Refers to a climate that has slightly more precipitation than an arid climate.

shale A fine-grained sedimentary rock formed from solidification (lithification) of clay or mud.

shear zone Refers to a roughly linear zone of rock that has been faulted by ductile or non-ductile processes in which the rock is sheared and both sides are displaced relative to one another.

shrink-swell clay See clay mineral.

siltstone A fine-grained clastic sedimentary rock composed of silt-sized rock and mineral material and more or less firmly cemented. Silt particles range from 1/16 to 1/256 mm in size.

sinkhole A roughly circular depression in a karst area measuring meters to tens of meters in diameter. It is funnel shaped and is formed by collapse of the surface material into an underlying void created by the dissolution of carbonate rock.

slope An inclined part of the earth's surface.

solution cavity A hole, channel or cave-like cavity formed by dissolution of rock.

stratigraphy The study of rock strata; also refers to the succession of rocks of a particular area.

surficial materials Unconsolidated glacial, wind-, or waterborne deposits occurring on the earth's surface.

tablelands General term for a broad, elevated region with a nearly level surface of considerable extent.

terrace gravel Gravel-sized material that caps ridges and terraces, left behind by a stream as it cuts down to a lower level.

terrain A tract or region of the Earth's surface considered as a physical feature or an ecological environment.

till Unsorted, generally unconsolidated and unbedded rock and mineral material deposited directly adjacent to and underneath a glacier, without reworking by meltwater. Size of grains varies greatly from clay to boulders.

uraniferous Containing uranium, usually more than 2 ppm.

vendor data Used in this report to refer to indoor radon data collected and measured by commercial vendors of radon measurement devices and/or services.

volcanic Pertaining to the activities, structures, and extrusive rock types of a volcano.

water table The surface forming the boundary between the zone of saturation and the zone of aeration; the top surface of a body of unconfined groundwater in rock or soil.

weathering The destructive process by which earth and rock materials, on exposure to atmospheric elements, are changed in color, texture, composition, firmness, or form with little or no transport of the material.

APPENDIX C
EPA REGIONAL OFFICES

EPA Regional Offices	State	EPA Region
EPA Region 1 JFK Federal Building Boston, MA 02203 (617) 565-4502	Alabama	4
	Alaska	10
	Arizona	9
	Arkansas	6
	California	9
EPA Region 2 (2AIR:RAD) 26 Federal Plaza New York, NY 10278 (212) 264-4110	Colorado	8
	Connecticut	1
	Delaware	3
	District of Columbia	3
	Florida	4
	Georgia	4
Region 3 (3AH14) 841 Chestnut Street Philadelphia, PA 19107 (215) 597-8326	Hawaii	9
	Idaho	10
	Illinois	5
	Indiana	5
	Iowa	7
EPA Region 4 345 Courtland Street, N.E. Atlanta, GA 30365 (404) 347-3907	Kansas	7
	Kentucky	4
	Louisiana	6
	Maine	1
	Maryland	3
EPA Region 5 (5AR26) 77 West Jackson Blvd. Chicago, IL 60604-3507 (312) 886-6175	Massachusetts	1
	Michigan	5
	Minnesota	5
	Mississippi	4
	Missouri	7
EPA Region 6 (6T-AS) 1445 Ross Avenue Dallas, TX 75202-2733 (214) 655-7224	Montana	8
	Nebraska	7
	Nevada	9
	New Hampshire	1
	New Jersey	2
EPA Region 7 726 Minnesota Avenue Kansas City, KS 66101 (913) 551-7604	New Mexico	6
	New York	2
	North Carolina	4
	North Dakota	8
	Ohio	5
EPA Region 8 (8HWM-RP) 999 18th Street One Denver Place, Suite 1300 Denver, CO 80202-2413 (303) 293-1713	Oklahoma	6
	Oregon	10
	Pennsylvania	3
	Rhode Island	1
	South Carolina	4
	South Dakota	8
	Tennessee	4
EPA Region 9 (A-3) 75 Hawthorne Street San Francisco, CA 94105 (415) 744-1048	Texas	6
	Utah	8
	Vermont	1
	Virginia	3
	Washington	10
EPA Region 10 1200 Sixth Avenue Seattle, WA 98101 (202) 442-7660	West Virginia	3
	Wisconsin	5
	Wyoming	8

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May, 1993

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EPA REGION 4 GEOLOGIC RADON POTENTIAL SUMMARY

by
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EPA Region 4 includes the states of Alabama, Florida, Georgia, Kentucky, Mississippi, North Carolina, South Carolina, and Tennessee. For each state, geologic radon potential areas were delineated and ranked on the basis of geologic, soil, housing construction, and other factors. Areas in which the average screening indoor radon level of all homes within the area is estimated to be greater than 4 pCi/L were ranked high. Areas in which the average screening indoor radon level of all homes within the area is estimated to be between 2 and 4 pCi/L were ranked moderate/variable, and areas in which the average screening indoor radon level of all homes within the area is estimated to be less than 2 pCi/L were ranked low. Information on the data used and on the radon potential ranking scheme is given in the introduction to this volume. More detailed information on the geology and radon potential of each state in Region 4 is given in the individual state chapters. The individual chapters describing the geology and radon potential of the states in EPA Region 4, though much more detailed than this summary, still are generalized assessments and there is no substitute for having a home tested. Within any radon potential area homes with indoor radon levels both above and below the predicted average will likely be found.

Major geologic/physiographic provinces for Region 4 are shown in figure 1 and are referred to in the summary that follows. The moderate climate, use of air conditioning, evaporative coolers, or open windows for ventilation, and the small proportion of homes with basements throughout much of Region 4 contribute to generally low indoor radon levels in spite of the fact that this area has substantial areas of high surface radioactivity.

Maps showing arithmetic means of measured indoor radon levels are shown in figure 2. Indoor radon data for Alabama, Georgia, Kentucky, Mississippi, North Carolina, South Carolina, and Tennessee are from the State/EPA Residential Radon Survey. Data for Florida are from the Florida Statewide Radon Study. County screening indoor radon averages range from less than 1 pCi/L to 4.6 pCi/L. The geologic radon potential areas in Region 4 have been summarized from the individual state chapters and are shown in figure 3.

ALABAMA

The Plateaus

The Interior Low Plateaus have been ranked high in geologic radon potential. The Mississippian carbonate rocks and shales that underlie this province appear to have high (>2.5 ppm eU) to moderate (1.5-2.5 ppm eU) radioactivity associated with them. The carbonates and shales are also associated with most of the highest county indoor radon averages for the State, particularly in Colbert, Madison, Lawrence, and Lauderdale Counties. The geologic units that may be the source of these problems, as indicated by the radioactivity, appear to be parts of the Fort Payne Chert, the Tuscumbia Limestone, the Monteagle, Bangor, Pride Mountain, and Parkwood Formations, and the Floyd Shale. Indoor radon levels in homes built on the St. Genevieve Limestone, Tuscumbia Limestone, and Fort Payne Chert averaged between 3.0 and 4.3 pCi/L. Soils developed from carbonate rocks are often elevated in uranium and radium. Carbonate soils are derived from the dissolution of the $CaCO_3$ that makes up the majority of the rock. When the $CaCO_3$ has been dissolved away, the soils are enriched in the remaining impurities, predominantly

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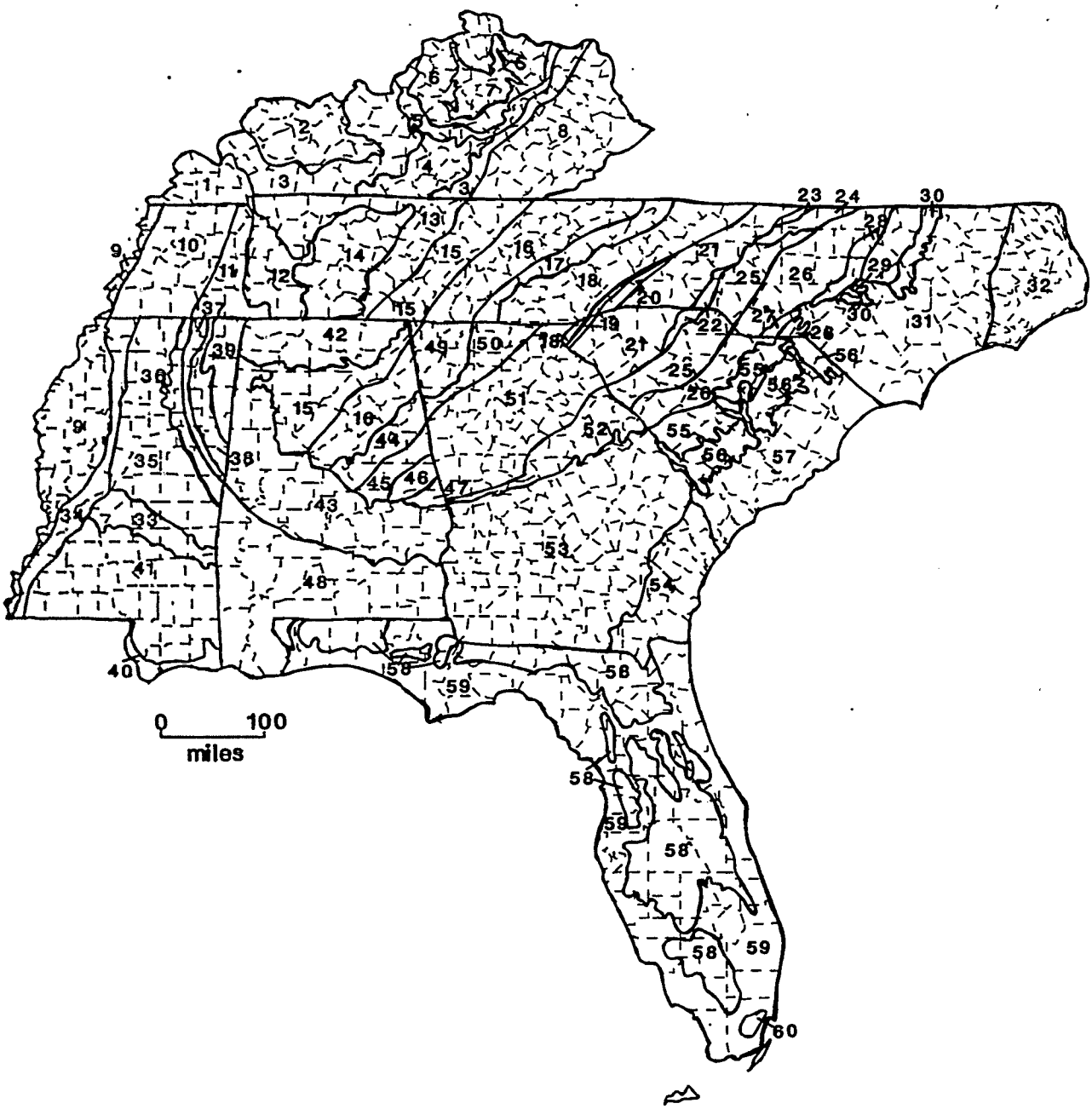

Figure 1. Geologic radon potential areas of EPA Region 4. See next page for names of numbered areas.

Figure 1 (continued). Geologic radon potential areas of EPA Region 4. Note: although some areas, for example, the Coastal Plain, are contiguous from state to state, they are sometimes referred to by slightly different names or are subdivided differently in different states, thus are numbered and labelled seperately on this figure.

1–Jackson Purchase (Coastal Plain)
2–Western Coalfield
3–Mississippian Plateau
4–Eastern Pennyroyal
5–New Albany Shale
6–Outer Bluegrass
7–Inner Bluegrass
8–Cumberland Plateau (Appalachian Plateau)
9–Mississippi alluvial plain
10–Loess-covered Coastal Plain
11–Eastern Coastal Plain
12–Cherty Highland
13–Highland Rim
14–Nashville Basin
15–Appalachian Plateau
16–Ridge and Valley
17–Unaka Mountains
18–Blue Ridge Belt
19–Brevard Fault Zone
20–Chauga Belt
21–Inner Piedmont
22–Kings Mountain Belt
23–Dan River Basin
24–Milton Belt
25–Charlotte Belt
26–Carolina Slate Belt
27–Wadesboro sub-basin
28–Sanford-Durham sub-basins
29–Raleigh Belt
30–Eastern Slate Belt
31–Inner Coastal Plain
32–Outer Coastal Plain
33–Jackson Prairies
34–Loess Hills
35–North Central Hills
36–Flatwoods
37–Pontotoc Ridge
38–Black Prairies
39–Tombigbee Hills
40–Coastal Pine Meadows
41–Pine Hills
42–Interior Low Plateaus
43–Inner Coastal Plain (Cretaceous)
44–Northern Piedmont (faults, phylite and granite rocks)
45–Wedowee and Emuckfaw Groups
46–Inner Piedmont/Dadeville Complex
47–Southern Piedmont
48–Inner and Outer Coastal Plain (Tertiary Rocks)
49–Rome-Kingston Thrust Stack
50–Georgiabama Thrust Stack (north of Allatoona Fault)
51–Georgiabama Thrust Stack (south of Allatoona Fault)
52–Little River Thrust Stack
53–Coastal Plain (Cretaceous/Tertiary)
54–Coastal Plain (Quaternary/Pliocene-Pleistocene gravels)
55–Upper Coastal Plain
56–Middle Coastal Plain
57–Lower Coastal Plain
58–Highlands
59–Lowlands
60–Dade County anomalous area.

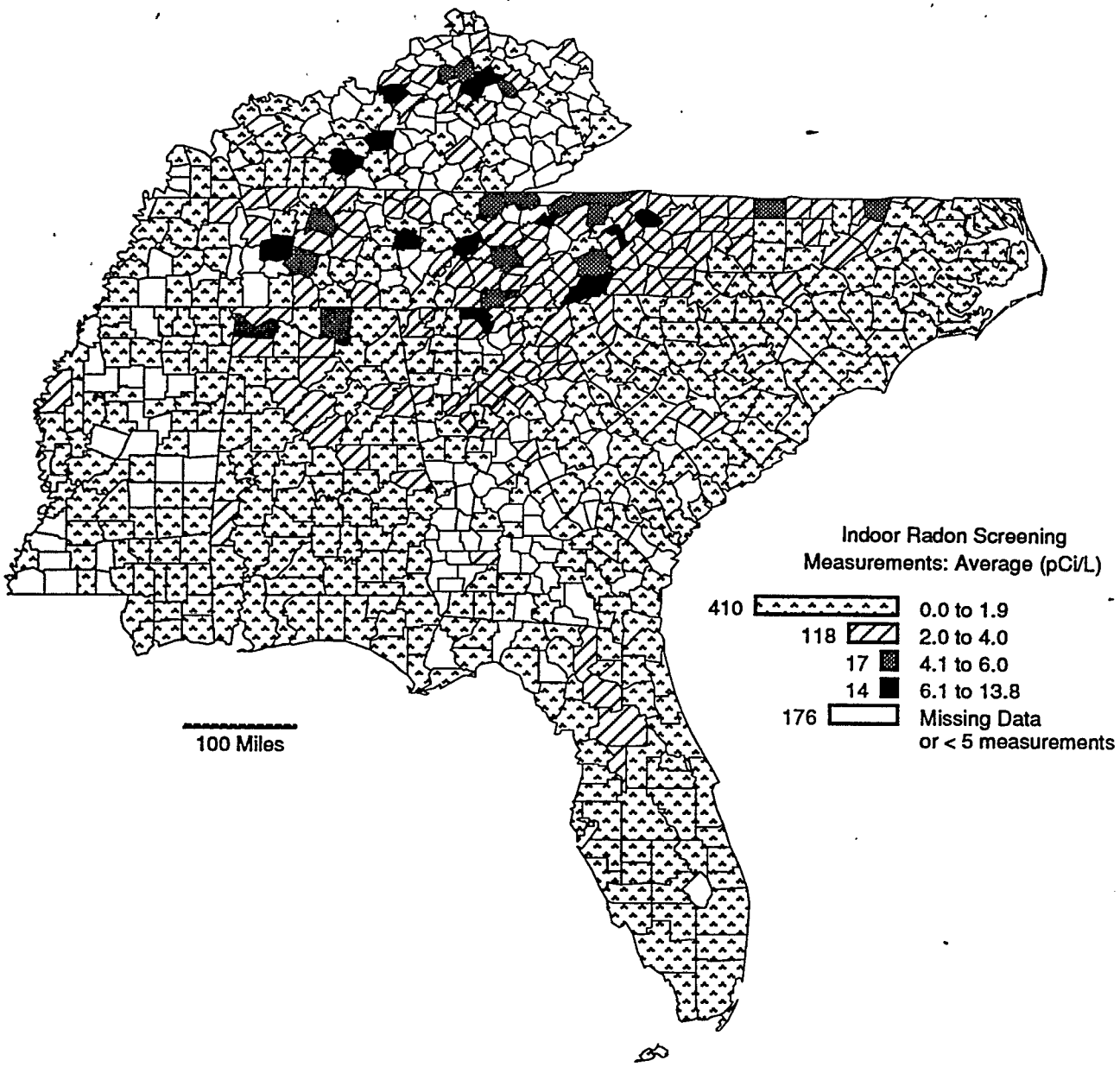

Figure 2. Screening indoor radon averages for counties with 5 or more measurements in EPA Region 4. Data for all states in Region4 except Florida from the State/EPA Residential Radon Survey. Data for Florida are from the Florida Statewide Radon Study. Histograms in map legend show the number of counties in each category.

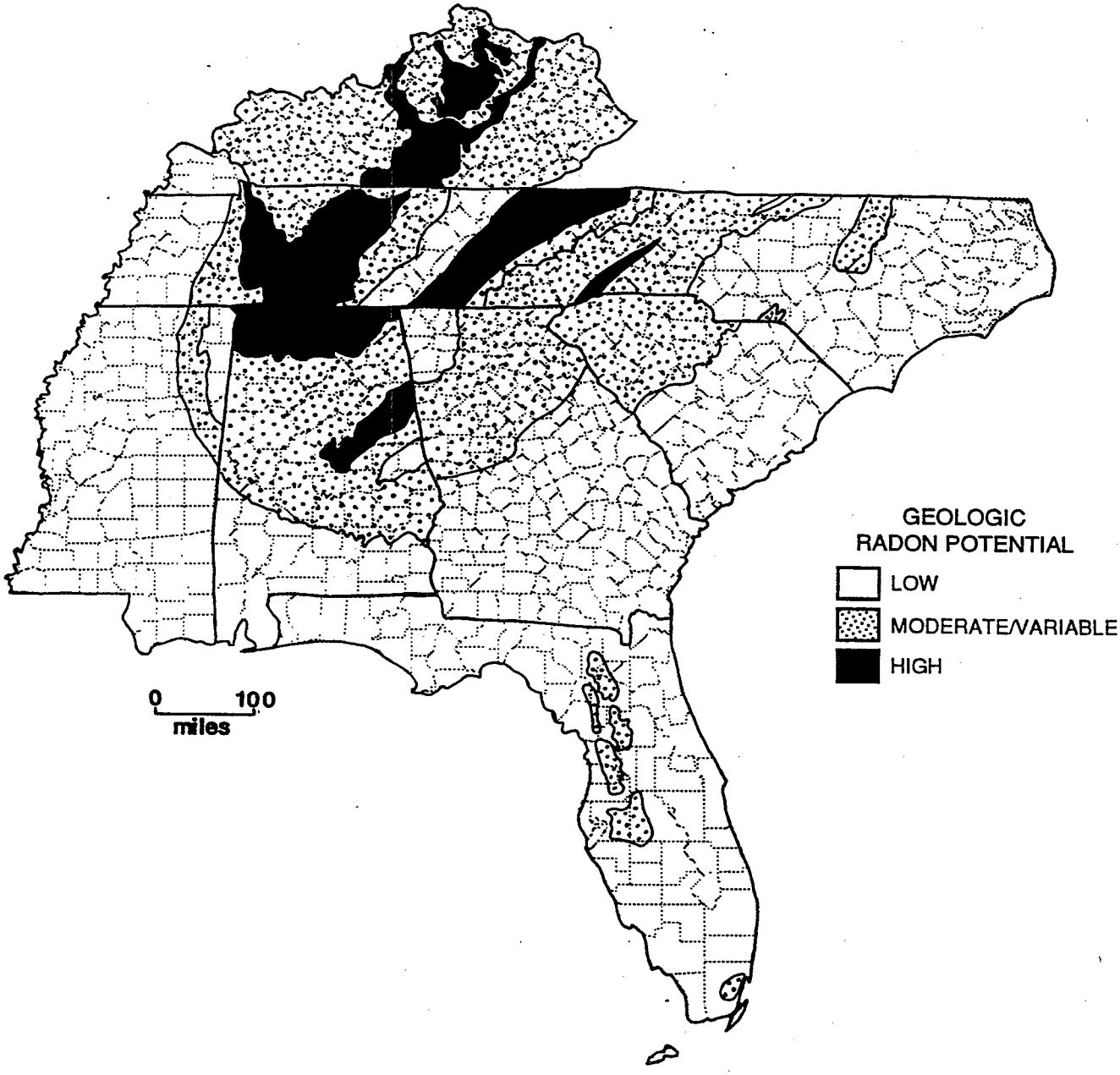

Figure 3. Geologic radon potential areas of EPA Region 4. For more detail, refer to individual state radon potential chapters.

base metals, including uranium. Rinds containing high concentrations of uranium and uranium minerals can be formed on the surfaces of rocks affected by $CaCO_3$ dissolution and karstification. Karst and cave morphology is also thought to promote the flow and accumulation of radon. Because carbonate soils are clayey, they have a tendency to crack when they dry and may develop very high permeability from the fractures. Under moist conditions, however, the soils derived from carbonates have generally low permeability.

The Appalachian Plateaus region is ranked moderate in radon potential. Indoor radon is generally low (< 2 pCi/L) to moderate (2-4 pCi/L). Radioactivity is low to moderate and soil permeability is moderate. The sandstone of the Pottsville Formation is not noted for being uranium-bearing, but uraniferous carbonaceous shales interbedded with the sandstone may be the cause of locally moderate to high (>4 pCi/L) indoor radon. Cullman County had several indoor radon measurements greater than 4 pCi/L, including one measurement of 19.8 pCi/L. Winston and Walker Counties also had several indoor radon levels greater than 4 pCi/L in the Alabama Department of Public Health data set.

Valley and Ridge

The Valley and Ridge province has been ranked moderate in geologic radon potential. Radioactivity is generally moderate in the Valley and Ridge, with high radioactivity occurring along the southeastern border with the Piedmont. Indoor radon is highly variable, with generally low county averages and one high county average. Most of the counties had a few readings greater than 4 pCi/L. The soils of the Valley and Ridge have low to moderate permeability. The permeability may be locally high in dry clayey soils and karst areas. Carbonate soils derived from Cambrian-Ordovician rock units of the Valley and Ridge province cause known indoor radon problems in eastern Tennessee, western New Jersey, western Virginia, eastern West Virginia and central and eastern Pennsylvania. Further, the Devonian Chattanooga Shale crops out locally in parts of the Valley and Ridge. This shale is widely known to be highly uraniferous and has been identified as a source of high indoor radon in Kentucky.

Piedmont

Where it is possible to associate high radioactivity and/or high indoor radon levels with particular areas, parts of the Piedmont have been ranked moderate to high in radon potential. Radiometric anomalies occur over the Talladega Fault zone, which separates the Paleozoic carbonates from the metamorphic rocks. Some of the metamorphic rocks in the Northern Piedmont, including the Poe Bridge Mountain Group, the Mad Indian Group, parts of the Wedowee Group, and the Higgins Ferry Group, also have high radioactivity associated with them. In many cases the radiometric anomalies appear to be associated with rocks in fault zones, graphitic schists and phyllites, felsic gneiss, and other granitic rocks. Furthermore, Talladega, Calhoun, Cleburne, and Randolph Counties all have some high indoor radon measurements. Uranium in graphitic phyllite with an assay value of 0.076 percent U_3O_8 has been reported from Cleburne County and similar graphitic phyllites from the Georgia Piedmont average 4.7 ppm uranium. Graphitic phyllites and schists in other parts of the Piedmont are known sources of radon and have high indoor radon levels associated with them . Another source of uranium in Piedmont metamorphic rocks is monazite, which contains high amounts of both uranium and thorium. It is a common accessory mineral in gneisses and granites throughout the Piedmont and its resistance to weathering and high density result in local monazite concentrations in saprolite. A uraniferous monazite belt that crosses the Piedmont in northern Chambers and Tallapoosa County may provide

a source of radon. Soils of the Northern and Southern Piedmont have moderate to high permeability, whereas soils developed from mafic rocks of the Dadeville Complex have low permeability. Because the Dadeville Complex consists primarily of mafic rocks with low radioactivity and low permeability, the Dadeville Complex was ranked separately from other Piedmont rocks and is ranked low in geologic radon potential.

Coastal Plain

More than half of Alabama is covered by the sediments of the Coastal Plain. Indoor radon levels are generally less than 4 pCi/L and commonly less than 2 pCi/L in this province. Soil permeability is variable–generally low in clays and moderate to high in silts and sands. A distinct radiometric high is located over the central belt of marly sandy clay and chalk known as the Selma Group. Within the Selma Group high radioactivity is associated with the Demopolis Chalk, Mooreville Chalk, Prairie Bluffs Chalk, and the Ripley Formation in central and western Alabama. In eastern Alabama and into Georgia these rocks are dominated by the glauconitic sands and clays of the Providence Sand, Cusseta Sand, and Blufftown Formation. These units have overall moderate geologic radon potential.

As part of a study by the U.S. Geological Survey and the U.S. EPA to assess the radon potential of the Coastal Plain sediments in the United States, data on radon in soil gas, surface gamma-ray activity, and soil permeability were collected and examined. Data were collected in the Alabama Coastal Plain along a transect running from just north of Montgomery, Alabama, to just south of De Funiak Springs, Florida. The highest soil-gas radon concentrations and equivalent uranium were found in the Cretaceous Mooreville Chalk, carbonaceous sands and clays of the Providence Sand, and the glauconitic sands of the Eutaw and Ripley Formations. However, permeability in many of these units is slow—generally less than $1x10^{-12}$ cm^2, and soil-gas radon was difficult to collect. Geologic units that have the lowest soil-gas radon concentrations and eU include the quartz sands of the Cretaceous Gordo Formation and quartz sands and residuum of the undifferentiated upper Tertiary sediments. Low to moderate radon and uranium concentrations were measured in the glauconitic sands and clays of the Tertiary Porters Creek Formation and in the glauconitic sands, limestones, and clays of the Tertiary Nanafalia, Lisbon Formation, and the Tuscahoma Sand. The indoor radon in some counties underlain by the Selma Group is in the 2-4 pCi/L range with a few measurements greater than 4 pCi/L, higher than in most other parts of the Alabama Coastal Plain. High uranium and radon concentrations in the sediments of the Jackson Group, locally exceeding 8 ppm U, but generally in the 1-4 ppm U range, and high soil-gas radon concentrations, are associated with faults and oil and gas wells in Choctaw County. Indoor radon measurements are generally low in these areas, but may be locally high.

FLORIDA

Florida lies entirely within the Coastal Plain, but there are six distinctive areas in Florida for which geologic radon potential may be evaluated—the Northern Highlands, Central Highlands, the Central and Northern Highlands anomalous areas, the Gulf Coastal Lowlands, Atlantic Coastal Lowlands, and an area here termed the Dade County anomalous area.

The Northern Highlands province has generally low geologic radon potential. All counties entirely within this province have average indoor radon levels less than 1 pCi/L. Leon County averaged 1.7 and 1.8 pCi/L in the two surveys of the Florida Statewide Radon Study. Most of these data likely come from Tallahassee, which lies within an area of moderately elevated eU. This

area and those parts of southern Columbia, western Union, and northern Alachua County which are underlain by phosphatic rocks, and limited areas where coarse gravels occur in river terraces in the western panhandle, are likely to have elevated radon potential.

The Central Highlands province has variable geologic radon potential. Generally low radon potential occurs in low eU areas in the eastern and southern parts of this province. Moderate radon potential occurs in the western part of this province where uraniferous phosphatic rocks are close to the surface. Localized areas in which uranium contents of soils and shallow subsoils exceed 100 ppm are likely, and indoor radon levels may exceed 20 pCi/L or more where this occurs. Alachua (lies in both the Central and Northern Highlands), Marion, and Sumter Counties report indoor radon values exceeding 20 pCi/L. Excessively well-drained hillslopes may also contribute to higher radon potential.

The Gulf Coastal Lowland Province generally has low radon potential. High rainfall and high water tables cause very moist soils which inhibit radon movement. Equivalent uranium is low in most areas except in some coastal bay areas of western peninsular Florida. Some isolated areas of elevated radon potential may occur in these areas of higher eU.

The Atlantic Coastal Lowland area generally has low radon potential. High rainfall and high water tables cause very moist soils that inhibit radon movement. Equivalent uranium is low in most areas. In some beach sand areas in northern Florida, elevated eU seems to be associated with heavy minerals; however, there is no evidence to suggest that elevated indoor radon occurs in these areas.

An area in southwestern Dade County, underlain by thin sandy soils covering shallow limestone bedrock, has equivalent uranium values as high as 3.5 ppm. Unusually high levels of radium are present in soils formed on the Pleistocene Key Largo Limestone and perhaps on other rock formations in certain areas of the Florida Keys and in southwestern Dade County. Areas of elevated eU and elevated indoor radon in Dade County are likely related to these unusual soils. These soils may be responsible for the modestly elevated eU in soils and for the elevated indoor radon levels, and they may extend into Collier County as well.

GEORGIA

Piedmont and Blue Ridge

The oldest rocks in Georgia form the mountains and rolling hills of the Blue Ridge Province and most of the Piedmont Province. These highly deformed rocks are separated by a series of thrust faults superimposing groups of older rocks over younger rocks, comprising the Georgiabama Thrust Stack. The igneous and metamorphic rocks in the Georgiabama Thrust Stack north of the Altoona Fault have been ranked moderate overall in geologic radon potential, but the radon potential of the area is variable. Mafic rocks are expected to have low radon potential whereas phyllite, slate, some metagraywacke, granitic gneiss and granite have moderate to high radon potential. Soil permeability is slow to moderate in most soils. Counties in this area have average indoor radon levels that vary from low to high (< 1 pCi/L to > 4 pCi/L), but the measurements are predominantly in the moderate range. The highest indoor radon reading, 18.7 pCi/L, was measured in the northern Blue Ridge in Fannin County, which is underlain predominantly by metagraywacke, slate, phyllite, and mica schists. Equivalent uranium concentrations in rocks and soils of this area are moderate to high.

The Georgiabama Thrust Stack south of the Alatoona Fault has also been ranked moderate in geologic radon potential. The majority of this part of the Georgiabama Thrust Stack is underlain

by schist and amphibolite of the Zebulon sheet, which have generally low radioactivity where not intruded by granites or where not highly sheared, particularly south of the Towaliga Fault. An area with distinctly low aeroradiometric readings which is underlain by mafic metamorphic rocks lies between the Brevard and Allatoona Faults in the northwestern Georgiabama Thrust Stack. All of these rocks have slow to moderate permeability, and indoor radon values are generally low to moderate. A central zone of biotite gneiss, granitic gneiss, and granite has elevated uranium concentrations and high equivalent uranium (>2.5 ppm) on the NURE map. Soil permeability is generally low to locally moderate. Indoor radon levels are generally moderate. Recent soil-gas radon studies in the Brevard zone and surrounding rocks show that this zone may yield unusually high soil-gas radon where the zone crosses the Ben Hill and Palmetto granites. Surface gamma-ray spectrometer measurements yielded equivalent uranium from 4 to 17 ppm over granite and granitic biotite gneiss (Lithonia gneiss). Soil-gas radon concentrations commonly exceeded 2,000 pCi/L and the highest soil-gas radon measured was 26,000 pCi/L in faulted Ben Hill granite. Undeformed Lithonia gneiss had average soil radon of more than 2,000 pCi/L. Mica schist averaged less than 1,000 pCi/L where it is undeformed. The Stone Mountain granite and mafic rocks yielded low soil-gas radon. The Grenville Basement granite and granite gneiss have moderate to locally high radon potential. Radioactivity is generally moderate to high and soil permeability is generally moderate.

The Little River Thrust Stack is generally low to moderate in geologic radon potential. It is underlain primarily by mafic metamorphic rocks with low radon potential, but each belt contains areas of rocks with moderate to locally high radon potential. Metadacites have moderate radon potential and moderate radioactivity. Faults and shear zones have local areas of mineralization and locally high permeability. Granite intrusives may also have moderate radon potential. Aeroradioactivity is generally low and soil permeability is generally moderate.

Ridge and Valley

The Rome-Kingston Thrust Stack is ranked low in geologic radon potential; however, some of the limestones and shales in this area may have moderate to high radon potential. Indoor radon is variable but generally low to moderate. Permeability of the soils is low to moderate. Equivalent uranium is moderate to locally high, especially along the Carters Dam and Emerson faults. Carbonate soils of the Valley and Ridge Province are likely to cause indoor radon problems. The Devonian Chattanooga Shale, which crops out locally in parts of the Valley and Ridge, is highly uraniferous and has been identified as a source of high indoor radon levels in Kentucky. Numerous gamma radioactivity anomalies are associated with the Pennington Formation, Bangor Limestone, Fort Paine Chert, Chattanooga Shale, Floyd Shale, the Knox Group, and the Rome Formation.

Appalachian Plateau

The Appalachian Plateau has been ranked low in geologic radon potential. Sandstone is the dominant rock type and it generally has low uranium concentrations. Equivalent uranium is low to moderate. Permeability of the soils is moderate and indoor radon levels are low.

Coastal Plain

The Coastal Plain has been ranked low in radon potential, but certain areas of the Coastal Plain in which glauconitic, carbonaceous, and phosphatic sediments are abundant may have moderate geologic radon potential. The highest soil-gas radon concentrations (>1000 pCi/L) and

equivalent uranium (eU) concentrations (>2 ppm) in studies of radon in soil-gas in the Coastal Plain of Alabama were found in the carbonaceous sands and clays of the Providence Sand and the glauconitic sands of the Eutaw and Ripley Formations. Low to moderate soil-gas radon and uranium concentrations were measured in the glauconitic sands, limestones, and clays of the Tertiary Nanafalia and Lisbon Formations, and the Tuscahoma Sand. Equivalent rock units in Georgia are also likely to be sources of high radon levels. Equivalent uranium is moderate in the Cretaceous and Tertiary-age sediments and low, with local highs, in the Quaternary sediments. Radioactivity highs in much of the Coastal Plain are related to phosphate and heavy-mineral concentrations. In the shoreline complexes and in several sediment units such as the Hawthorn Formation, the phosphate concentrations are naturally occurring. In the Black Lands and in many portions of the central Coastal Plain that have abundant agricultural activity, the radioactivity may be related to the use of phosphate fertilizers. Indoor radon in the Coastal Plain is generally low.

KENTUCKY

Three primary areas in Kentucky are identified as being underlain by rock types and geologic features suspected of producing elevated radon levels: (1) areas underlain by Devonian black shales in the Outer Bluegrass region; (2) areas underlain by the Ordovician Lexington Limestone, particularly the Tanglewood Member, in the Inner Bluegrass region; and (3) areas of the Mississippian Plateau underlain by karsted limestones or black shales. In addition, some homes underlain by, or in close proximity to, major faults in the Western Coalfield and Inner Bluegrass regions may have locally elevated indoor radon levels due to localized concentrations of radioactive minerals and higher permeability in fault and fracture zones.

Appalachian Plateau

The black shale and limestone areas in the Mississippian Plateau region have associated high surface radioactivity, and the Western Coalfield contains scattered radioactivity anomalies. The arcuate pattern of radioactivity anomalies along the southern edge of the Outer Bluegrass region corresponds closely to the outcrop pattern of the New Albany Shale. A group of radiometric anomalies in the vicinity of Warren and Logan counties appears to correspond to outcrops of the Mississippian Ste. Genevieve and St. Louis Limestones. The clastic sedimentary rocks of the Cumberland Plateau region are characterized by relatively low surface radioactivity and generally have low indoor radon levels.

In the Mississippian Plateau Region, locally elevated indoor radon levels are likely in areas with high soil permeability, solution cavities, or localized concentrations of radioactive minerals in karst regions, and in areas underlain black shale along the State's southern border. Of particular concern are the Devonian-Mississippian Chattanooga Shale (equivalent to the New Albany Shale), limestones in the Mississippian Fort Payne Formation, and the Mississippian Salem, Warsaw, Harrodsburg, St. Louis, and Ste. Genevieve Limestones in south-central Kentucky.

Caves, produced by limestone solution and relatively common in central Kentucky, are natural concentrators of radon and can be a local source of high radon levels. Levels of radon decay products approaching a maximum of 2.0 working levels (WL), which corresponds to about 400 pCi/L of radon (assuming that radon and its decay products are in 50 percent equilibrium), and averaging about 0.70 WL, or about 140 pCi/L of radon, have been recorded in Mammoth Cave. Although these levels are not considered hazardous if the exposure is of short duration, such as would be experienced by a visitor to the cave, it could be of concern to National Park Service employees and other persons that spend longer periods of time in the caves. Another potential hazard is the use of cave air for building air temperature control, as was formerly done at the

Mammoth Cave National Park visitor center. The cave air, which averages 54°F, was pumped into the visitor center for cooling, but this process has been discontinued due to the relatively high radioactivity associated with the cave air.

Coastal Plain

The majority of homes in the Jackson Purchase Region (Coastal Plain) have low indoor radon levels, although the area is underlain in part by loess with an eU signature in the 2.0-3.0 ppm range. The poor correspondence with surface radioactivity in this area appears to be due to a combination of low soil permeability and high water tables. The Coastal Plain is the only part of the State in which seasonal high water tables were consistently listed in the SCS soil surveys as less than 6 ft, and commonly less than 2 ft.

MISSISSIPPI

Examination of the available data reveals that Mississippi is generally an area of low radon potential. Indoor radon levels in Mississippi are generally low; however, several counties had individual homes with radon levels greater than 4 pCi/L. Counties with maximum levels greater than 4 pCi/L are concentrated in the northeastern part of the State within the glauconitic and phosphatic sediments of the Tombigbee Hills and Black Prairies. Readings greater than 4 pCi/L also occur in the Mississippi Alluvial Plain, the eastern part of the Pine Hills Province, and in loess-covered areas. Glauconitic and phosphatic sediments of the Coastal Plain, particularly the Cretaceous and lower Tertiary-age geologic units located in the northeastern portion of the State, have some geologic potential to produce radon. Based on radioactivity and studies of radon in other parts of the Coastal Plain, the Black Prairies and Pontotoc Ridge have been assigned moderate geologic radon potential; all other parts of Mississippi are considered to be low in geologic radon potential. The climate, soil, and lifestyle of the inhabitants of Mississippi have influenced building construction styles and building ventilation which, in general, do not allow high concentrations of radon to accumulate.

Coastal Plain

A study of the radon in the Coastal Plain of Texas, Tennessee, and Alabama suggests that glauconitic, phosphatic, and carbonaceous sediments and sedimentary rocks, equivalent to those in Mississippi, can cause elevated levels of indoor radon. Ground-based surveys of radioactivity and radon in soils in that study indicate that the Upper Cretaceous through Lower Tertiary Coastal Plain sediments are sources of high soil-gas radon (> 1,000 pCi/L) and soil uranium concentrations. The high equivalent uranium found over the Coastal Plain sediments in northeastern Mississippi supports the possibility of a similar source of high radon levels. Chalks, clays and marls tend to have low permeability when moist and higher permeability when dry due to desiccation fractures and joints.

The youngest Coastal Plain sediments, particularly Oligocene and younger, have decreasing amounts of glauconite and phosphate and become increasingly siliceous and therefore less likely to be significant sources of radon. Some carbonaceous units may be possible radon sources.

Loess in Tennessee, and probably elsewhere, is known to generate high levels of radon in both dry and saturated soils. Both thin and thick loess units can easily be traced on the

radioactivity map of Mississippi by following the highest of the moderate equivalent uranium anomalies. Loess tends to have low permeability when moist and higher permeability when dry.

Mississippi Alluvial Plain

The Mississippi Alluvial Plain contains several areas with locally high eU, as well as having moderate radioactivity overall. These high eU areas are located close to the river in Bolivar and Washington Counties. The highest indoor radon level recorded in Mississippi in the State/EPA Residential Radon Survey (22.8 pCi/L) occurs within Bolivar County and the second highest radon level of homes measured to date in the State (16.1 pCi/L) occurs in Washington County. It is not apparent from the data available whether the high eU and indoor radon levels are correlative, and only a few indoor radon readings in each county are greater than 4 pCi/L. The geology of the region is not generally conducive to high uranium concentrations, except possibly in heavy-mineral placer deposits. Further, elevated radioactivity in the Mississippi Alluvial Plain may be due in part to uranium in phosphatic fertilizers. Locally high soil permeability in some of the alluvial sediments may allow locally high indoor radon levels to occur.

The southeastern half of Mississippi has low radioactivity and low indoor radon levels. The few indoor radon readings greater than 4 pCi/L were between 4.1 and 5.8 pCi/L. The lowest eU is associated with the coastal deposits and the Citronelle Formation, which are predominantly quartz sands with low radon potential. Slightly higher eU, though still low overall, is associated with the Pascagoula and Hattiesburg Formations and Catahoula Formation. Soils in this area are variably poorly to well drained with slow to moderate permeabilities.

The Chattanooga Shale and related sedimentary rocks in the northeastern part of the State have the potential to be sources of high indoor radon levels. In Tennessee and Kentucky, the Chattanooga Shale has high uranium concentrations and is associated with high indoor radon levels in those states. The extent of these rocks in Mississippi is minor.

NORTH CAROLINA

Blue Ridge

The Blue Ridge has been ranked moderate overall in geologic radon potential, but it is actually variably moderate to high in radon potential. The province has highly variable geology and because of the constraints imposed by viewing the indoor radon data at the county level, it is impossible to assign specific geologic areas of the Blue Ridge to specific moderate or high indoor radon levels. Average indoor radon levels are moderate (2-4 pCi/L) in the majority of counties. However, two counties have indoor radon averages between 4.1 and 6 pCi/L (Cherokee and Buncomb Counties) and three counties in the northern Blue Ridge (Alleghany, Watauga, and Mitchell) have indoor radon averages greater than 6 pCi/L. These three counties are generally underlain by granitic gneiss, mica schist, and minor amphibolite and phyllite. Transylvania and Henderson Counties, which are underlain by parts of the Blue Ridge and Inner Piedmont, also have indoor radon averages greater than 6 pCi/L. The Brevard fault zone, Henderson Gneiss, and Ceasars Head Granite are possible sources of high indoor radon in these two counties. Equivalent uranium is variable from low to high in the Blue Ridge. The highest eU appears to be associated with the Ocoee Supergroup in the southern Blue Ridge, rocks in the Grandfather Mountain Window, and metamorphic rocks in parts of Haywood and Buncomb Counties. Soils are generally moderate in permeability.

The Chauga belt and Brevard fault zone are ranked high in geologic radon potential. The Chauga belt consists predominantly of the Henderson Gneiss. High eU and high uranium in stream sediments appears to be associated with the Brevard fault zone, Henderson Gneiss, and Ceasars Head Granite in this area. Average indoor radon levels in the two counties that the main part of the Chauga belt and the southern portion of the Brevard fault zone passes through are high. The soils have moderate permeability.

Piedmont

The Inner Piedmont and Kings Mountain belts have been ranked moderate in geologic radon potential. Indoor radon levels are generally moderate. Granitic plutons, granitic gneiss, monazite-rich gneiss and schist, pegmatites, and fault zones appear to have high eU and high uranium concentrations in stream sediment samples. Many of the granitic plutons are known to be enriched in uranium and recent studies suggest that the soils developed on many of the uraniferous granitic plutons and related fault zones in the Blue Ridge and Piedmont are possible sources of radon. Measured soil-gas radon concentrations commonly exceeded 1,000 pCi/L in soils developed on the Cherryville Granite, Rolesville Suite, and the Sims, Sandy Mush, and Castalia plutons. Soils developed on the Rocky Mount, Spruce Pine, Toluca, Mt. Airy, and Stone Mountain plutons had relatively low soil-gas radon concentrations. Soil permeabilities in the Inner Piedmont, Brevard fault zone, and Kings Mountain belt are variably low to moderate which, together with the large proportion of homes without basements, may account for the abundance of moderate indoor radon levels.

Most shear zones in the Piedmont and Blue Ridge should be regarded as having the potential to produce very localized moderate to high indoor radon levels. Geochemical and structural models developed from studies of shear zones in granitic metamorphic and igneous rocks from the Reading Prong in New York to the Piedmont in Virginia indicate that uranium enrichment, the redistribution of uranium into the rock foliation during deformation, and high radon emanation, are common to most shear zones. Because they are very localized sources of radon and uranium, shear zones may not always be detected by radiometric or stream sediments surveys.

The Charlotte belt has been ranked low in geologic radon potential but it is actually quite variable–dominantly low in the southern portion of the belt and higher in the northern portion of the belt. Equivalent uranium is generally low, with locally high eU occurring in the central and northern portions of the belt, associated with the Concord and Salisbury Plutonic Suites. Permeability of the soils is generally low to moderate and indoor radon levels are generally low.

The Carolina slate belt has been ranked low in radon potential where it is underlain primarily by metavolcanic rocks. Where it crops out east of the Mesozoic basins it has been ranked moderate. Aeroradioactivity over the Carolina slate belt, uranium in stream sediment samples, and indoor radon levels are markedly low. Permeability of many of the metavolcanic units is generally low to locally moderate. East of the Wadesboro subbasin in Anson and Richmond Counties lies a small area of the slate belt that is intruded by the Lilesville Granite and Peedee Gabbro. It has high eU and high uranium concentrations in stream sediments, and moderate to high permeability in the soils, and is a likely source of moderate to high indoor radon levels.

The Raleigh belt has been ranked moderate in geologic radon potential. Equivalent uranium in the Raleigh belt is generally moderate to high and appears to be associated with granitic intrusive rocks, including the Castalia and Wilton plutons and the Rolesville Suite. A belt of monazite-bearing rocks also passes through the Raleigh belt and may account for part of the observed high

radioactivity. Soils have variably low to moderate permeability. Indoor radon levels are generally moderate.

Coastal Plain

In the Coastal Plain province, moderate to high eU is associated with the Cretaceous and Tertiary sediments of the Inner Coastal Plain. Permeability of the soils is highly variable but generally moderate to low, and may be locally high in sands and gravels. Seasonally high water tables are common. Indoor radon levels in the Coastal Plain are generally low. The Inner Coastal Plain is ranked low in geologic radon potential but may be locally moderate to high, especially in areas underlain by Cretaceous sediments. Glauconitic, phosphatic, monazite-rich, and carbonaceous sediments and sedimentary rocks in the Coastal Plain of Texas, New Jersey, and Alabama, similar to some Coastal Plain sediments in North Carolina, are the source for moderate indoor radon levels seen in parts of the Inner Coastal Plain of these states.

The Outer Coastal Plain has low eU, low indoor radon levels, and is generally underlain by sediments with low uranium concentrations. Soil permeability is variable but generally moderate. Seasonally high water tables are common. A few isolated areas of high radioactivity in the Outer Coastal Plain may be related to heavy mineral and phosphate deposits in the shoreline sediments. The area has been ranked low in geologic radon potential, but may have local moderate or high indoor radon occurrences related to heavy minerals or phosphate deposits.

SOUTH CAROLINA

Blue Ridge and Piedmont

The Blue Ridge and Piedmont Provinces have moderate geologic radon potential. Possible sources of radon include uraniferous granites, biotite and granitic gneiss, and shear zones. Soils developed on many of the uraniferous granitic plutons and some fault zones within the Piedmont and Blue Ridge of North and South Carolina yield high soil-gas radon (>1,000 pCi/L). In the Blue Ridge, sheared graphitic rocks may be a local source for high indoor radon concentrations.

More than 10 percent of the homes tested in Greenville and Oconee Counties, in the Blue Ridge and Piedmont, have indoor radon levels greater than 4 pCi/L. Greenville County also has the highest indoor radon measurement in the State, 80.7 pCi/L, the highest radioactivity, associated with the Silurian-Devonian Ceasers Head Granitic Gneiss, and with biotite gneiss in the Carolina monazite belt. In Oconee County, the Toxaway Gneiss and graphitic rocks in the Brevard Fault Zone may account for the higher incidence of indoor radon levels exceeding 4 pCi/L and the higher overall indoor radon average of the county. Average indoor radon levels in the Blue Ridge and Piedmont are generally higher than in the rest of the State, and moderate to high radioactivity is common. Most of the soils formed on granitic rocks have moderate permeability and do not represent an impediment to radon mobility. Mafic rocks in the Blue Ridge and Piedmont have low radon potential. These rocks have low concentrations of uranium, and soils formed from them have low permeability.

Coastal Plain

In the Coastal Plain Province, moderate to high radioactivity is associated with the Cretaceous and Tertiary sediments of the upper Coastal Plain. Glauconitic, phosphatic, monazite-rich, and carbonaceous sediments and sedimentary rocks in the Coastal Plain of Texas, New Jersey, and Alabama, similar to some of those in South Carolina, cause elevated levels of indoor

radon. Orangeburg County is the only other county besides Greenville and Oconee Counties that has an average indoor radon level greater than 2 pCi/L. It is underlain by Lower Tertiary sediments in an extremely dissected part of the Coastal Plain. Radioactivity is moderate to low. Soils are highly variable in the county because of the complicated erosion patterns. The few high values of indoor radon for this county create an overall higher indoor radon average for the county. These locally high readings may be due to local accumulations of monazite, glauconite, or phosphate that can occur within these particular sediments.

The lower Coastal Plain has low to locally high radioactivity and low indoor radon levels. Most of the sediments have low uranium concentrations with the exception of the uraniferous, phosphatic sediments of the Cooper Group and local, heavy-mineral placer deposits within some of the Quaternary units. The area has been ranked low in geologic radon potential overall, but the radon potential may be locally high in areas underlain by these uraniferous sediments.

TENNESSEE

Coastal Plain and Mississippi Alluvial Plain

The Mississippi Alluvial Plain has low geologic radon potential. The high soil moisture, high water tables, and the lack of permeable soils lower the radon potential in spite of moderate eU values. Some areas with very sandy or excessively-drained soils may cause homes to have indoor radon levels exceeding 4 pCi/L.

The loess-covered parts of the Coastal Plain have low radon potential in spite of moderate eU values and elevated soil-gas radon concentrations. The radon potential is lowered by the high moisture content and low permeability of the soils. The lack of basements in homes also lowers the potential. If prolonged dry periods were to occur in this area, some homes might see a significant increase in indoor radon, especially those with basements or crawl spaces. The eastern Coastal Plain has moderate geologic radon potential. NURE data show elevated eU values compared to the rest of the Coastal Plain. Soil-gas radon levels are locally elevated.

Highland Rim and Nashville Basin

The Highland Rim and Nashville Basin are underlain by sedimentary rocks of Paleozoic age, principally limestone, shale, chert, and dolostone. The part of the Highland Rim that is underlain by cherty limestone (Fort Payne Formation) has high geologic radon potential. This area has moderate to locally high eU and soils that are cherty and excessively well drained. The limestone and shale part of the Highland Rim has moderate radon potential. The Nashville Basin has high geologic radon potential. The elevated eU, the presence of abundant phosphatic soils, local karst, and the presence of generally well-drained soils all contribute to this high geologic radon potential. Very high (>20 pCi/L) to extreme indoor radon values (>200 pCi/L) are possible where homes are sited on soils developed on the Chattanooga shale, on phosphate-rich residual soils, or on karst pinnacles.

Appalachian Plateau

Sandstones and shales underlie most of the Appalachian Plateau, which generally has moderate geologic radon potential. These rocks are typically not good sources of radon and values for eU are among the lowest in the State. However, many sandy, well-drained to excessively-drained soils are present in this region, and may be a source of locally elevated radon levels because of their high permeability.

Ridge and Valley

Folded and faulted Paleozoic limestone, shale, chert, dolostone, and sandstone underlie most of the Ridge and Valley region, with sandstone and cherty dolostone forming most of the ridges and limestone and shale forming most of the valleys. The Ridge and Valley region has high geologic radon potential because of elevated eU values, karst, and well drained soils. Very high (>20 pCi/L) to extreme indoor radon values (>200 pCi/L) are possible where homes are sited on soils developed on black shales, phosphate-rich residual soils, or karst pinnacles. Homes with basements are more likely to yield elevated indoor radon levels than homes with slab-on-grade construction.

Unaka Mountains

The Unaka Mountains are underlain by siltstone, sandstone, conglomerate, quartzite, phyllite, gneiss, granite, and metamorphosed volcanic rocks of Precambrian and Paleozoic age that have moderate geologic radon potential. Values of eU are generally moderate, although they are locally high. Some very high (>20 pCi/L) to extreme (>200 pCi/L) indoor radon levels are possible where homes are sited on phosphate-rich residual soils developed on phosphatic carbonate rocks, or on pegmatite in the metamorphic rock areas, but the former are much less common in this region than in the Nashville Basin and the Ridge and Valley region.

PRELIMINARY GEOLOGIC RADON POTENTIAL ASSESSMENT OF TENNESSEE

by
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U.S. Geological Survey

INTRODUCTION

This assessment of the radon potential of Tennessee is largely dependent on geologic information derived from publications of the Tennessee Department of Environment and Conservation, Division of Geology, and from publications of the U.S. Geological Survey. An analysis of indoor radon data collected as part of the State/EPA Residential Radon Survey in the winter of 1987-1988 by the U.S. EPA and the Tennessee Department of Health is also included in this report. The National Atlas of the United States of America provided much information on the geographic setting. Soil descriptions are developed from Springer and Elder (1980).

This is a generalized assessment of geologic radon potential of rocks, soils, and surficial deposits of Tennessee. The scale of this assessment is such that it is inappropriate for use in identifying the radon potential of small areas such as neighborhoods, individual building sites, or housing tracts. Any localized assessment of radon potential must be supplemented with additional data and information from the locality. Within any area of a given radon potential ranking, there are likely to be areas with higher or lower radon levels than characterized for the area as a whole. Indoor radon levels, both high and low, can be quite localized, and there is no substitute for testing individual homes. Elevated levels of indoor radon have been found in every state, and EPA recommends that all homes be tested. For more information on radon, the reader is urged to consult the local or State radon program or EPA regional office. More detailed information on state or local geology may be obtained from the State geological survey. Addresses and phone numbers for these agencies are listed in chapter 1 of this booklet.

GEOGRAPHIC SETTING

The State of Tennessee extends over 400 miles from the Mississippi River along its western boundary to the Unaka Mountains along its eastern edge. Tennessee can be subdivided into seven physiographic regions (fig. 1): the Mississippi Alluvial Plain; the Gulf Coastal Plain; the Highland Rim; the Nashville Basin; the Appalachian Plateau; the Appalachian Ridge and Valley Region; and the Unaka Mountains. The elevation of Tennessee ranges from less than 300 feet above sea level along the Mississippi River to 6,642 feet at Clingman's Dome in the Unaka Mountains.

The Mississippi Alluvial Plain is a narrow section of swamp and floodplain of very low relief (less than 50 feet) along the Mississippi River. The Gulf Coastal Plain is characterized by a very gently rolling plain dissected by streams that drain westward into the Mississippi River. The eastern part is very hilly. Relief ranges 50-150 feet. The Highland Rim is characterized by areas of gently rolling plains and low open hills with 300-500 feet of relief. The Nashville Basin has a gently rolling plain with small hills known as knobs. Relief ranges from 300 to 500 feet. Both the Highland Rim and the Nashville Basin are drained by the Tennessee and Cumberland Rivers and their tributaries.

The Appalachian Plateau is an area of high plateaus and open low mountains with 500-1000 feet of relief. A few deep, narrow valleys dissect this surface. The Appalachian Ridge and

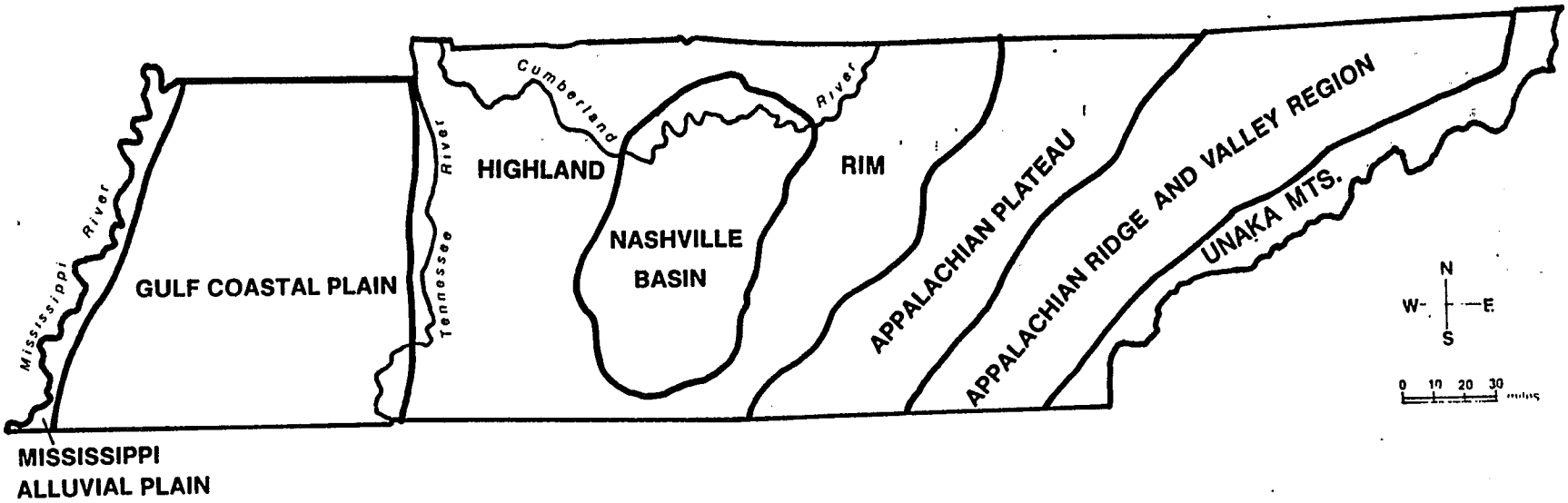

Figure 1- Physiographic provinces of Tennessee (modified from Facts on File, 1984).

Valley Region has a series of low northeast-trending ridges that rise up to 800 feet above the adjacent valleys. The Unaka Mountains are typified by rugged mountains and narrow valleys with 1000-3000 feet of relief.

Population distribution and land use in Tennessee reflects in part the geology, topography, and climate of the State. In 1990,the population was 4,877,185, including 60 percent urban population (fig. 2) concentrated mostly in Memphis, Nashville, and Knoxville. The average population density is approximately 116 per square mile. The climate is humid continental to the north and humid sub-tropical to the south. Precipitation averages about 50 inches per year (fig. 3).

Much of the State consists of cropland with lesser pasture and woodland. Hilly and mountainous areas are predominantly grazed and ungrazed woodland and forest. The principal crops include soybeans, tobacco, wheat, cotton, and corn.

GEOLOGIC SETTING

The Coastal Plain is underlain by unconsolidated Cretaceous and younger marine and fluvial sand, clay, and marl that dip gently westward into the Mississippi Embayment, a structural sag cutting into the southern edge of the mid-continent (fig. 4). Deposits of loess derived from the Mississippi River valley cover the western part of the Coastal Plain. They are more than 20 m thick near the Mississippi River but thin to zero to the east. Along the eastern edge of the Coastal Plain the marine sediments are locally phosphatic.

The Highland Rim, the Nashville Basin, and the Appalachian Plateau of central Tennessee lie within the stable midcontinent area of the United States. The first two regions are underlain primarily by gently folded sedimentary rocks of Paleozoic age, principally limestone, shale, chert, and dolostone, whereas gently dipping sandstones and shales underlie most of the Appalachian Plateau (fig. 4). Limestones dominate the Nashville Basin. Folded and faulted Paleozoic limestone, shale, chert, dolostone, and sandstone underlie most of the Ridge and Valley region, with sandstone and cherty dolostone forming most of the ridges and limestone and shale forming most of the valleys. The Unaka Mountains are underlain by siltstone, sandstone, conglomerate, quartzite, phyllite, gneiss, granite, and metamorphosed volcanic rocks of Precambrian and Paleozoic age (fig. 4). Rocks of the Unaka Mountains have been folded, thrusted, and metamorphosed.

Uranium occurrences are uncommon in Tennessee and are largely associated with pegmatites in Precambrian crystalline rocks or radioactive residual soils on crystalline rocks in Carter, Johnson, and Unicoi Counties (fig. 5). The Mississippian and Devonian Chattanooga Shale is the most uraniferous sedimentary rock formation in the State. The Chattanooga shale crops out widely around the periphery of the Nashville Basin, along the west side of the Highland Rim, along the Sequatchie Valley within the Appalachian Plateau, and along several ridges in the southern half of the Ridge and Valley region. The Chattanooga Shale also occurs in the northern half of the Valley and Ridge. It varies in thickness across Tennessee and ranges from 8 feet in Chattanooga to 1500 feet near the Kentucky and Virginia state lines. The Chattanooga Shale may contain as much as 200 ppm U but it typically contains 10-100 ppm U. Thickness of the Gassaway Member, the most uraniferous part of the Chattanooga Shale, is generally less than 30 feet. In some areas, the Chattanooga Shale is enriched enough in uranium that it was considered a uranium exploration target during World War II and during the 1970s.

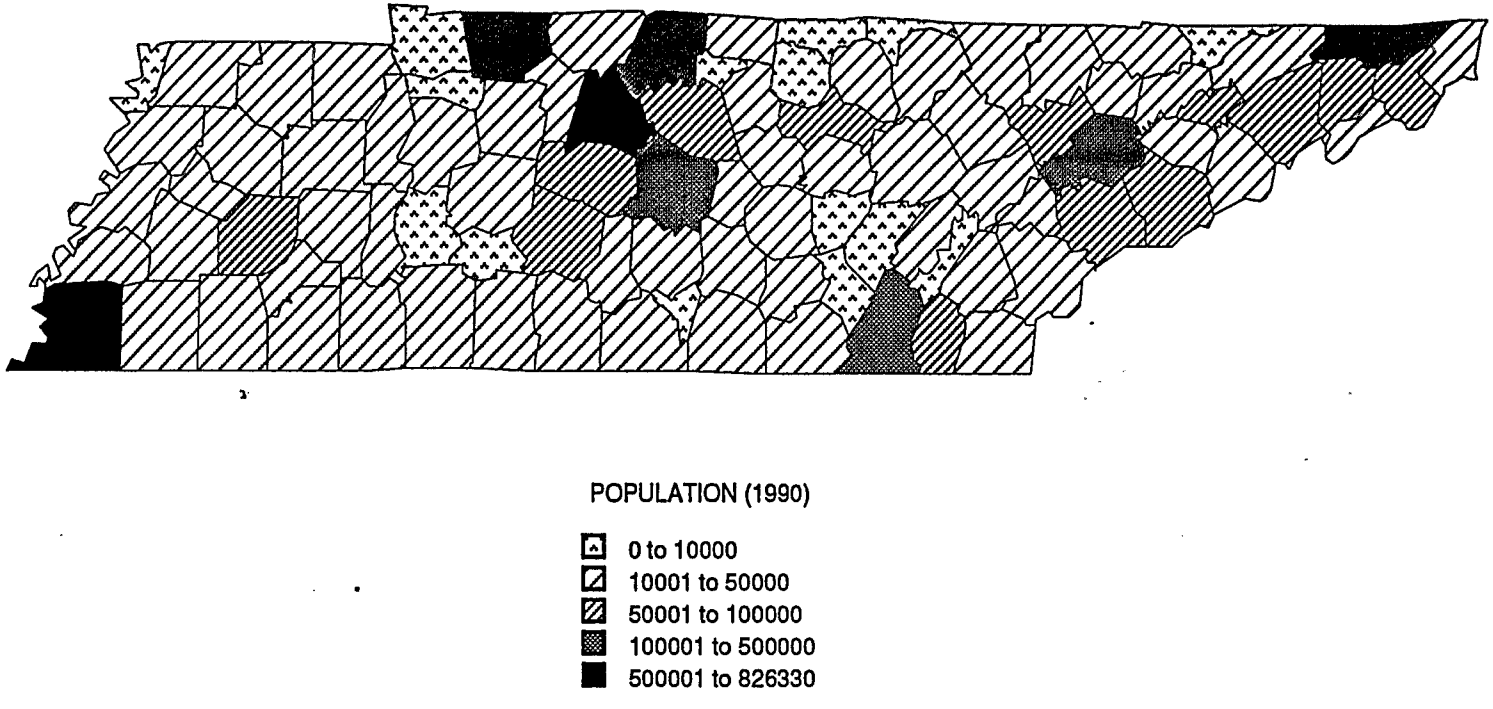

Figure 2. Population of counties in Tennessee (1990 U.S. Census data).

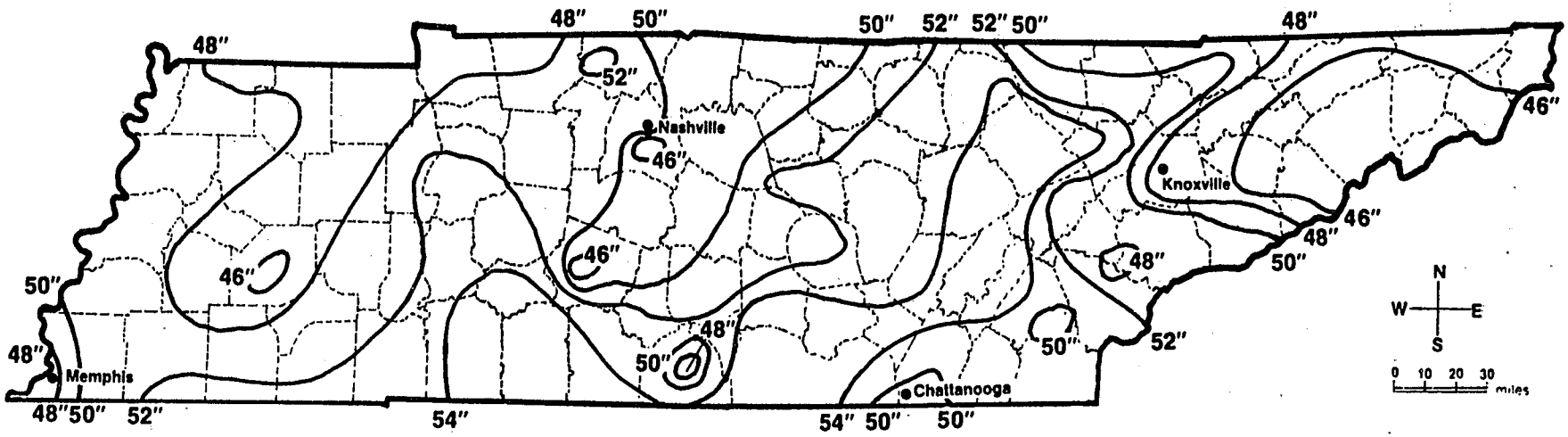

Figure 3- Map showing precipitation in Tennessee (modified from Facts on File, 1984).

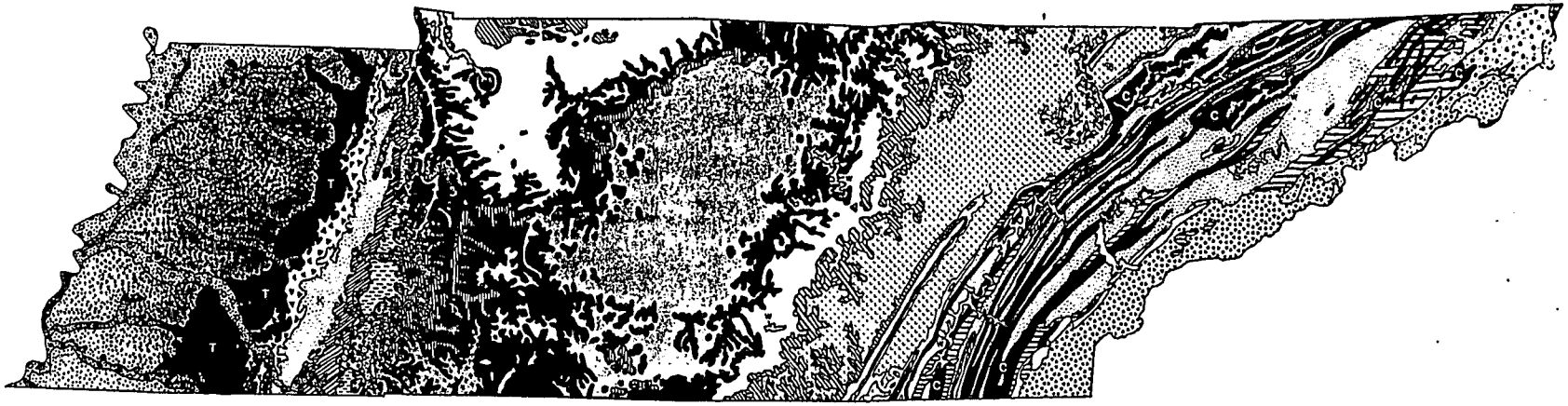

Figure 4- Generalized bedrock geologic map of Tennessee. From Bennison (1989). w- water. See accompanying explanation for map units.

EXPLANATION FOR THE GEOLOGIC MAP

QUATERNARY

Holocene alluvium

Holocene loess, dune sand

Holocene terrace deposits

TERTIARY

Jackson Formation- Sand, with layers of gray clay, silt, and lignite; thickness at least 60 ft

Claiborne and Wilcox Formations- Irregularly bedded sand, locally interbedded with lenses and beds of gray to white clay, silty clay, lignite clay, and lignite. Thickness more than 400 ft.

Porters Creek Clay- Pale-brown to brownish-gray, massive, blocky clay; locally contains glauconitic sand. Thickness 130 to 170 ft.

Clayton Formation- Glauconitic sand, argillaceous and locally fossiliferous impure fossiliferous limestone. Thickness 30 to 70 ft.

CRETACEOUS

Owl Creek Formation- Sandy clay, greenish-gray, glauconitic, fossiliferous; merges northward into unfossiliferous clays and sands. Thickness 0 to about 40 ft.

McNairy Sand- Predominantly sand, in places interbedded with silty light-gray clays. Fine-grained sand at base, locally contains heavy minerals. Thickness about 300 ft.

Coon Creek Formation- Fossiliferous, micaceous sand, silty and glauconitic; locally fossiliferous sandy clay at base. Siderite concretions common in upper part. Thickness about 140 ft.

Demopolis Formation- Marl and calcareous clay, light-gray, fossiliferous, locally glauconitic and sandy. Merges northward into sands mapped as Coon Creek Formation. Maximum thickness 180 ft.

Sardis Formation- Quartz sand and glauconite sand, argillaceous and locally fossiliferous. (Mapped with Kcc north of Beech River.) Maximum thickness 70 ft.

Coffee Sand- Loose fine-grained sand, light-gray, sparsely glauconitic, locally interbedded with laminated lignitic clay. Thickness 25 to 200 ft; thins northward.

Eutaw Formation- Grayish-green sand, fine-grained, glauconitic, micaceous; interbedded with gray laminated clays which commonly contain carbonized or silicified wood. (Mapped with Coffee except in Hardin County and southeastern Decatur County.) Thickness 0 to 180 ft; thins northward.

Tuscaloosa Formation- Poorly sorted, light-gray, chert gravel in a matrix of silt and sand; locally interbedded with sand and clay lenses. Thickness 0 to 140 ft.

PENNSYLVANIAN

Cross Mountain Formation- Mostly shale, interbedded with sandstone, siltstone, and thin coal beds; base at top of Frozen Head Sandstone. Maximum preserved thickness 554 ft.

Vowell Mountain Formation- Shale, sandstone, siltstone, and coal; from Frozen Head Sandstone Member to Pewee coal. Thickness 230 to 375 ft.

Redoak Mountain Formation- Shale, sandstone, siltstone, and several important coals; from Pewee coal to Windrock coal. Thickness 340 to 420 ft.

Graves Gap Formation- Shale, sandstone, siltstone, and coal; from Windrock coal to top of Pioneer Sandstone. Thickness 200 to 350 feet.

Indian Bluff Formation- Shale, sandstone, siltstone, and thin coal beds; from Pioneer Sandstone Member to Fellico coal. Thickness 150 to 250 feet.

Slatestone Formation- Shale, sandstone, siltstone, and several important coals; from Fellico coal to Poplar Creek coal. Thickness 500 to 650 feet.

Crooked Fork Group- Shale, sandstone, siltstone, and thin coal beds; from top down group includes Poplar Creek coal, Wartburg Sandstone, Glenmary Shale, Coalfield Sandstone, Burnt Mill Shale, Crossville Sandstone, and Dorton Shale. Thickness 200 to 450 feet.

Crab Orchard Mountains and Gizzard Groups- Sandstone, conglomerate, siltstone, shale, and thin coal beds. From top down Crab Orchard Mountains Group includes Rockcastle Conglomerate, Vandever Formation, Newton Sandstone, Whitwell Shale, and Sewanee Conglomerate; Gizzard Group includes Signal Point Shale, Warren Point Sandstone, and Racoon Mountain Formation. Thickness about 1,200 to 1,400 feet.

MISSISSIPPIAN

Mississippian undivided, but includes units named below.

Pennington Formation- Redish and greenish shale and siltstone; fine-grained dolomite; dark-gray limestone; and thin-bedded sandstone. Persistent dolomite bed at base. Thickness 150 to 400 ft.

Bangor Limestone- Dark brownish-gray limestone, thick-bedded. Thickness 70 to 400 ft.

Hartselle Formation- Thin-bedded, fine-grained sandstone interbedded with gray shale; with oolitic and coarse-grained limestone beds locally. Thickness 0 to 80 ft.

Monteagle Limestone- Mainly fragmental and oolitic, light-gray limestone; blocky bryozoan chert weathers from base. Thickness 180 to 300 ft.

St. Louis Limestone- Fine-grained, brownish-gray limestone, dolomitic and cherty. Thickness 80 to 160 ft.

Warsaw Limestone- Mainly medium-to coarse-grained, gray limestone, crossbedded. Includes much calcareous sandstone and shale to the north. thickness 100 to 130 ft.

Fort Payne Formation- Bedded chert, calcareous and dolomitic, somewhat crinoidal; and minor shale. Thin green shale (Maury) at base. Thickness about 300 ft.

Grainger Formation- Gray to green shale with siltstone and fine-grained glauconitic sandstone; in some areas quartz-pebble conglomerate. Thickness 500 to 1,000 ft.

DEVONIAN

Chattanooga Shale (Devonian-Mississippian)- Black carbonaceous shale, fissile. Thickness 100 to 900 feet; about 25 feet on Chilhowee Mountain.

Pegram Formation- Thick-bedded, gray limestone and gray sandstone. Thickness 0 to 15 ft.

(Devonian continued on next page.)

Camden Formation- Light-gray novaculitic chert and tripolitic clay; and minor siliceous limestone. Thickness 0 to about 100 ft.

Flat Gap Limestone- Thick-bedded, coarse-grained limestone, gray with red and brown grains. Thickness 0 to 55 ft.

Ross Formation- Siliceous limestone; gray and variegated shale; and medium-grained glauconitic limestone. Thickness 0 to 75 ft.

SILURIAN

Sneedville Limestone- Gray silty limestone and dolomite, minor shale, and fine-grained, greenish-gray sandstone; fossils locally abundant. Thickness 100 to 300 feet.

Rockwood Formation- Grown to maroon shale, thin gray siltstone and sandstone, and thin lenticular layers of oolitic and fossiliferous red hematite. Thickness 350 to 550 feet.

Clinch Sandstone- Clean, white, well-sorted sandstone; locally gray siltstone and shale. Average thickness about 600 feet.

ORDOVICIAN

Sequatchie Formation- Maroon and gray shaley limestone, mottled greenish; with interbeds of calcareous, olive to maroon shale and siltstone. Average thickness about 200 ft.

Reedsville Shale- Greenish-gray calcareous shale. Thickness 0 to 400 ft.

Moccasin Formation- Maroon calcareous shale, siltstone, and limestone; thin metabentonite layers in upper part; mud cracks ripple marks common. Thickness 800 to 1,000 ft.

Juniata Formation- Maroon claystone, siltstone, and shale; uniformly bedded; some faint greenish mottling; less calcareous than Sequatchie Formation. Thickness about 300 feet.

Martinsburg Shale- Bluish-gray, calcareous clay shale, weathers yellowish-brown; with thin beds of nodular gray, fossiliferous limestone; thin layers of metabentonite near base. Thickness about 1,000 feet.

Bays Formation- Maroon claystone and siltstone, commonly mottled greenish, evenly bedded; to northeast, light-gray to white, thick-bedded sandstone; metabentonite in upper part. Maximum thickness 1,000 feet.

Ottosee Shale- Bluish-gray calcareous shale, weathers yellow; with reef lenses of coarsely crystalline reddish fossiliferous limestone ("marble"). Thickness about 1,000 feet.

Holston Formation- Pink, gray, and red coarsely crystalline limestone (Holston Marble); in many areas upper part is sandy, crossbedded ferruginous limestone and brown to greenish calcareous shale. Thickness 200 to 600 feet.

Mannie Shale- Shale with thin beds of argillaceous limestone. Thickness 0 to 20 ft.

Hermitage Formation- Thin-bedded to laminated, sandy and argillaceous limestone with shale; nodular shaly limestone; coquina; and phosphatic calcarenite. Thickness 50 to 100 ft.

Carters Limestone- Fine-grained, yellowish-brown limestone; thin-bedded in upper part; thicker bedded and very slightly cherty with scattered mottlings of magnesian limestone in lower part. Contains thin bentonite beds. Thickness 50 to 100 ft.

Lebanon Limestone- Thin-bedded, gray limestone with calc shale partings. Thickness 80 to 100 ft.

Ridley Limestone- Thick-bedded, brownish-gray limestone, fine-grained,, with minor mottlings of magnesian limestone; slightly cherty. Thickness 90 to 150 ft.

Pierce Limestone- gray, thin-bedded limestone with shale partings. Thickness 25 ft.

Murfreesboro Limestone- Thick-bedded, dark-gray, fine-grained limestone, with mottlings of magnesian limestone; somewhat cherty in upper part. Maximum exposed thickness 70 ft.

Arnheim Formation- Nodular, shaly, gray limestone. Thickness 0 to 20 ft.

Leipers Formation- Nodular, shaly limestone; fine- to coarse-grained limestone; and phosphatic calcarenite locally. Thickness 0 to 150 ft.

Inman Formation- Thin-bedded to laminated, fine-grained, gray limestone with shale partings. Thickness 0 to 50 ft.

Catheys Formation- Nodular, shaly limestone; fine- to coarse-grained limestone; phosphatic calcarenite; and light-gray cryptograined limestone. Thickness 50 to 175 ft.

Bigby-Cannon Limestone- Brownish-gray phosphatic calcarenite and light-gray to brownish-gray, cryptograined to medium-grained, even-bedded limestone. Thickness 50 to 125 ft.

Sequatchie Formation- Maroon and gray shaley limestone, mottled greenish; with interbeds of calcareous, olive to maroon shale and siltstone. Average thickness about 200 ft.

Reedsville Shale- Greenish-gray calcareous shale. Thickness 0 to 400 ft.

Unnamed Limestone Unit- Medium-grained, fossiliferous, gray limestone, shaly in part. Thickness as much as 600 ft.

Moccasin Formation- Maroon calcareous shale, siltstone, and limestone; thin metabentonite layers in upper part; mud cracks ripple marks common. Thickness 800 to 1,000 ft.

Juniata Formation- Maroon claystone, siltstone, and shale; uniformly bedded; some faint greenish mottling; less calcareous than Sequatchie Formation. Thickness about 300 feet.

Martinsburg Shale- Bluish-gray, calcareous clay shale, weathers yellowish-brown; with thin beds of nodular gray, fossiliferous limestone; thin layers of metabentonite near base. Thickness about 1,000 feet.

Bays Formation- Maroon claystone and siltstone, commonly mottled greenish, evenly bedded; to northeast, light-gray to white, thick-bedded sandstone; metabentonite in upper part. Maximum thickness 1,000 feet.

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Carters Limestone- Fine-grained, yellowish-brown limestone; thin-bedded in upper part; thicker bedded and very slightly cherty with scattered mottlings of magnesian limestone in lower part. Contains thin bentonite beds. Thickness 50 to 100 ft.

Lebanon Limestone- Thin-bedded, gray limestone with calcareous shale partings. Thickness 80 to 100 ft.

Ridley Limestone- Thick-bedded, brownish-gray limestone, fine-grained,, with minor mottlings of magnesian limestone; slightly cherty. Thickness 90 to 150 ft.

Pierce Limestone- gray, thin-bedded limestone with shale partings. Thickness 25 ft.

Murfreesboro Limestone- Thick-bedded, dark-gray, fine-grained limestone, with mottlings of magnesian limestone; somewhat cherty in upper part. Maximum exposed thickness 70 ft.

Arnheim Formation- Nodular, shaly, gray limestone. Thickness 0 to 20 ft.

Leipers Formation- Nodular, shaly limestone; fine- to coarse-grained limestone; and phosphatic calcarenite locally. Thickness 0 to 150 ft.

Inman Formation- Thin-bedded to laminated, fine-grained, gray limestone with shale partings. Thickness 0 to 50 ft.

Catheys Formation- Nodular, shaly limestone; fine- to coarse-grained limestone; phosphatic calcarentie; and light-gray cryptograined limestone. Thickness 50 to 175 ft.

Bigby-Cannon Limestone- Brownish-gray phosphatic calcarenite and light-gray to brownish-gray, cryptograined to medium-grained, even-bedded limestone. Thickness 50 to 125 ft.

Lenoir Limestone- Nodular, argillaceous, gray limestone; in places basal sedimentary breccia conglomerate, quartz sand; Mosheim Limestone Member (dense light- to medium-gray limestone) near base. Thickness 25 to 500 ft.

Athens Shale- Medium- to dark-gray, calcareous, graptolitic shale; calcareous gray sandstone, siltstone, and locally fine-pebble quartz conglomerate; nodules of shaly limestone near base. Maximum thickness 1,500 feet.

Sevier Shale- Calcareous, bluish-gray shale, weathers yellowish-brown; with thin gray limestone layers; sandstone, siltstone, and locally conglomerate to the east. Thickness 2,000 to 7,000 feet.

Mascot Dolomite- Light-gray, fine-grained, well-bedded cherty dolomite; mottled (red and green) dolomite characteristic; interbeds of bluish-gray limestone in upper part; chert-matrix quartz sandstone at base. Erosional unconformity at top. Thickness 350 to 800 feet.

Kingsport Formation- Gray, fine-grained, sparingly cherty dolomite with basal dense, gray limestone sequence. Thickness about 250 feet.

(Ordovician continued on next page.)

Longview Dolomite- Siliceous, gray, fine-grained, medium-bedded dolomite; interbeds of gray limestone in upper part. Thickness about 300 feet.

Chepultepec Dolomite- Light-gray, fine-grained, well-bedded dolomite, moderately cherty; fine-grained limestone locally in upper part; quartz sandstone beds at base. Thickness about 800 feet.

Jonesboro Limestone- Dark bluish-gray, ribboned (silt and dolomite) limestone; numerous interbeds of dark-gray dolomite; quartz sandstone at base. Erosional unconformity at top. Thickness about 2,000 feet.

Wells Creek Dolomite and Knox Group- Yellowish-gray and light olive-gray dolomite with thin partings of grayish-green shale, and pale-orange to yellowish-gray limestone; thin- to thick-bedded, micrograined to coarse-grained. Present only in Wells Creek Basin. Exposed thickness at least 600 ft.

ORDOVICIAN-CAMBRIAN

Mascot Dolomite- Light-gray, fine-grained, well-bedded cherty dolomite; mottled (red and green) dolomite characteristic; interbeds of bluish-gray limestone in upper part; chert-matrix quartz sandstone at base. Erosional unconformity at top. Thickness 350 to 800 feet.

Kingsport Formation- Gray, fine-grained, sparingly cherty dolomite with basal dense, gray limestone sequence. Thickness about 250 feet.

Longview Dolomite- Siliceous, gray, fine-grained, medium-bedded dolomite; interbeds of gray limestone in upper part. Thickness about 300 feet.

Chepultepec Dolomite- Light-gray, fine-grained, well-bedded dolomite, moderately cherty; fine-grained limestone locally in upper part; quartz sandstone beds at base. Thickness about 800 feet.

Copper Ridge Dolomite- Coarse, dark-gray, knotty dolomite, asphaltic in places, with much gray, medium-grained, well-bedded dolomite; abundant chert; cryptozoans typical. Thickness about 1,000 ft.

Conococheague Limestone- Well-bedded, ribboned (silt and dolomite), dark-gray limestone; interbeds of fine-grained, light- to dark-gray dolomite; sparingly cherty; cryptozoans typical. Thickness about 1,500 feet.

CAMBRIAN

C

Copper Ridge Dolomite- Coarse, dark-gray, knotty dolomite, asphaltic in places, with much gray, medium-grained, well-bedded dolomite; abundant chert; cryptozoans typical. Thickness about 1,000 ft.

Conococheague Limestone- Well-bedded, ribboned (silt and dolomite), dark-gray limestone; interbeds of fine-grained, light- to dark-gray dolomite; sparingly cherty; cryptozoans typical. Thickness about 1,500 feet.

Maynardville Limestone- Thick-bedded, bluish-gray, ribboned (silt and dolomite), nodular limestone; light-gray, fine-grained, laminated to thin-bedded, noncherty dolomite in upper part. Thickness 150 to 400 ft.

Nolichucky Shale- Pastel-colored (pink, greenish, olive), flaky clay shale; gray, commonly oolitic, shaly limestone lenses; locally stromatolitic limestone layers; thin, blocky siltstone near middle. Thickness 100 ft in the east to 900 ft in the west.

Maryville Limestone- Gray, ribboned (silt and dolomite), fine-grained, evenly bedded limestone; intraformational conglomerate and oolitic layers common; clay shale and light-gray dolomite locally. Thickness 300 to 800 feet.

Rogersville Shale- Light-green, fissile clay shale; in places limestone (Craig Member) in upper part. Commonly 25 to 80 feet thick; maximum thickness 250 feet.

Rutledge Limestone- Medium- to dark-gray, ribboned (silt and dolomite), medium-grained, well-bedded limestone; locally dark-gray, coarse-grained, medium-bedded dolomite in upper part. Thickness 100 to 500 feet.

Honaker Dolomite- Dark-gray, medium-bedded dolomite with minor dark limestone beds; locally cherty; cryptozoans abundant in places. Thickness about 1,500 feet.

Pumpkin Valley Shale- Dull-brown to maroon shale with numerous interbeds of thin, blocky, sandy siltstone. Thickness 100 to 600 ft.

(Cambrian continued on next page.)

C

Conasauga Group- Mostly shale northwest of a line connecting Knoxville and Tazewell; dominantly dolomite with minor shale southeast of a line from Newport to Kingsport; between these lines it consists six formations. Thickness about 2,000 feet.

Rome Formation- Variegated (red, green, yellow) shale and siltstone with beds of gray, fine-grained sandstone in middle and west part of Valley and Ridge; abundant limestone and dolomite in east. Thickness 1,500-2,000 feet..

Shady Dolomite- Light-gray, well-bedded dolomite with thin- to medium-bedded, gray limestone; yellowish-brown residual clays with "jasperoid" diagnostic. Thickness about 1,000 feet.

Hesse Sandstone- White, vitreous quartzite, medium- to coarse-grained, occurs in massive ledges; Helenmode Member at top is gray to greenish sandstone and shale. Thickness about 600 ft.

Murray Shale- Shale, silty, sandy, dull-green to brown, micaceous. Thickness about 500 ft.

Nebo Sandstone- Medium-bedded, fine-grained, white vitreous quartzite, in part feldspathic. Thickness 250 ft.

Erwin Formation- White, vitreous quartzite, massive with interbeds of dark-green silty and sandy shale, minor siltstone, and very fine-grained sandstone. Thickness 1,000 to 1,500 feet.

Nichols Shale- Olive-gray to green, silty and sandy, micaceous shale and siltstone; local lenses of fine-grained feldspathic quartzite. Thickness about 700 ft.

Cochran Conglomerate- Quartz-pebble conglomerate, gray pebbly arkose, siltstone and shale; irregular bedding, scour features crossbedding common; maroon micaceous arkose and shale near middle and base. Thickness about 1,200 ft.

Hampton Formation- Dark greenish-gray, silty and sandy, micaceous shale; numerous layers of medium-grained, feldspathic, thinly bedded sandstone. Thickness 500 to 2,000 feet.

Unicoi Formation- Sequence of gray feldspathic sandstone, arkose, conglomerate, graywacke, siltstone and shale; greenish amygdaloidal basalt flows near middle and base. Thickness 2,000 to 5,000 feet.

PRECAMBRIAN- PROTEROZOIC "Z"

Sandsuck Formation- Olive-green and gray, argillaceous, micaceous shale with coarse feldspathic sandstone and quartz-pebble conglomerate. Thickness about 2,000 ft.

Wilhite Formation- Gray to green siltstone and slate with interbeds of pebble conglomerate, sandstone, and quartzite. Thickness about 4,000 ft.

Shields Formation- Massive conglomerate, sandstone, argillaceous slate; conglomerate (pebbles of various rock types) characteristic. Thickness about 1,500 ft.

Licklog Formation- Feldspathic sandstone, greenish phyllite, and bluish-gray slate. Thickness about 1,500 ft.

Anakeesta Formation- Dark-gray, bluish-gray, and black slate with dark-gray interbeds of fine-grained sandstone. Thickness 1,500 to 4,500 feet.

Thunderhead Sandstone- Coarse, gray feldspathic sandstone, graywacke, and conglomerate; occurs in massive ledges; graded bedding and blue quartz characteristic. Thickness 5,500 to 6,300 feet.

Elkmont Sandstone- Coarse to fine, gray feldspathic sandstone, graywacke, and fine conglomerate; generally finer grained beds in lower part; graded bedding typical. Thickness 1,000 to 8,000 feet.

Cades Sandstone- Gray, well-bedded, fine- to medium-grained feldspathic metasandstone, with interbeds of dark slate and metasiltstone; precise stratigraphic position unknown. Thickness about 1,500 feet.

Rich Butt Sandstone- Gray, massive beds of feldspathic, fine- to medium-grained sandstone, with interbeds of dark slate and arkosic conglomerate; exact stratigraphic position unknown. Thickness about 1,500 feet.

Unnamed Sandstone Unit- Gray, coarse sandstone and fine conglomerate, similar to Thunderhead Sandstone. Thickness about 4,500 feet.

Pigeon Siltstone- Laminated, greenish quartzose and feldspathic siltstone; minor fine-grained gray sandstone. Thickness as much as 10,000 feet.

Roaring Fork Sandstone- Interbedded massive feldspathic sandstone, greenish siltstone, and greenish phyllite. Maximum thickness 7,000 feet.

(Proterozoic "Z" continued on next page.)

Metcalf Phyllite- Lustrous, pale-green and silvery sericitic and chloritic phyllite; siltstone interbeds abundant. Thickness uncertain; at least 5,000 feet.

Longarm Quartzite- Feldspathic quartzite and arkose, conspicuously light-colored current bedded and crossbedded. Thickness about 5,000 feet.

Wading Branch Formation- Medium- to dark-gray sandy slate to coarse, pebbly feldspathic sandstone and graywacke; basal part is quartz-sericite phyllite; graded bedding common. Thickness about 1,500 feet.

PROTEROZOIC "Y" AND "Z"

Mount Rogers Group- Metavolcanics, typically purplish and reddish; massive lavas and tuffs, altered rhyolites and quartz latites; strongly foliated; interbedded arkose, shale, and conglomerate. Thickness 1,000 to 3,000 feet.

Bakersville Gabbro- Metagabbro, dark, porphyritic; contains diorite, basalt, anorthosite, and diabase; occurs as thin to massive dikes and lenticular masses.

Beech Granite- Granite, porphyritic, light-gray to reddish; coarse potash feldspar crystals with clustered interstitial mafics (chloritized biotite and hornblende) give spotted appearance; includes Max Patch Granite.

Cranberry Granite- Complex of intertonguing rock types including migmatite, granitic gneisses, monzonite, quartz diorite, green stone, mica and hornblende schists, abundant granitic pegmatite.

Roan Gneiss- Layered hornblende and garnet gneiss and granitic migmatite with zones of mica schist and amphibolite, foliation commonly contorted; contains numerous granitic and gabbroic dikes.

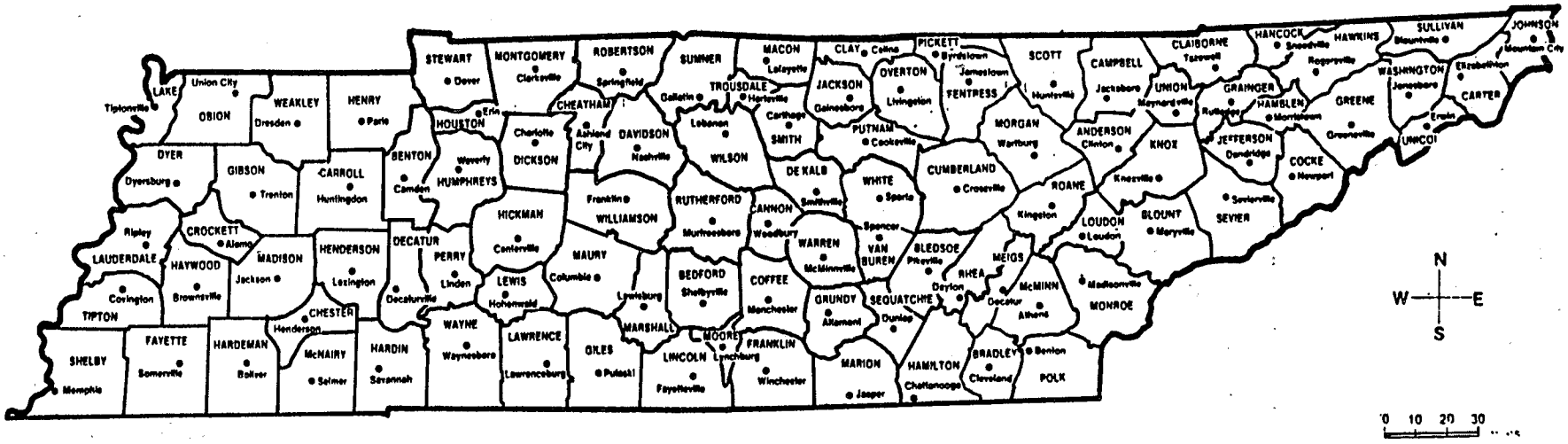

Figure 5- Map showing counties in Tennessee (modified from Facts on File, 1984).

An aeroradiometric map of Tennessee (fig. 6) compiled from NURE flightline data (Duval and others, 1989) shows that low radioactivity (<1.5 ppm eU) is associated with Pennsylvanian sandstones and sandy shales of the Appalachian Plateau. Moderate radioactivity (1.5-2.5 ppm eU) is found throughout the State. Areas of high radioactivity (> 2.5 ppm eU) are found in the northern part of the loess-covered Coastal Plain, in an outcrop belt of Cretaceous sedimentary rocks containing heavy mineral sands in the eastern Coastal Plain, in the Nashville Basin and the eastern part of the Highland Rim, the Valley and Ridge, and parts of the Unaka Mountains. The areas of high radioactivity are most commonly associated with black shales, phosphatic carbonate rocks of Paleozoic age, and granites.

SOILS

Most Tennessee soils lie in the thermic udic soil moisture-temperature regime (Rose and others, 1991). Soils in the the Appalachian Plateau Province and the northern part of the Highland Rim Province lie in the mesic udic regime. Thermic udic soils are very moist in the winter (56-96 percent pore saturation in sandy loams, and 74-99 percent saturation in silty clay loams) and are slightly moist in the summer (24-44 percent pore saturation in sandy loams, and 39-58 percent pore saturation in silty clay loams). Mesic udic soils are very moist (56-96 percent pore saturation in sandy loams, and 74-99 percent saturation in silty clay loams) in the winter and are moderately moist (44-56 percent saturation in sandy loams, and 58-74 percent in silty clay loams) in the summer.

In places where soils are moderately moist to very moist, soil moisture will tend to inhibit radon migration by diffusion and flow. However, soils in which the water drains rapidly from the soil profile because of high intrinsic permeability, steep slopes, or both, may be areas in which radon may migrate more readily and the radon potential of that area is increased. Conversely, soils in which the water drains away slowly because of low intrinsic permeability, low slopes, or both, may be areas where radon migrates very slowly, and the radon potential is lowered.

On the Mississippi Alluvial Plain most soils are seasonally wet and gently sloping. They have thin organic surface layers, although thick organic surface layers are present locally. In the subsurface the soils are poorly differentiated. Many soils are mottled. Mineral matter is generally being lost from the profiles. Accumulations of clays and Fe-Mn minerals in subsurface layers occur only locally. Some soils have developed on recently deposited alluvial sand, silt, and clay and have no discernable pedogenic development. The soils are mostly clayey to silty and are poorly drained in bottomlands away from the river, and clayey to locally sandy and poorly to locally excessively drained near the river (Springer and Elder, 1980).

Soils developed on the Gulf Coastal Plain are usually moist and gently to moderately sloping. These soils are generally high in bases in the western loess-covered part of the area, but are more acid in the east. Surface layers are gray to brown. Thin clay layers have developed in the subsurface and, locally, dense, brittle layers of fragipan have formed. In the eastern part of the area, thick clay layers have formed locally in the subsurface. Soils throughout the loess-covered area are generally well drained or moderately well drained. Soils developed on the Tertiary sedimentary rocks are generally well drained. Those soils in the latter area that contain abundant clay are well drained because they tend to be formed on slopes.

Soils in the western part of the Highland Rim area are usually moist with no or very short dry periods. Slopes are gentle to locally steep. Organic matter is rare in the subsurface and thick clayey horizons are common, although locally they are thinner and have associated dense brittle

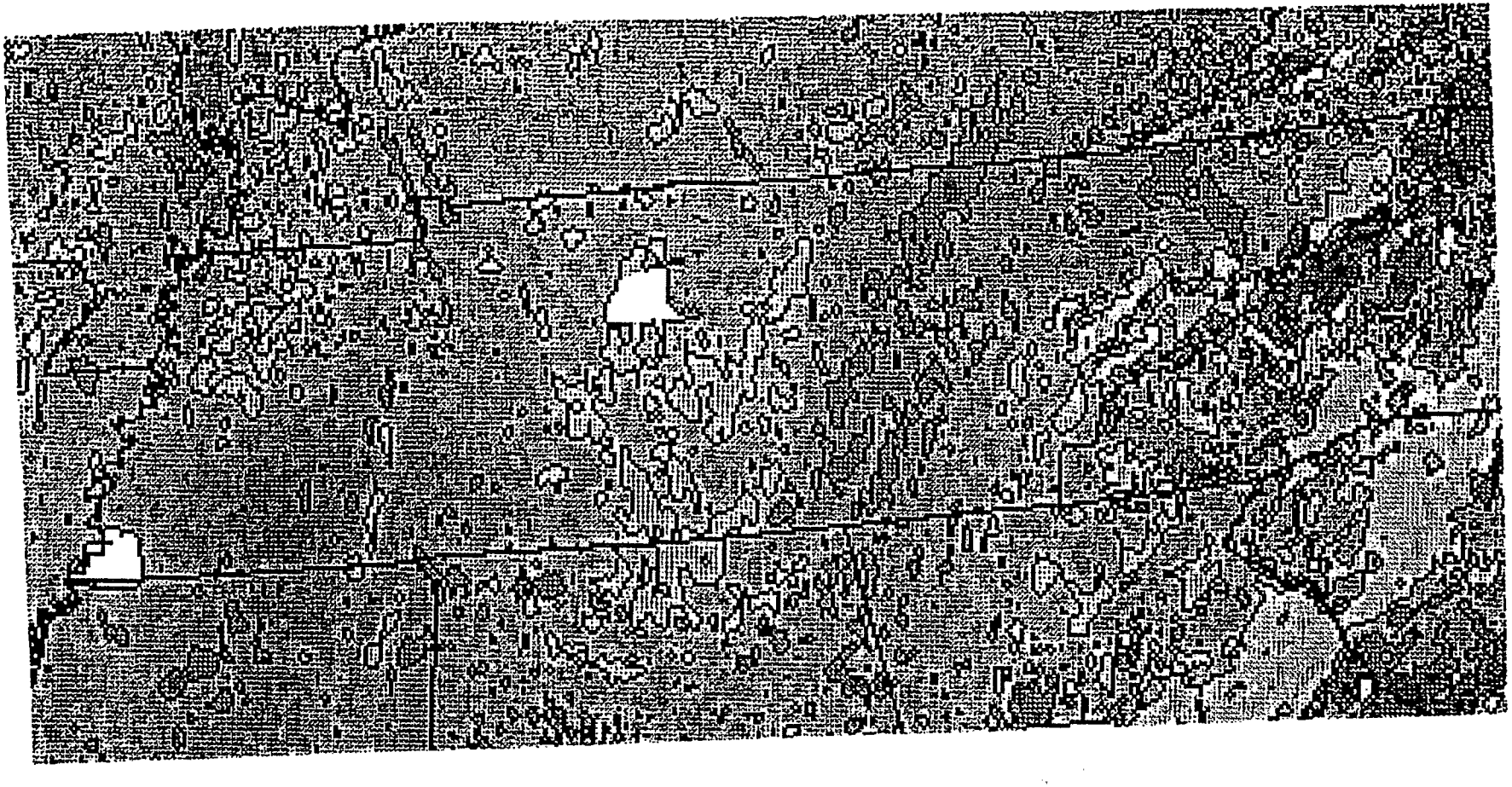

Figure 6- Aeroradiometric map of Tennessee (after Duval and others, 1989). Contour lines at 1.5 and 2.5 ppm equivalent uranium (eU). Pixels shaded at 0.5 ppm eU increments.

fragipan layers. Soils developed on cherty limestone are steep and excessively drained, whereas those developed on limestone or thin loess are well drained and silty. The Bodine soil map unit, which occurs on steep slopes in widespread areas of the western half of the Highland Rim, is excessively drained and very rapidly permeable. In the eastern part of the Highland Rim, soils are usually moist, but some soils dry out during warm periods of the year. Thin clays accumulate in the subsurface. Some steep, excessively drained soils developed on cherty limestone occur in a thin belt, but well drained silty to clayey soils developed on shale, siltstone, limestone, and thin loess are most common. Well drained silty soils developed on alluvium occur adjacent to the Cumberland Plateau area (Springer and Elder, 1980).

The Nashville Basin is characterized by gentle to moderate slopes and soils that are usually moist except during extended dry periods. Soils are dominated by thick red clay horizons except in some areas where clay horizons are thin. Rocky areas with thin or no soils occur locally. Well-drained silty to clayey soils developed on phosphatic limestones rim the center part of the Nashville Basin (fig. 7). Cherty soils occur locally in this area. The central part of the Nashville Basin has well drained, clayey to locally silty soils developed on limestone.

Soils of the Appalachian Plateau are moderately sloping in the high plateau areas and steeply sloping in the areas transitional to nearby valleys. The soils are dominated by loams and stony loams developed from sandstone, shale and, locally, limestone. Soils developed on colluvium are common near the base of steeply sloping areas. Sandstone outcrops are common in steep terrains. The soils are very acid and typically well drained. In the Sequatchie Valley, well-drained clayey to loamy soils developed on alluvium and colluvium dominate.

Soils of the Appalachian Ridge and Valley region are highly varied because of differences in underlying rock types and topography. Slopes are moderate in most valley areas and steep on the ridges. The area is underlain variously by 1) deep, cherty to clayey soils and deep, red, clayey soils developed from dolomitic limestone; 2) deep, locally red to dark red, clayey and loamy soils developed from limestone, alluvium, and colluvium; 3) shallow to moderately deep shaley, clayey, and loamy soils developed from calcareous shale; 4) shallow to moderately deep clayey to loamy soils developed from limestone and shale; and 5) stony, loamy, and clayey soils from sandstone, shale, and limestone (Springer and Elder, 1980). The soils are generally well drained, however some of the soils derived from sandstone on ridges north and northeast of Knoxville are excessively drained.

Soils of the Unaka Mountains are also highly varied. Most of the soils are formed on steep slopes. The soils include: 1) loamy and stony soils derived from metamorphic and igneous rocks and colluvium; 2) stony, loamy, and channery soils derived from phyllite, slate, shale, sandstone, quartzite, and colluvium; and 3) loamy to clayey soils derived from colluvium, alluvium, shale and limestone (Springer and Elder, 1980). All of these soils are well drained and some have high intrinsic permeability.

Because of the abundance of clay- and silt-sized materials and the generally moist conditions of soils throughout Tennessee, soil permeability is generally moderate to low throughout the State. However, most soils throughout the State tend to be well drained because most soils are sloping. High intrinsic soil permeability is likely to occur in areas of steep, well-drained, sandy soils derived from weathering of sandstone in the Appalachian Plateau and some areas of the Valley and Ridge, cherty soils developed on cherty limestone in the Highland Rim or soils developed on alluvial sand and gravel deposits in areas along streams and rivers. Some soils developed on karst terrains in Tennessee may have high permeability where pinnacles of bedrock occur locally.

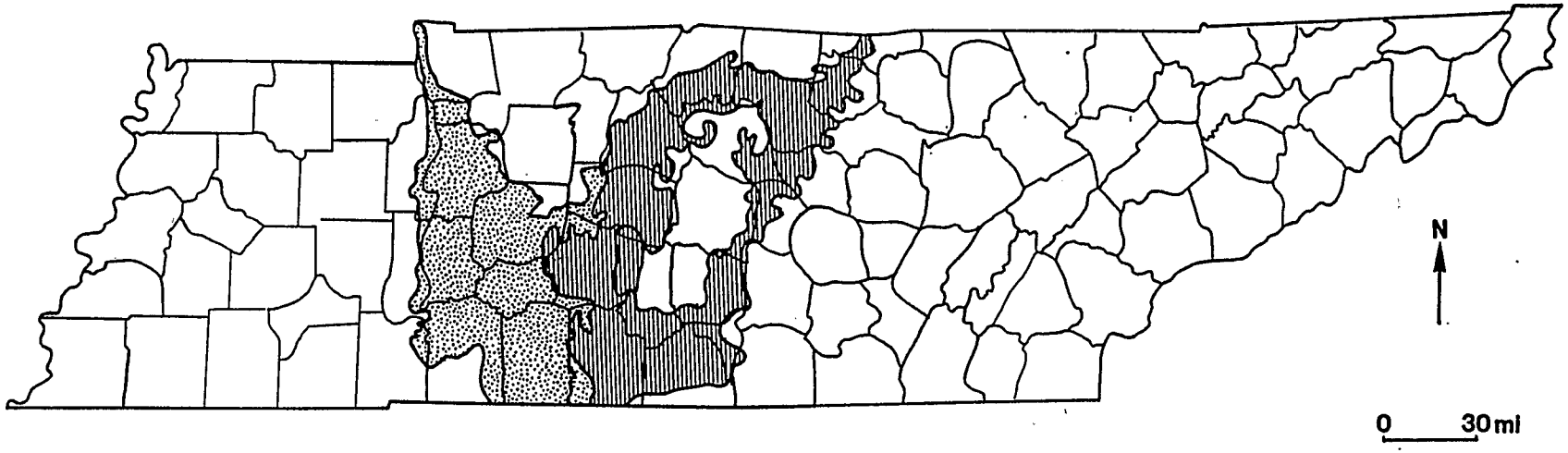

Figure 7- Map showing distribution of soil associations where the excessively drained Bodine soils in the Highland Rim are a princpal part of the mapped soils (dotted areas) and where phosphatic soils of the Nashville Basin form a major part of the mapped soils (lined pattern). Many areas of river bottomlands and inclusions of other mapped soil associations are not shown on this map. Derived from Springer and Elder (1980). The reader should refer to Springer and Elder (1980) or to individual county soil surveys for more complete descriptions of radon-prone soil units and maps of their distribution.

Phosphate enrichment has occurred in many soils developed on certain Precambrian, Cambrian, Ordovician, and Mississippian carbonate and shale units. Phosphate-enriched material was mined for fertilizer in the early 1900s in various parts of Tennessee and is still mined in the Nashville Basin. Because phosphate precipitation in the original limestone and shale units was commonly accompanied by uranium precipitation, these residual soils are also similarly enriched in uranium. Geologic units of concern and areas with past phosphate mining include: a) the Mississippian Maury Formation, Smith County (Nashville Basin); b) Precambrian Yellow Breeches Member, Wilhite Formation, Sevier County (Unaka Mountains); c) Cambrian Shady and Rome Formations, Johnson, Unicoi, Greene, and Cocke Counties (Unaka Mountains); d) Cambrian Helenmode Formation, Johnson County (Unaka Mountains); and e) Ordovician limestones and shales, Blount, De Kalb, Smith, Jackson, and Macon Counties (Unaka Mountains, Nashville Basin). Phosphatic soils underlie much of the outer edge of the Nashville Basin. The distribution of soil associations in which phosphatic soils form a significant part of the association are shown in figure 7.

INDOOR RADON DATA

The Tennessee Department of Health and Environment and the U.S. EPA conducted a population-based survey of indoor radon levels in 1773 homes in Tennessee during the winter of 1987-88 (fig. 8, Table 1). In figure 8, data are shown only for those counties with 5 or more data values in the State/EPA Residential Radon Survey. Geologic interpretations of population-based data must be made with caution because the measured houses are typically only from a relatively few population centers in a given county and the distribution of these houses do not necessarily reflect the variation in geology in the county. For example, a county may have a relatively high radon potential on well-drained, uraniferous soils on hillslopes that occur over a widespread area, but if housing is generally located on poorly drained soils with low uranium contents on the valley floor, a population-based survey for that area will contain relatively low indoor radon values.

The maximum indoor radaon value recorded in the survey was 99.9 pCi/L in Roane County. Maximum values exceeding 20 pCi/L occur in counties within the Ridge and Valley region, the Nashville Basin, and the Highland Rim. The average (arithmetic mean) for the State was 3.0 pCi/L and the 18.2 percent of the homes tested had indoor radon levels exceeding 4 pCi/L.

Some persistent patterns are present in the survey data. Average indoor radon levels less than 2 pCi/L occur in counties throughout the Mississippi Alluvial Plain and the Coastal Plain, except in Henry County, where a single value of 11.9 pCi/L raised the county average from 1.1 to 2.0 pCi/L (N=12), and in Fayette County where a 4.2 pCi/L value raised the average to 2.3 pCi/L (N=2). There are very few homes with basements in these two physiographic regions and only a few basement measurements were made in the Memphis area.

Average indoor radon levels for counties within the Highland Rim range from less than 1 pCi/L to 13.8 pCi/L; however, most counties average 1.0-3.0 pCi/L. Maximum values are 54.5 pCi/L in White County and 76.8 pCi/L in Hickman County. Removal of these single high values from the datasets of both counties lowers the county averages from 13.8 to 2.9 pCi/L and 8.0 to 1.7 pCi/L, respectively. In the Nashville Basin, the county averages tend to be somewhat higher with most counties averaging 2-5 pCi/L. Values exceeding 4 pCi/L in these two regions can largely be attributed to radium-rich soils developed on normal and phosphatic carbonate rocks, to karstic topography, and to soils developed on uraniferous shales.

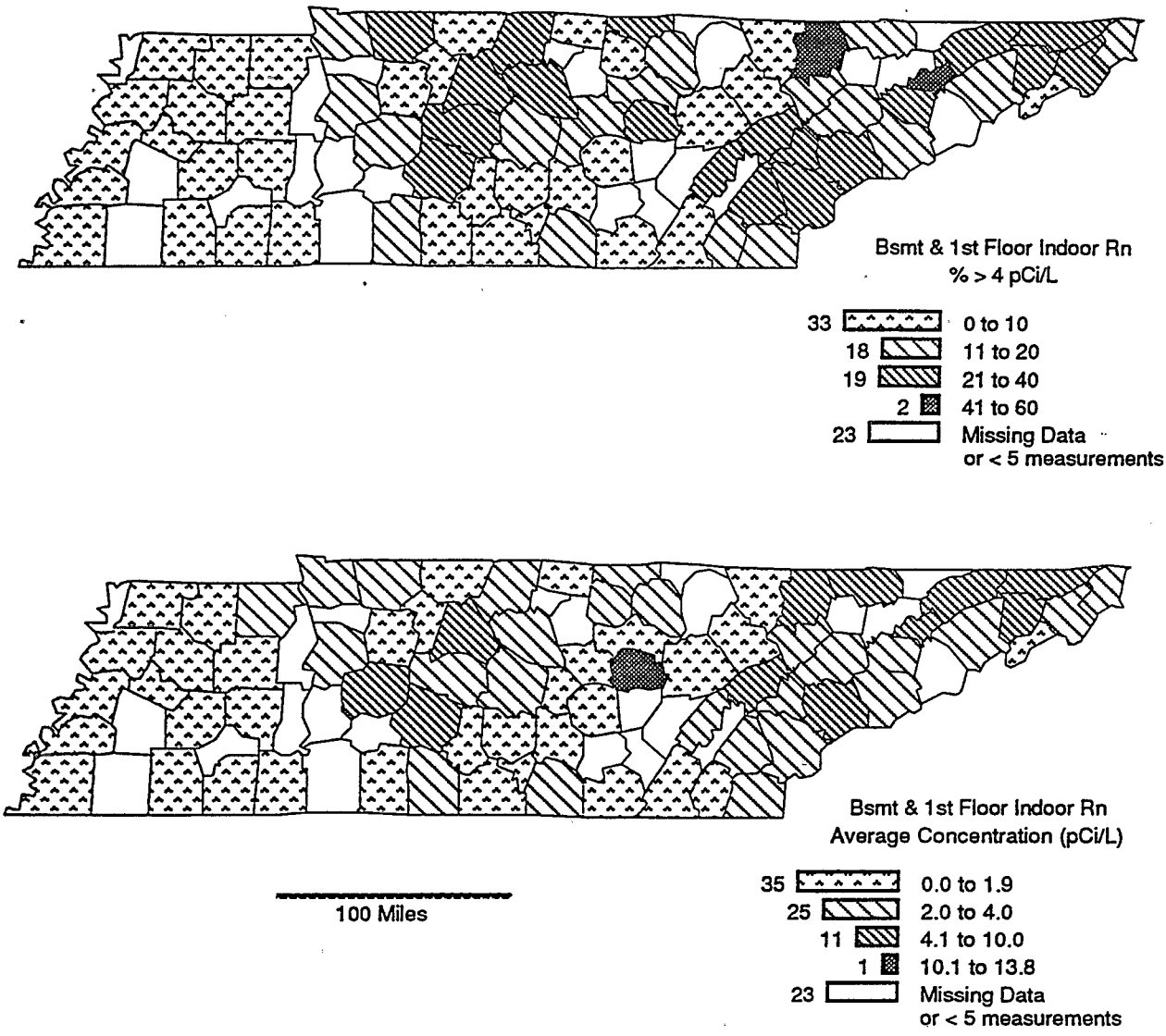

Figure 8. Screening indoor radon data from the EPA/State Residential Radon Survey of Tennessee, 1987-88, for counties with 5 or more measurements. Data are from 2-7 day charcoal canister tests. Histograms in map legends show the number of counties in each category. The number of samples in each county (See Table 1) may not be sufficient to statistically characterize the radon levels of the counties, but they do suggest general trends. Unequal category intervals were chosen to provide reference to decision and action levels.

TABLE 1. Screening indoor radon data from the EPA/State Residential Radon Survey of Tennessee conducted during 1986-87. Data represent 2-7 day charcoal canister measurements from the lowest level of each home tested.

COUNTY	NO. OF MEAS.	MEAN	GEOM. MEAN	MEDIAN	STD. DEV.	MAXIMUM	%>4 pCi/L	%>20 pCi/L
ANDERSON	35	3.0	1.5	1.7	4.1	21.2	20	3
BEDFORD	10	1.5	1.2	1.3	1.3	4.9	10	0
BENTON	2	1.0	0.6	1.0	1.1	1.7	0	0
BLEDSOE	1	0.4	0.4	0.4	0.0	0.4	0	0
BLOUNT	40	4.1	2.4	2.5	5.1	23.1	23	5
BRADLEY	35	1.7	1.0	1.2	1.5	5.7	11	0
CAMPBELL	8	5.3	3.5	4.0	5.5	18.1	50	0
CANNON	7	1.7	1.2	1.2	1.4	4.5	14	0
CARROLL	9	1.2	1.1	1.2	0.6	2.2	0	0
CARTER	30	3.7	2.4	3.0	3.4	16.1	33	0
CHEATHAM	15	1.6	0.9	1.1	1.5	5.6	7	0
CHESTER	1	0.6	0.6	0.6	0.0	0.6	0	0
CLAIBORNE	10	4.3	2.6	2.6	5.5	19.2	20	0
CLAY	9	2.2	1.6	1.4	2.0	6.8	22	0
COCKE	4	2.2	2.1	2.3	0.5	2.7	0	0
COFFEE	30	1.6	1.1	1.1	1.5	6.6	10	0
CROCKETT	6	0.7	0.5	0.5	0.6	1.6	0	0
CUMBERLAND	5	1.6	1.1	1.3	1.4	3.9	0	0
DAVIDSON	245	5.1	2.3	2.0	8.6	64.2	30	5
DE KALB	8	1.6	1.2	0.8	1.4	4.3	13	0
DECATUR	3	0.3	0.2	0.1	0.3	0.7	0	0
DICKSON	12	1.7	1.1	1.1	2.1	8.0	8	0
DYER	6	0.9	0.8	1.0	0.5	1.6	0	0
FAYETTE	2	2.3	1.3	2.3	2.7	4.2	50	0
FENTRESS	4	1.2	0.8	1.2	1.0	2.2	0	0
FRANKLIN	26	2.2	1.6	1.9	2.0	9.9	12	0
GIBSON	13	0.7	0.6	0.5	0.3	1.1	0	0
GILES	14	2.1	1.6	1.5	1.4	4.5	7	0
GRAINGER	3	1.1	0.6	0.3	1.3	2.6	0	0
GREENE	20	2.2	1.4	1.1	2.3	8.7	20	0
GRUNDY	3	0.9	0.8	1.1	0.3	1.1	0	0
HAMBLEN	22	7.0	3.2	3.6	12.4	59.0	45	5
HAMILTON	120	1.6	1.0	1.2	1.7	8.0	7	0
HANCOCK	4	2.3	2.0	1.9	1.4	4.2	25	0
HARDEMAN	5	0.9	0.7	0.5	0.9	2.6	0	0
HARDIN	6	1.1	1.0	1.1	0.6	1.8	0	0
HAWKINS	18	4.5	2.3	2.2	6.0	24.0	28	6
HAYWOOD	4	0.5	0.4	0.5	0.3	0.8	0	0
HENDERSON	9	0.8	0.7	0.7	0.3	1.4	0	0
HENRY	12	2.0	1.0	1.0	3.2	11.9	8	0
HICKMAN	12	8.0	2.0	1.8	21.7	76.8	17	8

TABLE 1 (continued). Screening indoor radon data for Tennessee.

COUNTY	NO. OF MEAS.	MEAN	GEOM. MEAN	MEDIAN	STD. DEV.	MAXIMUM	%>4 pCi/L	%>20 pCi/L
HOUSTON	2	0.8	0.8	0.8	0.3	1.0	0	0
HUMPHREYS	6	2.1	1.3	1.5	2.2	6.3	17	0
JACKSON	11	2.5	2.1	2.0	1.5	5.1	9	0
JEFFERSON	13	3.7	2.3	2.7	4.2	15.9	31	0
JOHNSON	6	3.4	1.8	1.6	4.9	13.2	17	0
KNOX	131	2.6	1.7	1.9	3.1	26.6	18	1
LAKE	2	1.9	1.8	1.9	1.0	2.6	0	0
LAUDERDALE	8	0.7	0.5	0.8	0.5	1.4	0	0
LAWRENCE	6	1.9	1.4	1.3	1.9	5.7	17	0
LINCOLN	16	1.1	0.5	0.5	1.3	4.1	6	0
LOUDON	13	2.8	2.2	2.4	1.9	6.5	23	0
MACON	10	1.4	1.1	1.2	1.2	4.5	10	0
MADISON	15	0.9	0.7	0.7	0.7	2.8	0	0
MARION	5	1.2	1.0	0.8	0.9	2.8	0	0
MARSHALL	5	1.1	1.0	1.0	0.6	2.0	0	0
MAURY	39	4.1	2.6	3.1	5.3	33.2	38	3
MCMINN	18	2.7	1.7	2.0	2.3	8.8	28	0
MCNAIRY	5	0.5	0.4	0.4	0.2	0.7	0	0
MEIGS	3	2.2	1.7	1.5	1.9	4.4	33	0
MONROE	9	3.0	2.1	2.7	2.3	7.6	22	0
MONTGOMERY	18	2.8	1.7	1.5	2.6	8.8	33	0
MOORE	5	1.3	0.7	1.5	1.0	2.7	0	0
MORGAN	5	1.5	1.3	1.6	0.6	2.2	0	0
OBION	6	0.7	0.5	0.6	0.6	1.7	0	0
OVERTON	5	2.6	2.1	1.9	1.7	4.8	20	0
PERRY	1	0.6	0.6	0.6	0.0	0.6	0	0
POLK	8	2.5	1.6	1.4	3.3	10.4	13	0
PUTNAM	27	1.8	1.2	1.4	1.4	4.9	11	0
RHEA	13	3.6	2.2	2.4	4.6	18.1	31	0
ROANE	22	7.1	1.8	1.5	21.0	99.9	23	5
ROBERTSON	12	1.2	1.0	0.9	1.2	4.8	8	0
RUTHERFORD	23	3.1	2.0	1.6	4.1	17.5	17	0
SCOTT	5	0.5	0.4	0.5	0.2	0.6	0	0
SEQUATCHIE	2	2.8	2.6	2.8	1.6	3.9	0	0
SEVIER	16	3.2	2.0	2.4	3.8	16.5	13	0
SHELBY	144	1.0	0.8	0.8	0.8	4.4	1	0
SMITH	2	3.8	3.7	3.8	1.1	4.5	50	0
STEWART	5	2.4	2.0	2.2	1.3	4.3	20	0
SULLIVAN	73	5.6	2.7	2.7	9.4	67.3	36	5
SUMNER	70	2.7	1.6	1.9	2.9	17.0	24	0
TIPTON	6	1.1	0.7	1.0	0.9	2.7	0	0
TROUSDALE	3	2.2	1.4	0.8	2.5	5.1	33	0
UNICOI	14	1.9	1.6	1.8	1.1	4.9	7	0

TABLE 1 (continued). Screening indoor radon data for Tennessee.

COUNTY	NO. OF MEAS.	MEAN	GEOM. MEAN	MEDIAN	STD. DEV.	MAXIMUM	%>4 pCi/L	%>20 pCi/L
UNION	3	2.2	1.5	1.3	2.3	4.8	33	0
WARREN	10	1.9	1.3	1.4	1.5	4.3	10	0
WASHINGTON	35	4.1	2.5	2.6	4.7	24.9	29	3
WAYNE	4	1.1	1.0	1.2	0.4	1.4	0	0
WEAKLEY	7	0.9	0.8	0.7	0.6	2.1	0	0
WHITE	5	13.8	5.0	3.0	22.9	54.5	40	20
WILLIAMSON	56	3.1	2.2	2.3	2.7	15.6	25	0
WILSON	17	2.6	1.0	0.7	3.1	9.3	29	0

Counties in the Appalachian Plateau are sparsely sampled, but county indoor radon averages range from 0.4 to 5.3 pCi/L. County averages in the Ridge and Valley region range from 1.6 pCi/L to 7.0 pCi/L with about 1/3 of the counties having indoor radon averages exceeding 4.0 pCi/L. County averages in the Unaka Mountains are hard to characterize because only Johnson and Unicoi Counties, which average 3.4 and 1.9 pCi/L, respectively, lie exclusively in the Unaka Mountains. All other counties lie astride the Unaka Mountains and the Ridge and Valley regions and most samples are inferred to have come from the more highly populated Ridge and Valley.

GEOLOGIC RADON POTENTIAL

A study of soil-gas radon in the Coastal Plain of Tennessee and other states (Gundersen and Peake, 1992) suggests that soils developed on loess and glauconitic, phosphatic, and carbonaceous sedimentary rocks produce average measured soil-gas radon concentrations of 2200 pCi/L (range 120-3950 pCi/L); however, 34 of 59 sample sites were so moist that no sample could be drawn. Soil-gas radon levels above 2000 pCi/L tend to produce significant percentages of homes with more than 4 pCi/L radon indoors in other areas of the United States. However, loess and the clayey soils common in this area tend to have low permeability, especially under the soil moisture conditions that prevail in much of Tennessee, so even though these sediments may be a possible source of high radon levels, slow permeability probably inhibits radon availability.

The Mississippian and Devonian Chattanooga Shale is a source of high indoor radon levels in Kentucky (Peake and Schumann, 1991). Glacial soils with fragments of the uraniferous Ohio Shale, an equivalent formation in central Ohio, produce a high percentage of homes above 4 pCi/L (80-90 percent) and levels as much as 200 pCi/L indoors in the Columbus, Ohio area. In various areas of Tennessee, the Chattanooga Shale forms wide outcrop (eastern Highland Rim) or underlies small valleys (eastern Tennessee). In many areas in central Tennessee it occurs on hillslopes capped by more resistant units such as the Fort Payne chert and thus tends to be covered by colluvial material. Because of its high swelling clay content, the Chattanooga Shale provides poor foundation conditions for structures and historically it has been less developed. However, housing pressures are forcing many developers to build in areas not previously considered. Structures sited on the Chattanooga Shale are very likely to have elevated indoor radon levels, especially where fractures increase the bedrock permeability and slopes tend to increase drainage and keep soils drier. Such structures may locally have indoor radon levels exceeding 200 pCi/L.

Soils developed from carbonate rocks are often elevated in uranium and radium. When the carbonate minerals in the original rock dissolve away, the soils are enriched in the remaining clay and iron oxides which collect impurities including base metals, uranium, and radium. The accumulation of uranium is strongly enhanced where the carbonate rocks are phosphatic because phosphatic carbonate rocks contain more uranium initially, and the phosphate and associated uranium concentrate readily in the residual soils. Karst terrains that develop on carbonate rocks also enhance radon potential because the bedrock contains numerous solution openings that accumulate radon and increase the bedrock permeability. Carbonate soils derived from Cambrian-Ordovician rock units of the Valley and Ridge Province cause indoor radon problems in eastern Tennessee (Goldsmith and others, 1983), western New Jersey, western Virginia, eastern West Virginia (Schultz and others, 1992), and central and eastern Pennsylvania (Greeman and others, 1990; Sachs and others, 1982).

SUMMARY

For the purpose of this assessment, Tennessee has been divided into nine geologic radon potential areas and each area assigned a Radon Index (RI) and a Confidence Index (CI) score (fig. 9, Table 2). The RI is a semi-quantitative measure of radon potential based on geology, soils, radioactivity, architecture, and indoor radon. The CI is a measure of the relative confidence of the RI assessment based on the quality and quantity of the data used to assess geologic radon potential (see the introduction chapter to this regional booklet for more information).

The Mississippi Alluvial Plain has low geologic radon potential. The high soil moisture, high water table, and the lack of permeable soils lower the radon potential in spite of moderate eU values. Some areas with very sandy or excessively drained soils may cause structures to have indoor radon values over 4 pCi/L.

The loess-covered parts of the Coastal Plain have geologic low radon potential in spite of moderate eU values and elevated soil-gas radon concentrations. The potential is lowered by the high moisture content, low soil permeability, and lack of basements in homes. If prolonged dry periods were to occur in this area, some homes might see a significant increase in indoor radon, especially those with basements or crawl spaces. The eastern Coastal Plain has moderate radon potential. NURE data show elevated eU values compared to the rest of the Coastal Plain. Soil-gas radon levels are locally elevated.

The part of the Highland Rim underlain by cherty limestone (Fort Payne Formation) has high radon potential. This area has moderate to locally high eU and soils that are cherty and well drained. The limestone and shale part of the Highland Rim has moderate radon potential.

The Nashville Basin has high radon potential. The elevated eU, the presence of abundant phosphatic soils, local karst, and the presence of generally well drained soils all contribute to this geologic radon potential. Very high (>20 pCi/L) to extreme indoor radon values (>200 pCi/L) are possible where homes are sited on soils developed on the Chattanooga shale, on phosphate-rich residual soils, or karst pinnacles.

The Appalachian Plateau has moderate radon potential. Values for eU are among the lowest in the State, but many sandy, well drained to excessively drained soils are present.

The Ridge and Valley region has high radon potential because of elevated eU values, karst, and well drained soils. Very high to extreme indoor radon levels are possible where homes are sited on soils developed on black shales, on phosphate-rich residual soils, or karst pinnacles. Home with basements are more likely to yield elevated indoor radon levels.

The Unaka Mountains have moderate radon potential. Values of eU are generally moderate to locally high. Some very high (>20 pCi/L) to extreme (>200 pCi/L) indoor radon levels are possible where homes are sited on phosphate-rich residual soils developed on phosphatic carbonate rocks or pegmatite in the metamorphic rock areas, but the former are much less common in this region than in the Nashville Basin and the Ridge and Valley region.

This is a generalized assessment of the State's geologic radon potential and there is no substitute for having a home tested. The conclusions about radon potential presented in this report cannot be applied to individual homes or building sites. Indoor radon levels, both high and low, can be quite localized, and within any radon potential area there will likely be areas with higher or lower radon potential than assigned to the area as a whole. Any local decisions about radon should not be made without consulting all available local data. For additional information on radon and how to test, contact your State radon program or EPA regional office. More detailed information on state or local geology may be obtained from the State geological survey. Addresses and phone numbers for these agencies are listed in chapter 1 of this booklet.

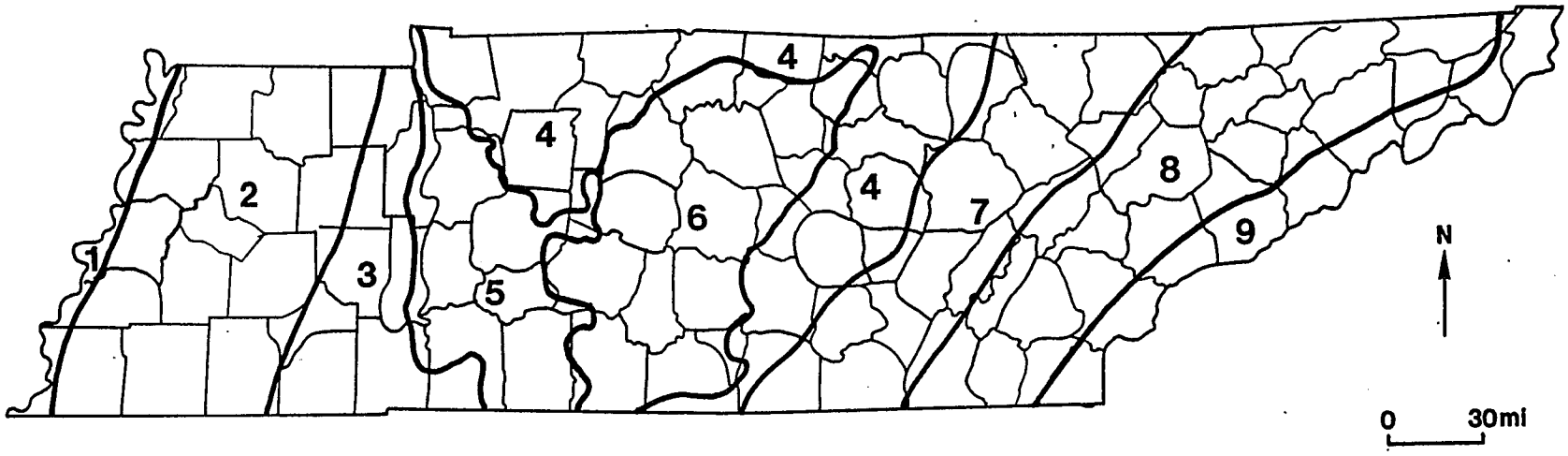

Figure 9- Radon potential areas in Tennessee. (1) Mississippi Alluvial Plain; (2) Loess-covered Coastal Plain; (3) East Coastal Plain; (4) Highland Rim; (5) Cherty Highland; (6) Nashville Basin; (7) Appalachian Plateau; (8) Ridge and Valley; (9) Unaka Mountains.

TABLE 2. Radon Index (RI) and Confidence Index (CI) scores for geologic radon potential areas of Tennessee. See figure 9 for locations of areas.

FACTOR	Mississippi alluvial plain RI	CI	Loess-covered Coastal Plain RI	CI	Eastern Coastal Plain RI	CI
INDOOR RADON	1	3	1	3	1	3
RADIOACTIVITY	2	3	2	3	2	3
GEOLOGY	1	2	2	3	2	3
SOIL PERM.	2	3	2	3	2	3
ARCHITECTURE	1	-	1	-	1	-
GFE POINTS	0	-	0	-	1	-
TOTAL	7	11	8	12	9	12
RANKING	LOW	HIGH	LOW	HIGH	MOD	HIGH

FACTOR	Highland Rim RI	CI	Cherty Highland RI	CI	Nashville Basin RI	CI
INDOOR RADON	2	3	2	3	2	3
RADIOACTIVITY	2	3	2	3	3	3
GEOLOGY	3	2	3	2	3	3
SOIL PERM.	2	3	3	3	2	3
ARCHITECTURE	1	-	1	-	2	-
GFE POINTS	1	-	1	-	2	-
TOTAL	11	11	12	11	14	12
RANKING	MOD	HIGH	HIGH	HIGH	HIGH	HIGH

FACTOR	Appalachian Plateau RI	CI	Ridge and Valley RI	CI	Unaka Mountains RI	CI
INDOOR RADON	1	3	2	3	2	3
RADIOACTIVITY	2	3	3	3	2	3
GEOLOGY	2	3	3	3	2	2
SOIL PERM.	2	3	2	3	2	3
ARCHITECTURE	1	-	2	-	3	-
GFE POINTS	0	-	2	-	0	-
TOTAL	8	12	14	12	11	11
RANKING	LOW	HIGH	HIGH	HIGH	MOD	HIGH

RADON INDEX SCORING:

Radon potential category	Point range	Probable screening indoor radon average for area
LOW	3-8 points	< 2 pCi/L
MODERATE/VARIABLE	9-11 points	2 - 4 pCi/L
HIGH	> 11 points	> 4 pCi/L

Possible range of points = 3 to 17

CONFIDENCE INDEX SCORING:

LOW CONFIDENCE	4 - 6 points
MODERATE CONFIDENCE	7 - 9 points
HIGH CONFIDENCE	10 - 12 points

Possible range of points = 4 to 12

REFERENCES CITED IN THIS REPORT AND GENERAL REFERENCES PERTAINING TO RADON IN TENNESSEE

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EPA's Map of Radon Zones

The USGS' Geologic Radon Province Map is the technical foundation for EPA's Map of Radon Zones. The Geologic Radon Province Map defines the radon potential for approximately 360 geologic provinces. EPA has adapted this information to fit a county boundary map in order to produce the Map of Radon Zones.

The Map of Radon Zones is based on the same range of predicted screening levels of indoor radon as USGS' Geologic Radon Province Map. EPA defines the three zones as follows: Zone One areas have an average predicted indoor radon screening potential greater than 4 pCi/L. Zone Two areas are predicted to have an average indoor radon screening potential between 2 pCi/L and 4 pCi/L. Zone Three areas are predicted to have an average indoor radon screening potential less than 2 pCi/L.

Since the geologic province boundaries cross state and county boundaries, a strict translation of counties from the Geologic Radon Province Map to the Map of Radon Zones was not possible. For counties that have variable radon potential (i.e., are located in two or more provinces of different rankings), the counties were assigned to a zone based on the predicted radon potential of the province in which most of its area lies. (See Part I for more details.)

TENNESSEE MAP OF RADON ZONES

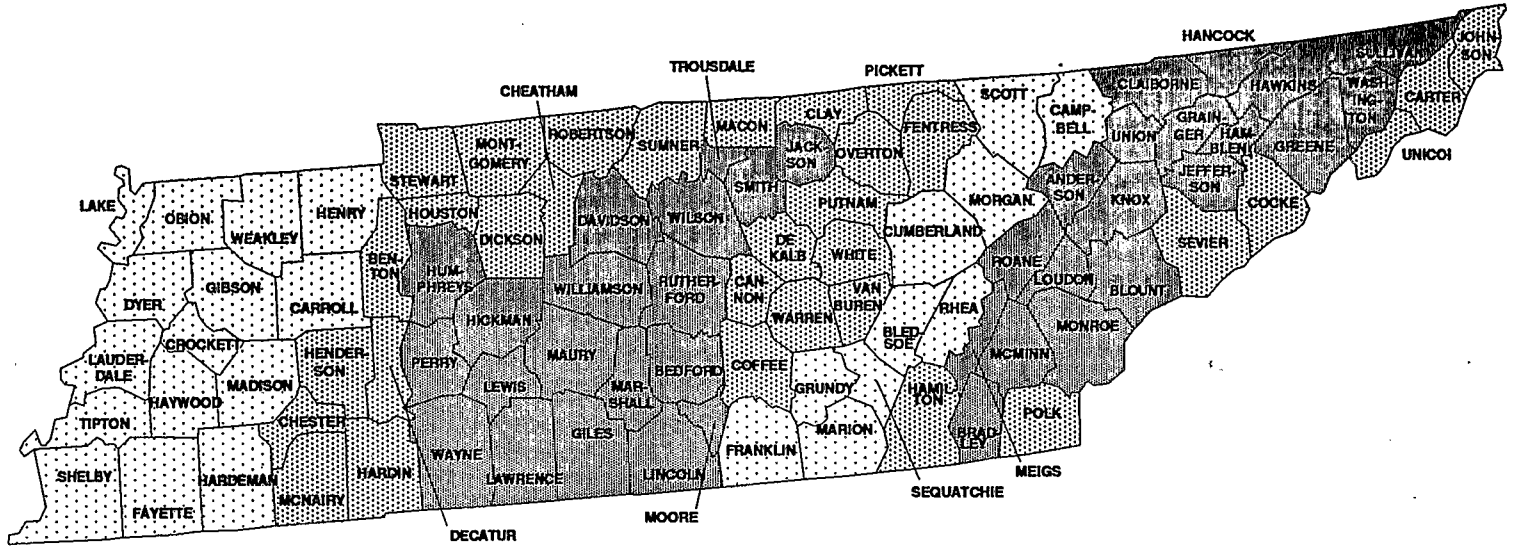
The Tennessee Map of Radon Zones and its supporting documentation (Part IV of this report) have received extensive review by Tennessee geologists and radon program experts. The map for Tennessee generally reflects current State knowledge about radon for its counties. Some States have been able to conduct radon investigations in areas smaller than geologic provinces and counties, so it is important to consult locally available data.

Although the information provided in Part IV of this report -- the State chapter entitled "Preliminary Geologic Radon Potential Assessment of Tennessee" -- may appear to be quite specific, it cannot be applied to determine the radon levels of a neighborhood, housing tract, individual house, etc. **THE ONLY WAY TO DETERMINE IF A HOUSE HAS ELEVATED INDOOR RADON IS TO TEST.** Contact the Region 4 EPA office or the Tennessee radon program for information on testing and fixing homes. Telephone numbers and addresses can be found in Part II of this report.

TENNESSEE - EPA Map of Radon Zones

The purpose of this map is to assist National, State and local organizations to target their resources and to implement radon-resistant building codes.

This map is not intended to determine if a home in a given zone should be tested for radon. Homes with elevated levels of radon have been found in all three zones. ***All homes should be tested, regardless of zone designation.***



Zone 1 Zone 2 Zone 3

IMPORTANT: Consult the publication entitled "Preliminary Geologic Radon Potential Assessment of Tennessee" before using this map. This document contains information on radon potential variations within counties. EPA also recommends that this map be supplemented with any available local data in order to further understand and predict the radon potential of a specific area.